

*Presentation at GABLS4 workshop, Toulouse, 12 September 2018*

# Background and Overview of the GEWEX Atmospheric Boundary Layer Study (GABLS)

Bert Holtslag

(thanks to Gert-Jan Steeneveld and GABLS community)



WAGENINGEN UNIVERSITY  
WAGENINGEN UR

Meteorology and Air Quality

# STABLE ATMOSPHERIC BOUNDARY LAYERS AND DIURNAL CYCLES

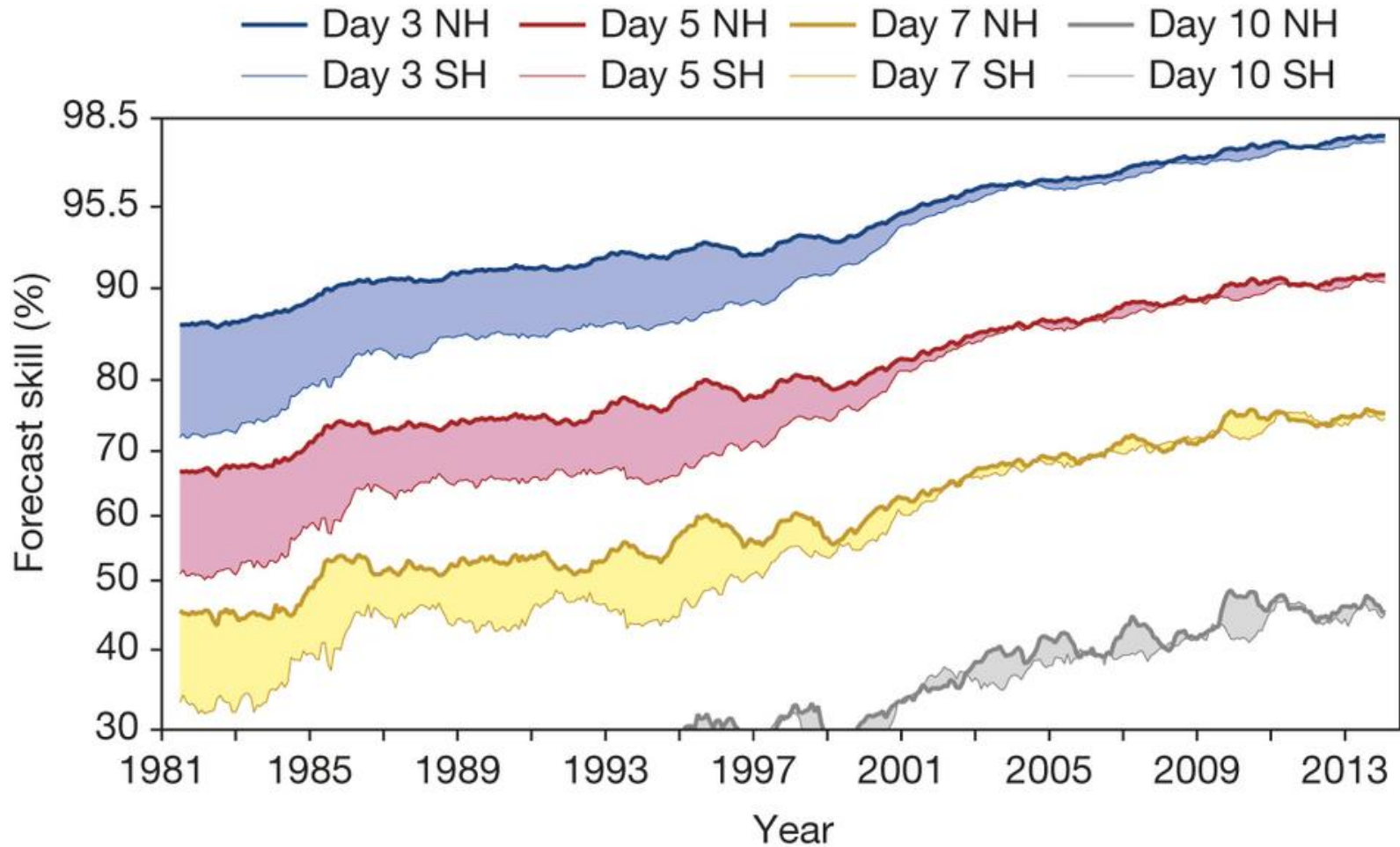
Challenges for Weather and Climate Models

BY A. A. M. HOLTSLAG, G. SVENSSON, P. BAAS, S. BASU, B. BEARE, A. C. M. BELJAARS, F. C. BOSVELD,  
J. CUXART, J. LINDVALL, G. J. STEENEVELD, M. TJERNSTRÖM, AND B. J. H. VAN DE WIEL

The atmospheric boundary layer impacts strongly the model performance for temperature and wind, yet stable situations, such as in clear, calm conditions at night or over ice, remain problematic.

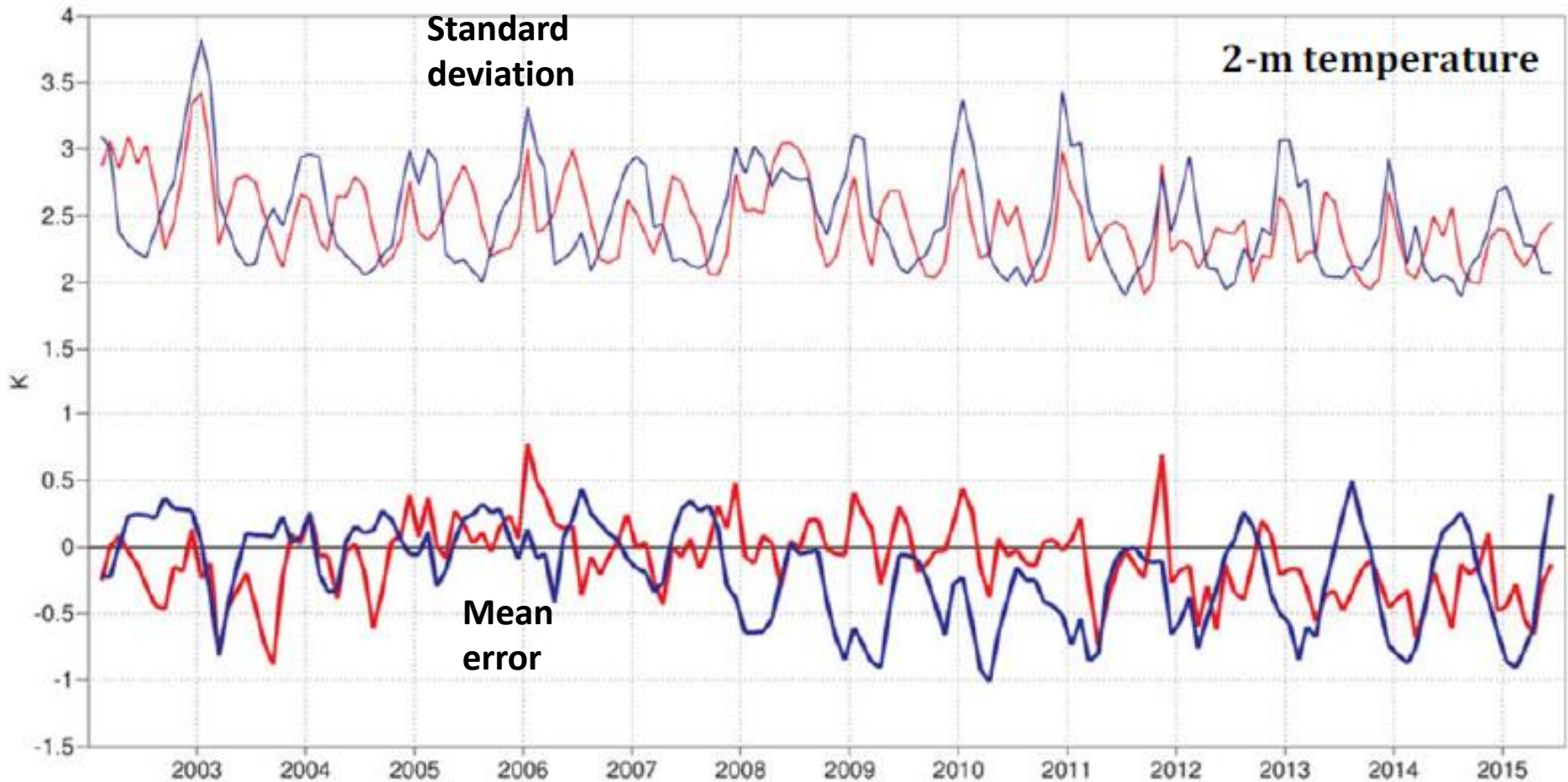


Improvement of forecast skill during the years  
as the correlation between the forecasts and the verifying analysis of the height  
of the 500-hPa level, expressed as the anomaly with respect to the  
climatological height. Values greater than 60% indicate useful forecasts, while  
those greater than 80% represent a high degree of accuracy

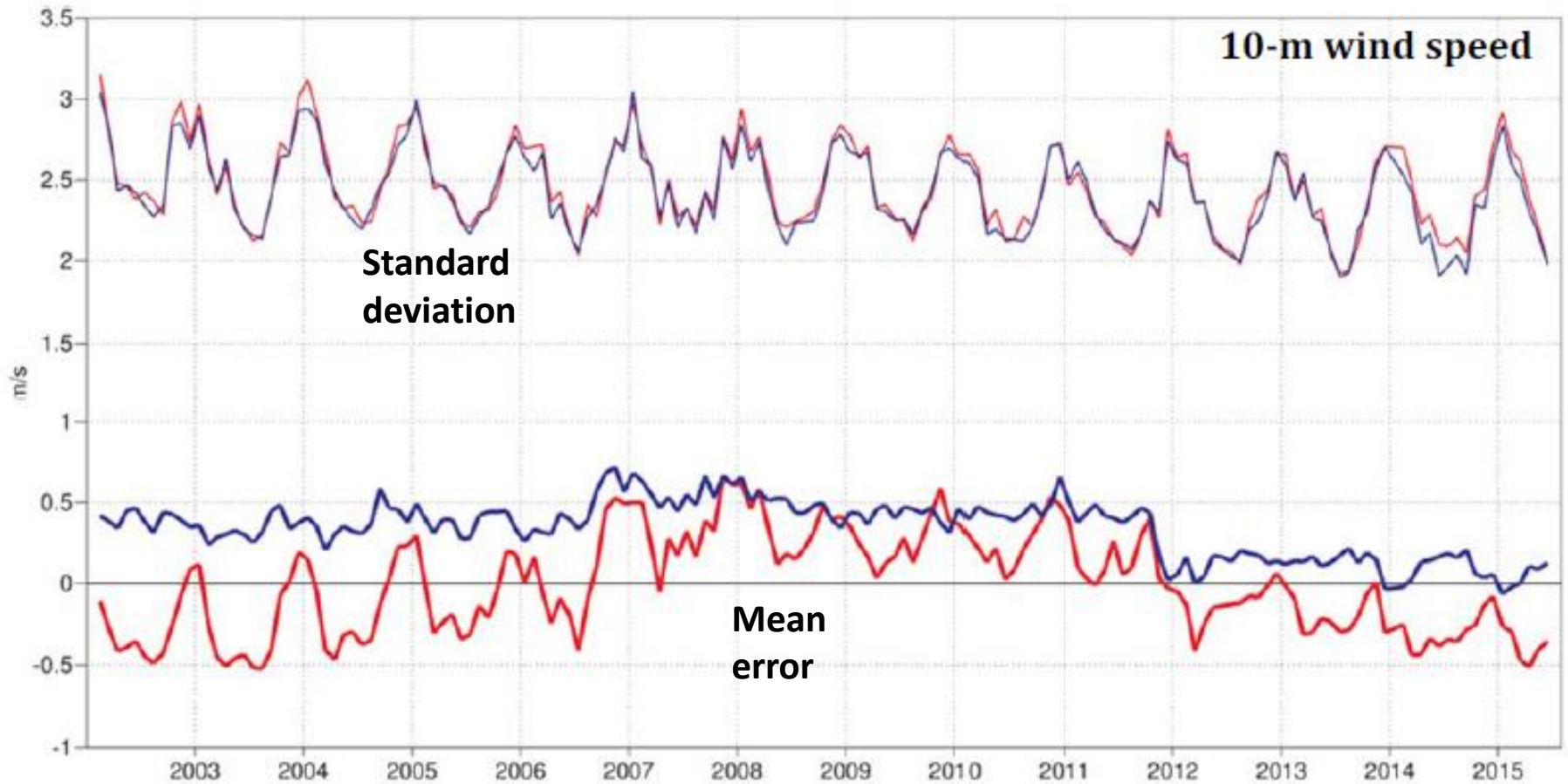


(Bauer et al., 2015)

# Still significant day (+72 h) and night (+60h) temperature errors in weather forecasts (Monthly averages over Europe by ECMWF)

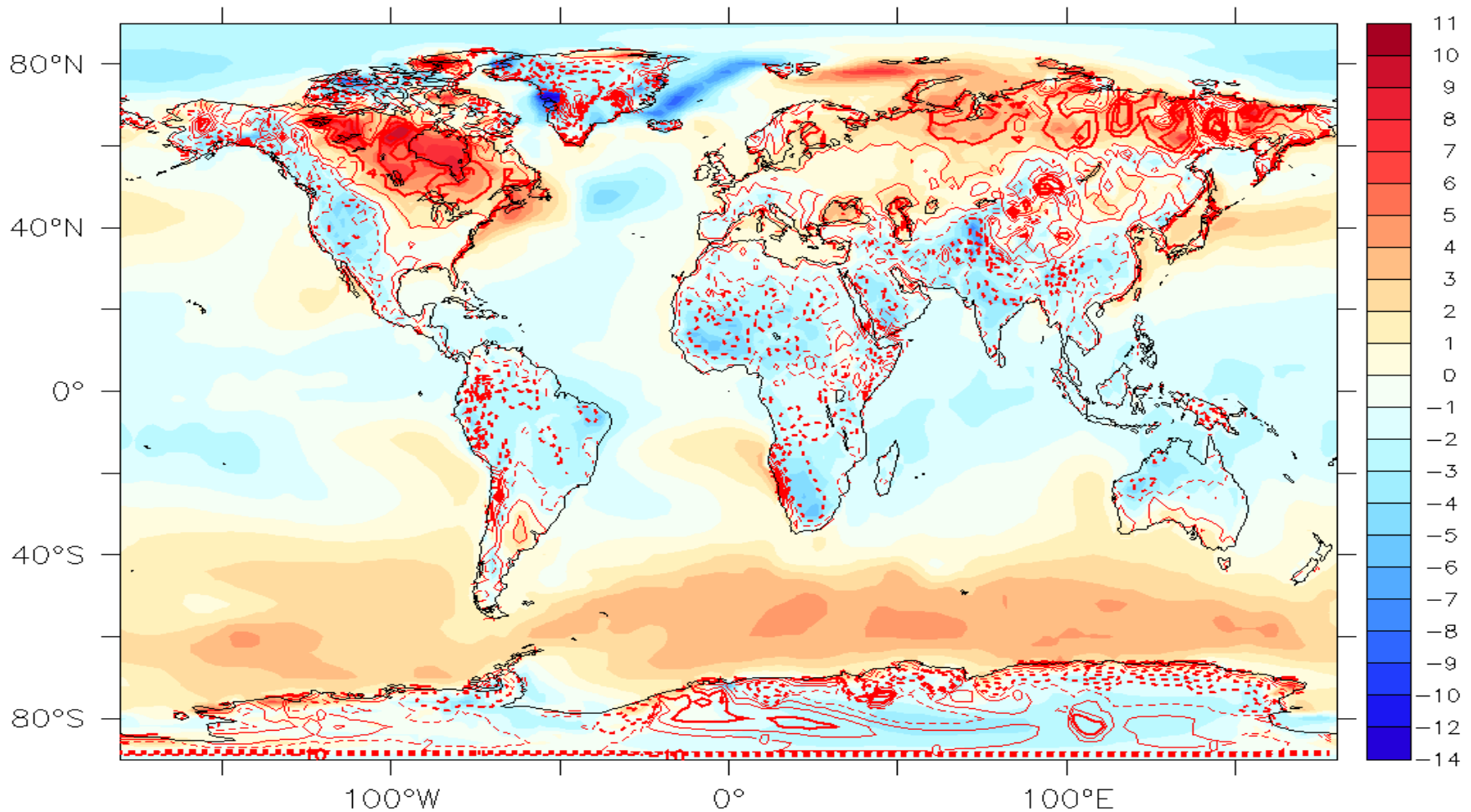


# Still significant day (+72 h) and night (+60h) wind errors in weather forecasts (Monthly averages over Europe by ECMWF)

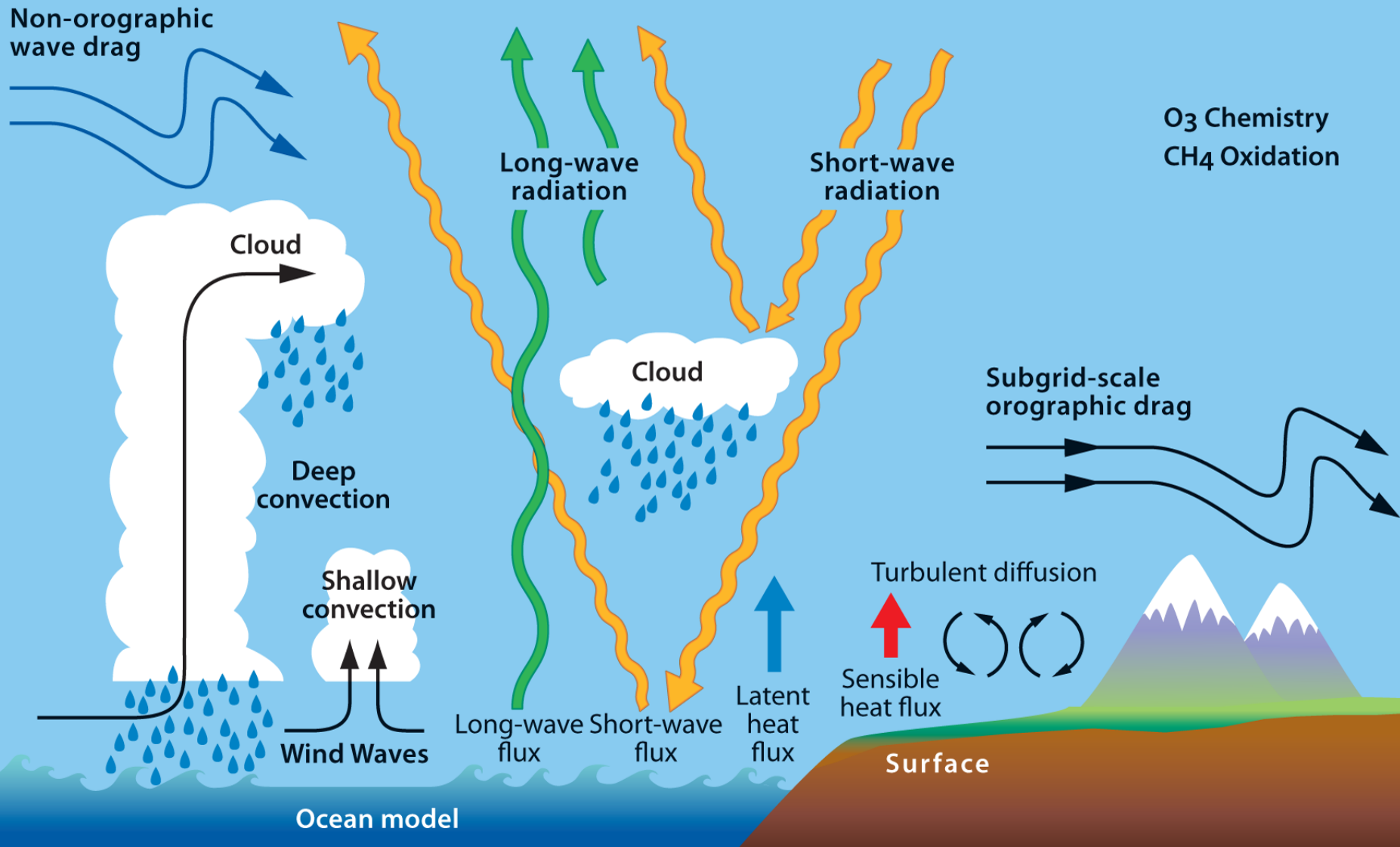


# Bias of 2 m temperature in DJF

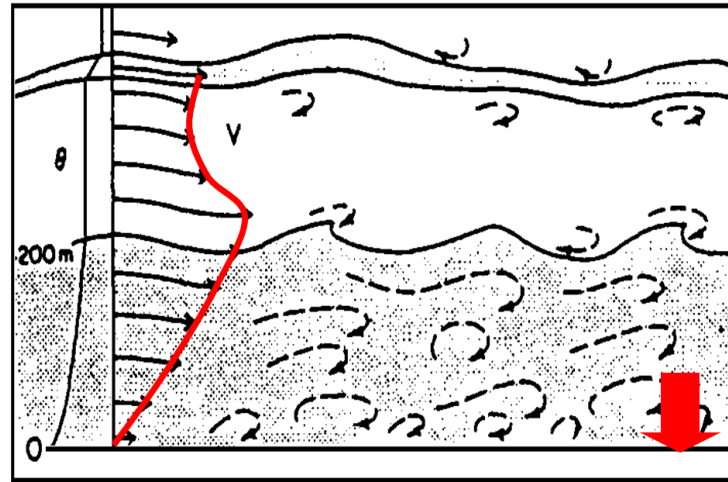
EC-Earth climate model – ERA-Interim (Jan 1989-Jan 2009)



# Small-scale processes relevant for weather and climate



# Stable Atmospheric Boundary Layers



(Figure by  
John Wyngaard, 1985)

## Long history of research and applications

(e.g., Zilitinkevich, 1972; Brost and Wyngaard, 1978; Nieuwstadt, 1984,...)

Progress in understanding through observations, simulations and modelling, in particular for cases with stronger turbulence

Stable boundary layers with weak and intermittent turbulence still very challenging (e.g., vd Wiel et al., 2002; Mahrt, 2014,...)



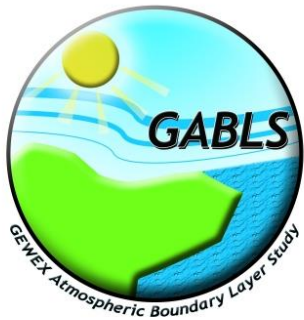


# Understanding and Modeling Atmospheric Boundary Layers: It is still a challenge!

Atmospheric models do have difficulties in representing  
the stable boundary layer in particular

Sensitivity to details in turbulence mixing formulations  
and interaction with other small scale processes

## Strategy




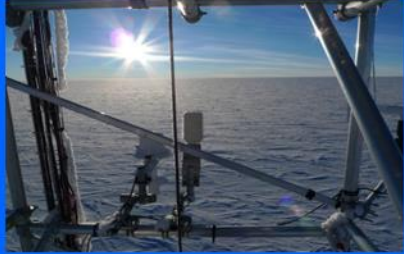


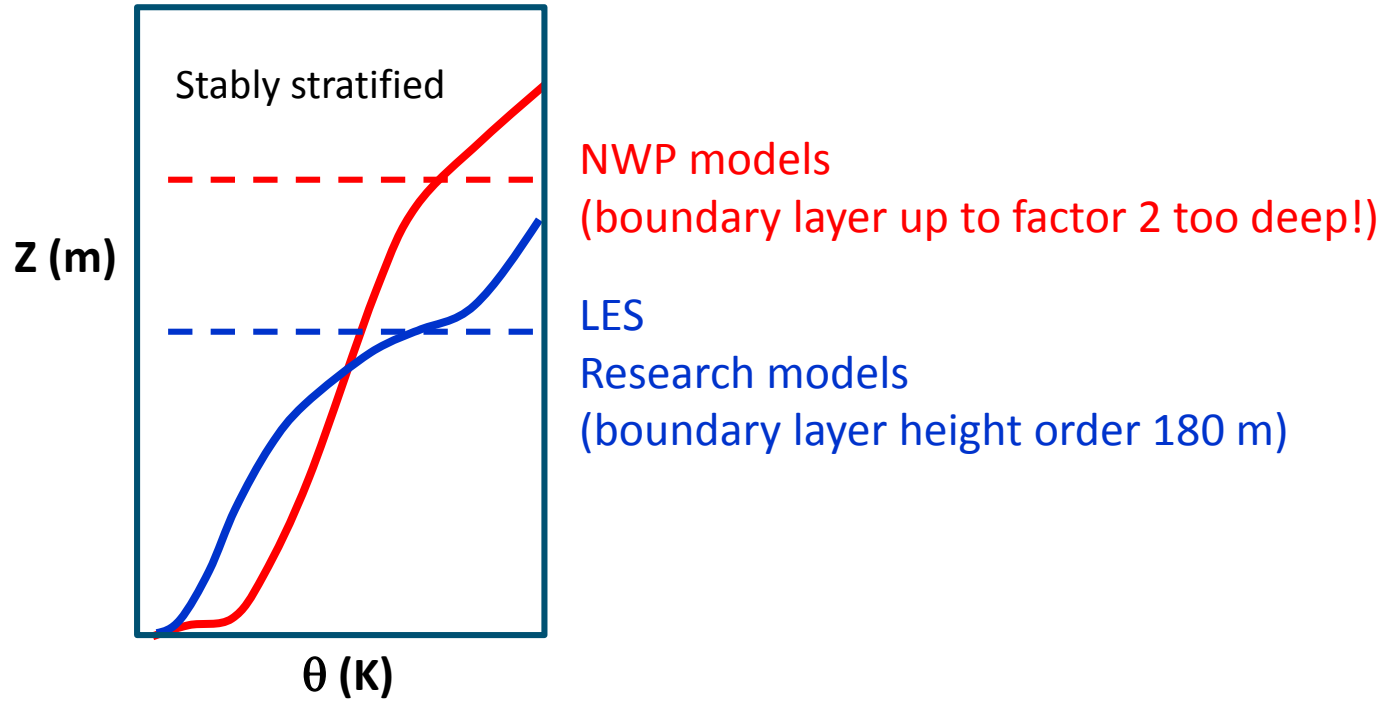
Enhance understanding by benchmark studies  
over land and ice in comparison with observations  
and fine scale numerical model results





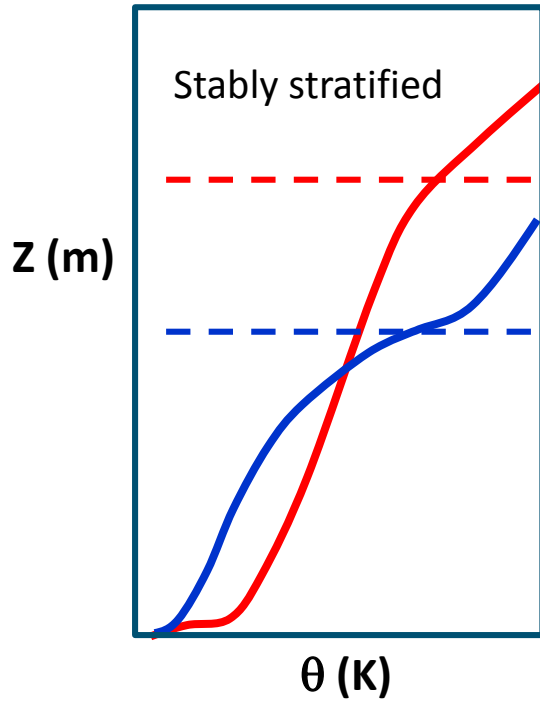
# GEWEX Atmospheric Boundary Layer Study (GABLS) provides platform for model intercomparisons

| GABLS1   | GABLS2  | GABLS3   | GABLS4  |
|--|---|--|---|
|  |  |  |  |
| <i>LES</i> as reference  | Data (CASES99)  | Data (CABA UW)   | Data (DOME-C)   |
| Academic set up  | Idealized forcings  | Realistic forcings   | Realistic forcings  |
| Prescribed $T_s$   | Prescribed $T_s$  | Full coupling ( <i>SCM</i> )<br>Prescribed $T_s$ ( <i>LES</i> )                    | Full coupling<br>Prescribed $T_s$   |
| No Radiation   | No Radiation  | Radiation included   | Radiation included  |
| Turbulent mixing   | Diurnal cycle   | Low level jet +<br>transitions   | Very Stable   |



(Beare et al. 2006; Cuxart et al. 2006;  
Svensson and Holtslag, 2009)



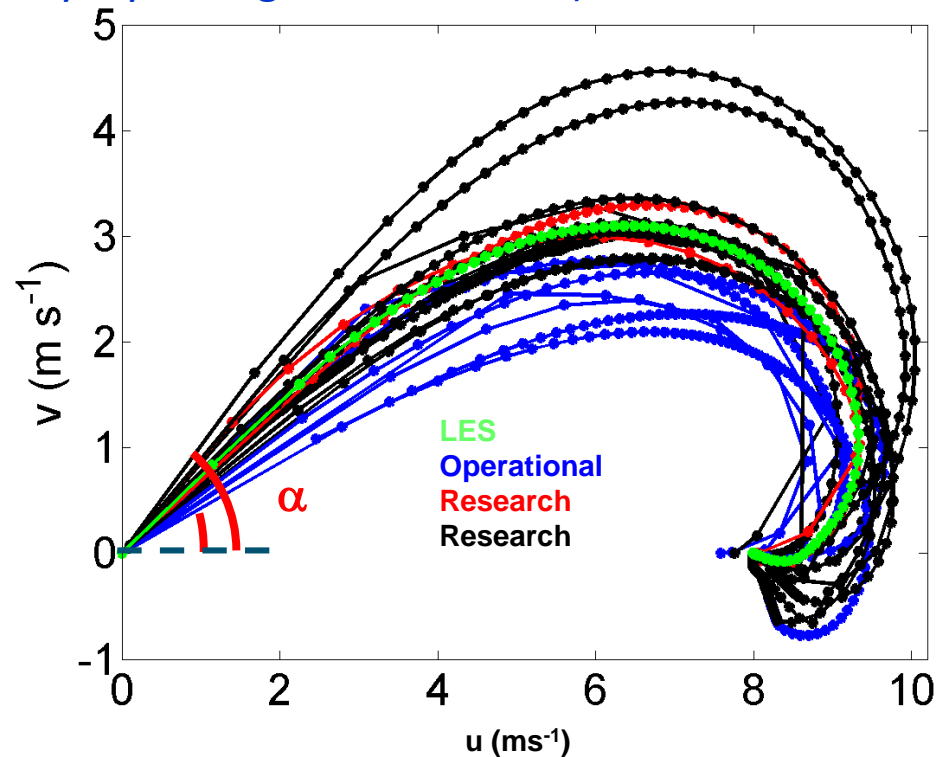


NWP models  
(boundary layer up to factor 2 too deep!)

LES  
Research models  
(boundary layer height order 180 m)

Too small variation of wind direction with height in operational NWP models

(Beare et al. 2006; Cuxart et al. 2006; Svensson and Holtslag, 2009)

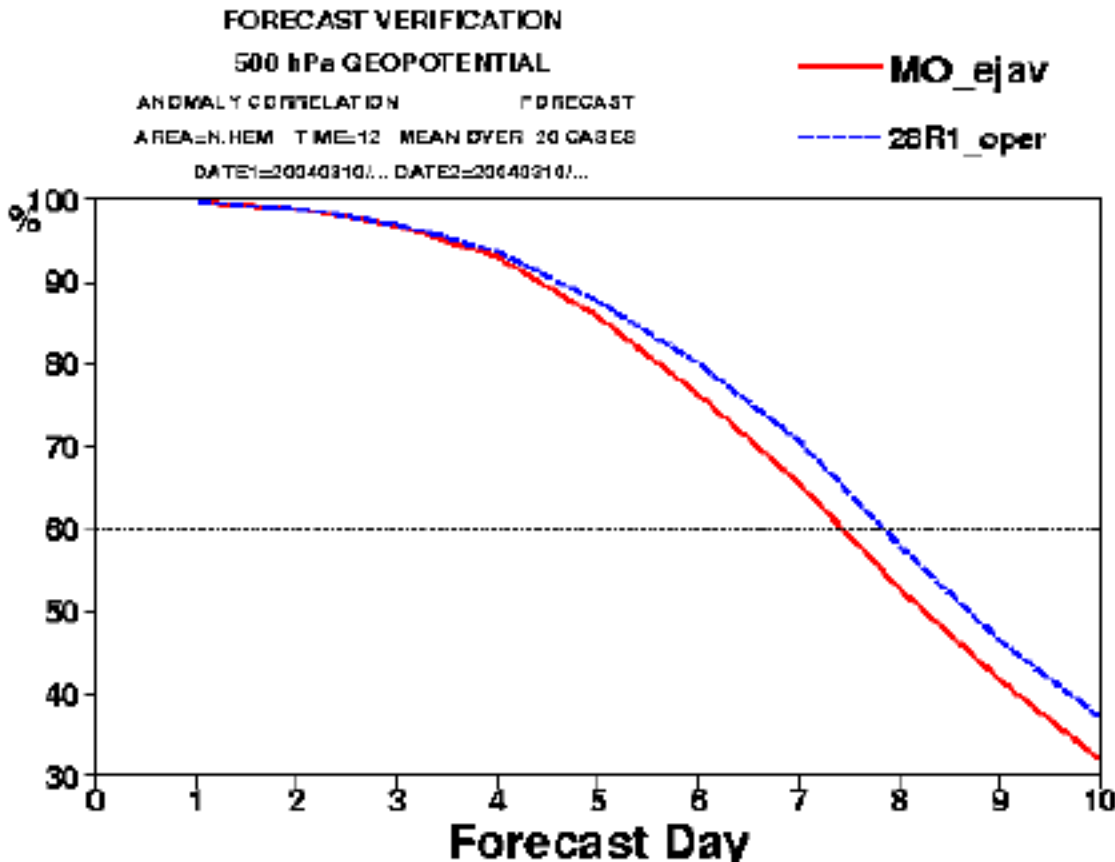


# Stable boundary layer diffusion affects large scale scores

Effect of **MO-stability functions** (reduced diffusion) instead of **operational formulation**, on 500hPa NH height scores

Model somehow needs larger drag over land than can be obtained from schemes that produce reasonable stable boundary layer structure.

Ground truth for drag over land does not exist.



Courtesy

Anton Beljaars



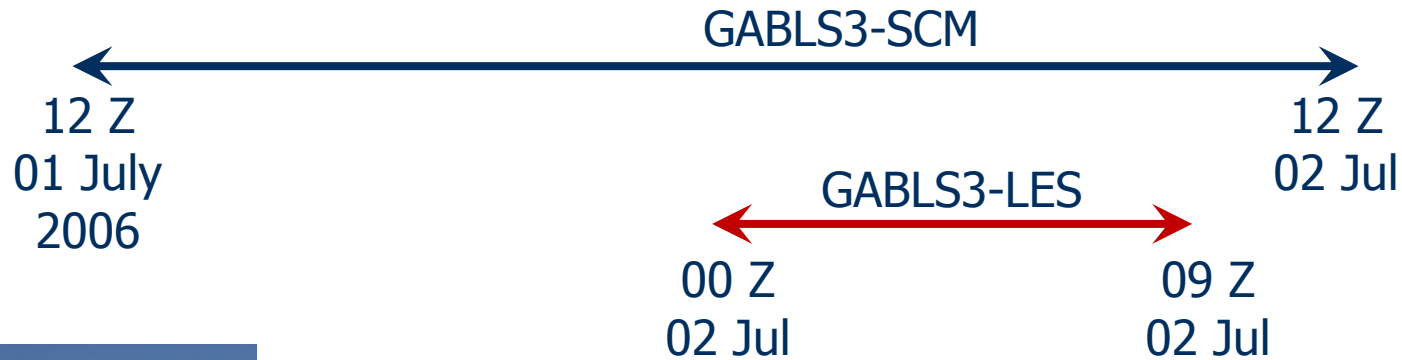


*Can small-scale orographic gravity waves provide the missing drag  
in the stable boundary layer?*

Aristofanis Tsiringakis, Gert-Jan Steeneveld,  
and Bert Holtslag  
Meteorology and Air Quality  
Wageningen University, NL

See paper in QJRMS, 2017

# GABLS3: SCM and LES model studies



Cabauw tower  
(KNMI, NL)

Initialization: Cabauw tower, Profiler, Nearby De Bilt Sounding

Geostrophic Wind (time-height dependent)

Large-scale Advection (time-height dependent)

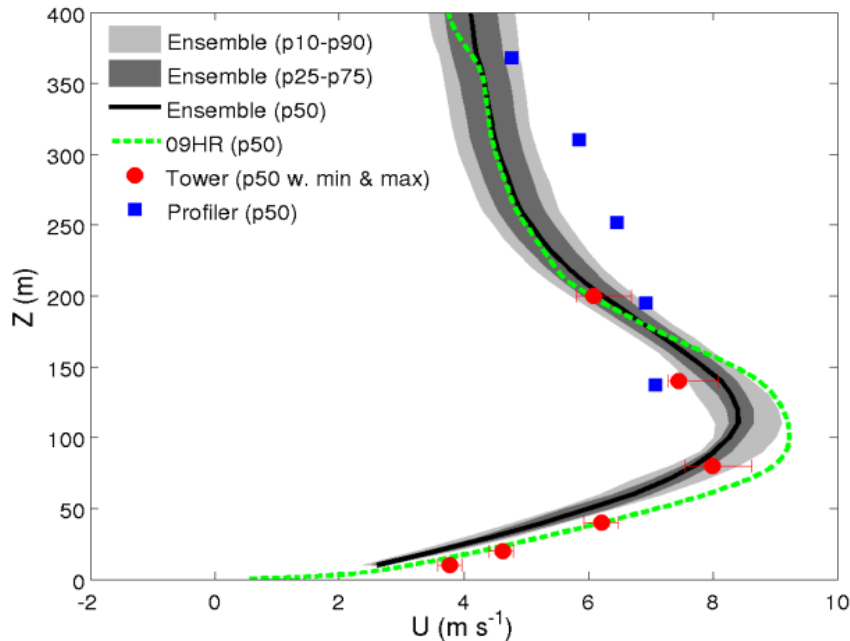
LES: Prescribed Observed Near Surface Temperature (0.5 m)

SCM: Full interaction with Land Scheme and Radiation

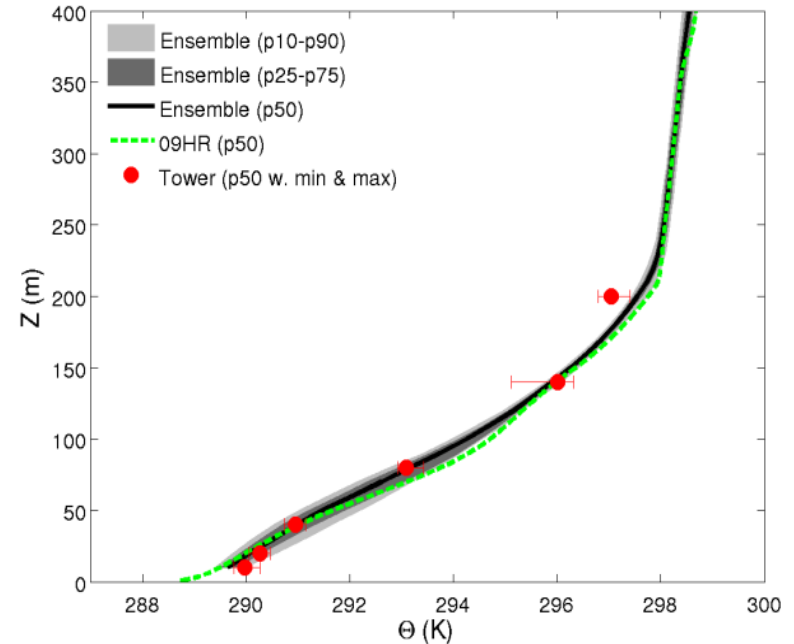
# GABLS3 Large Eddy Simulation intercomparison

(coordinated by Sukanta Basu)

## Wind magnitude (m/s)



## Potential Temp (K)



Ensemble of 11 LES models for mean profiles 03-04 UTC (6.25 m resolution)

Green dashed line: 1m vertical resolution LES run (Courtesy Siegfried Raasch)

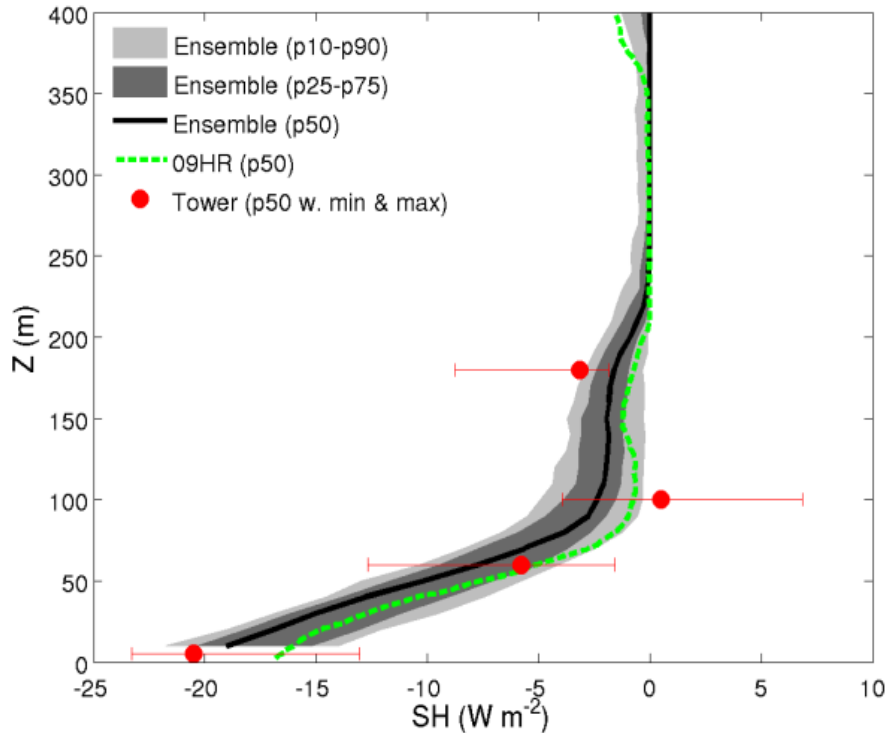
Red dots: Tower observations; Blue squares: Wind Profiler



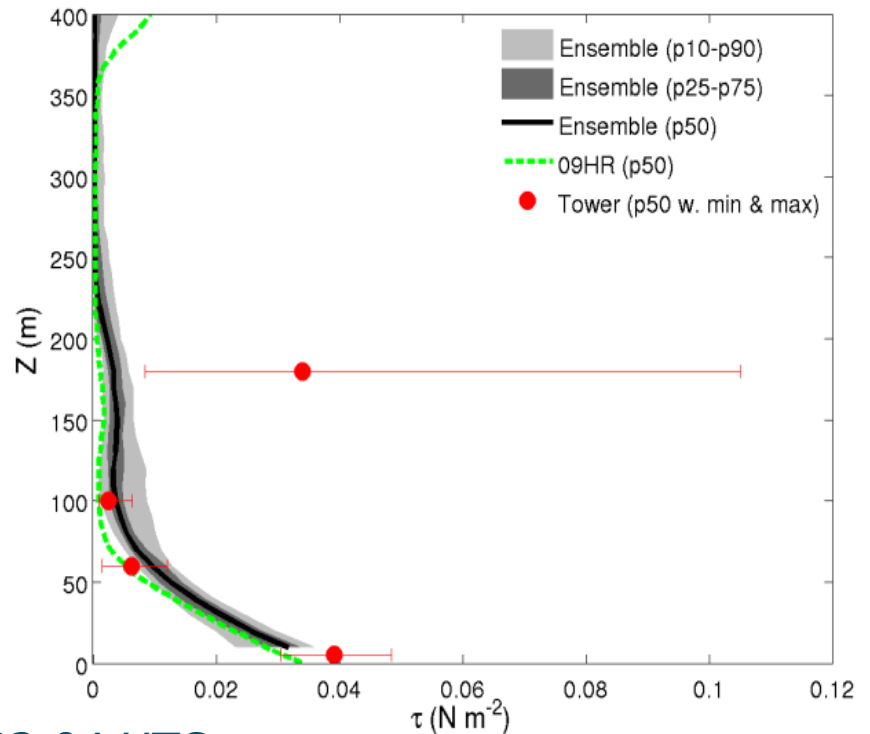
# GABLS3 Large Eddy Simulation intercomparison

(coordinated by Sukanta Basu)

## Sensible Heat flux ( $\text{W}/\text{m}^2$ )



## Momentum flux ( $\text{N}/\text{m}^2$ )



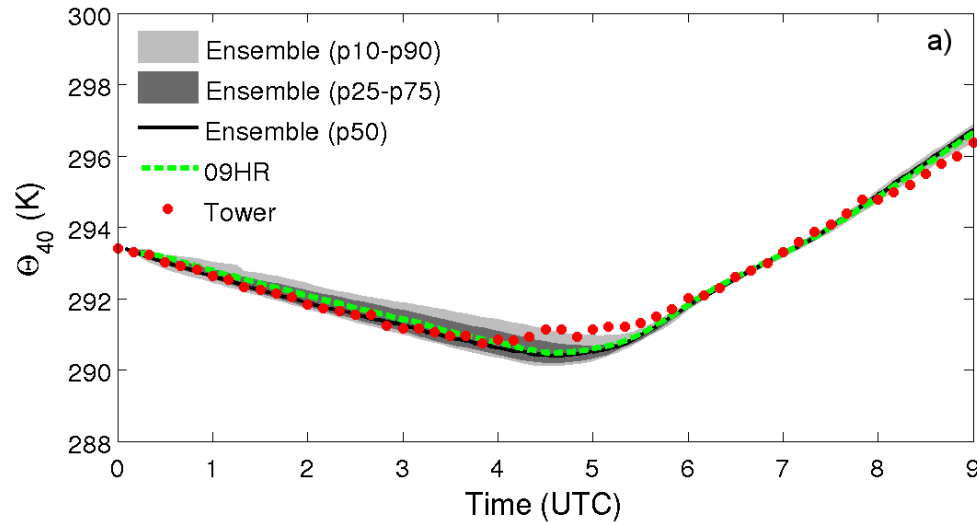
Flux profiles 03-04 UTC

Green dashed line: 1m LES run (Courtesy Siegfried Raasch)

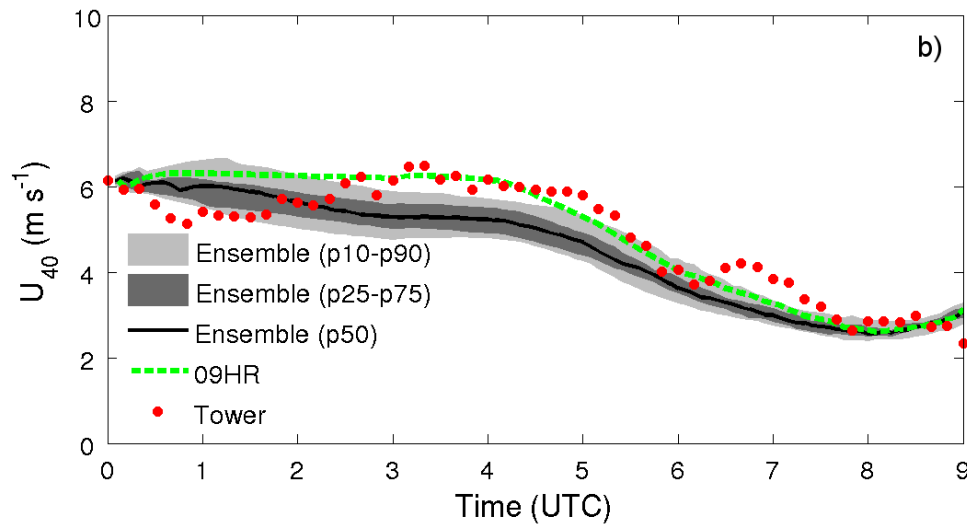
Red dots: Cabauw Tower observations



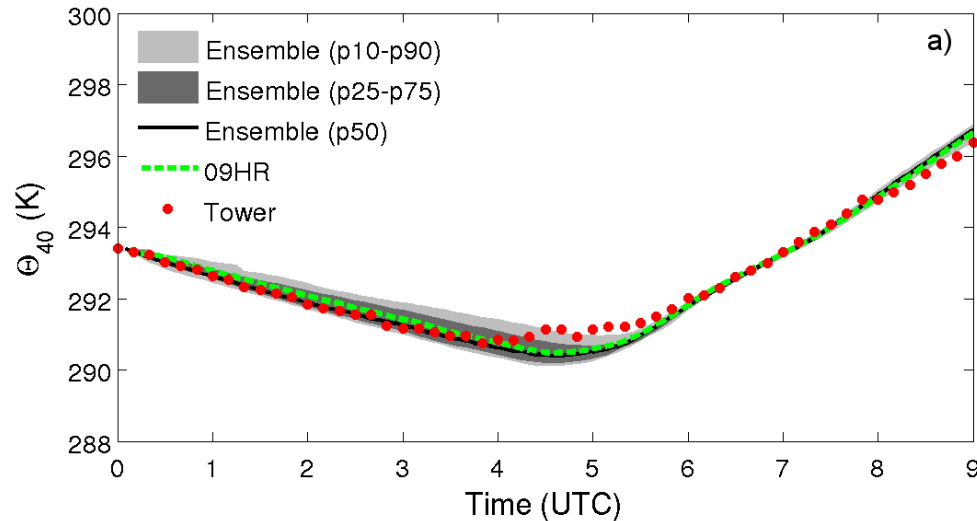
# GABLS3 Large Eddy Simulation intercomparison



Temporal evolution



# GABLS3 Large Eddy Simulation intercomparison



Temporal  
evolution

Overall LES results in good agreement  
with Cabauw mean and turbulent flux profiles  
(using observed near surface temperatures)

*But there are issues...*



# GABLS3

intercomparison of 19  
Single Column versions  
(SCM) of operational  
and research models  
(Coordinated by  
Fred Bosveld, KNMI)

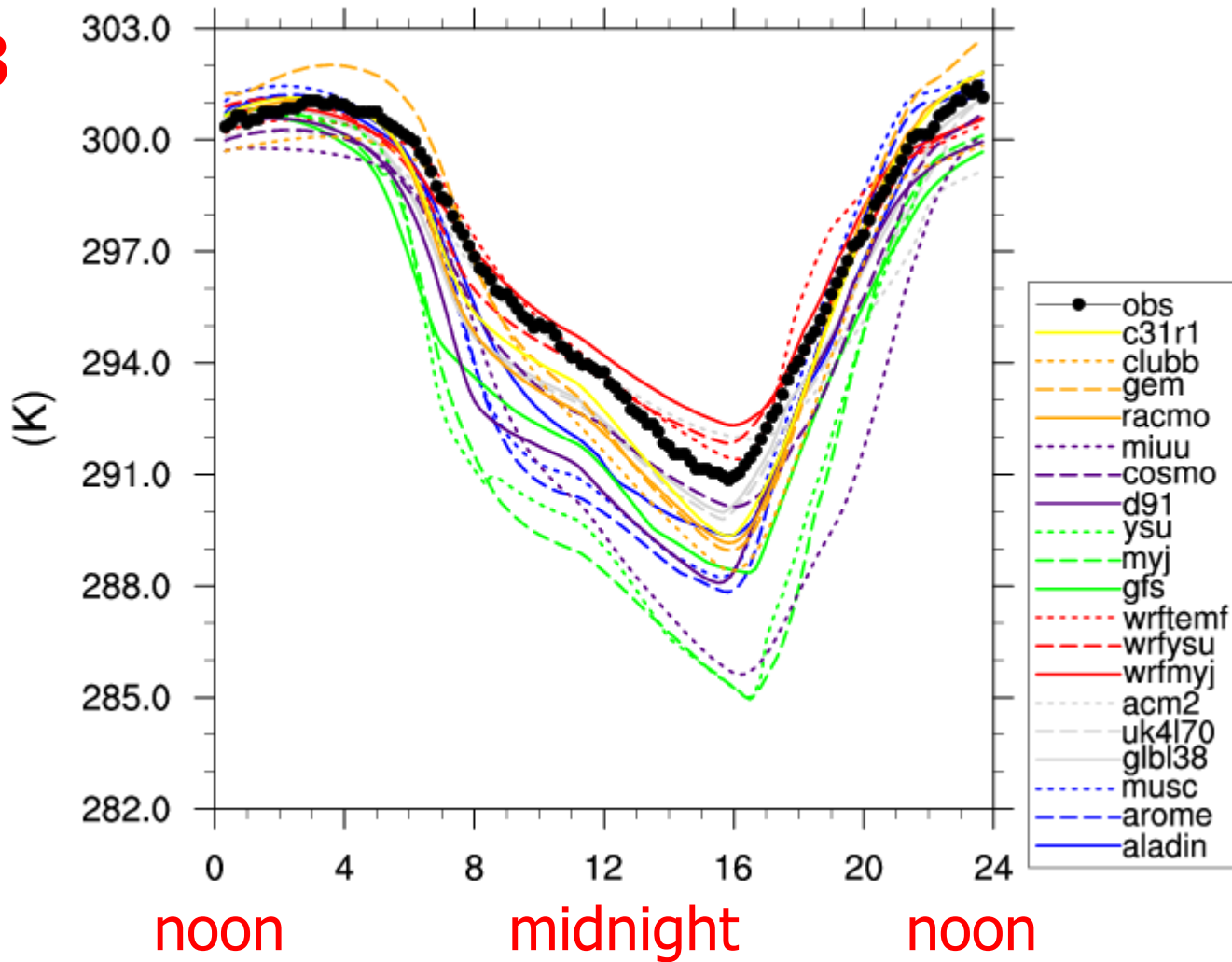
Note:

Each SCM uses its  
own radiation and land  
surface scheme  
interacting with the  
boundary layer scheme on  
usual resolution!  
(Nlev is number of vertical  
levels in whole atmosphere)

| <i>Name</i> | <i>Institute</i> | <i>Nlev</i> | <i>BL.Scheme</i> | <i>Skin</i> |
|-------------|------------------|-------------|------------------|-------------|
| ALADIN      | Meteo France     | 41          | TKE              | No          |
| AROME       | Meteo France     | 41          | TKE              | No          |
| GLBL38      | Met Office       | 38          | K (long tail)    | Yes         |
| UK4L70      | Met Office       | 70          | K (short tail)   | Yes         |
| D91         | WUR              | 91          | K                | Yes         |
| GEM         | Env. Canada      | 89          | TKE-I            | No          |
| ACM2        | NOAA             | 155         | K+non-local      | No          |
| WRF YSU     | NOAA             | 61          | K                | No          |
| WRF MYJ     | NOAA             | 61          | TKE-I            | No          |
| WRFTMF      | NOAA             | 61          | Total E          | No          |
| COSMO       | DWD              | 41          |                  |             |
| GFS         | NCEP             | 57          | K                | Yes         |
| WRF MYJ     | NCEP             | 57          | TKE-I            | Yes         |
| WRF YSU     | NCEP             | 57          | K                | Yes         |
| MIUU        | MISU             | 65          | 2nd order        |             |
| MUSC        | KNMI             | 41          | TKE-I            | No          |
| RACMO       | KNMI             | 80          | TKE              | Yes         |
| C31R1       | ECMWF            | 80          | K                | Yes         |
| CLUBB       | UWM              | 250         | Higher order     | No          |

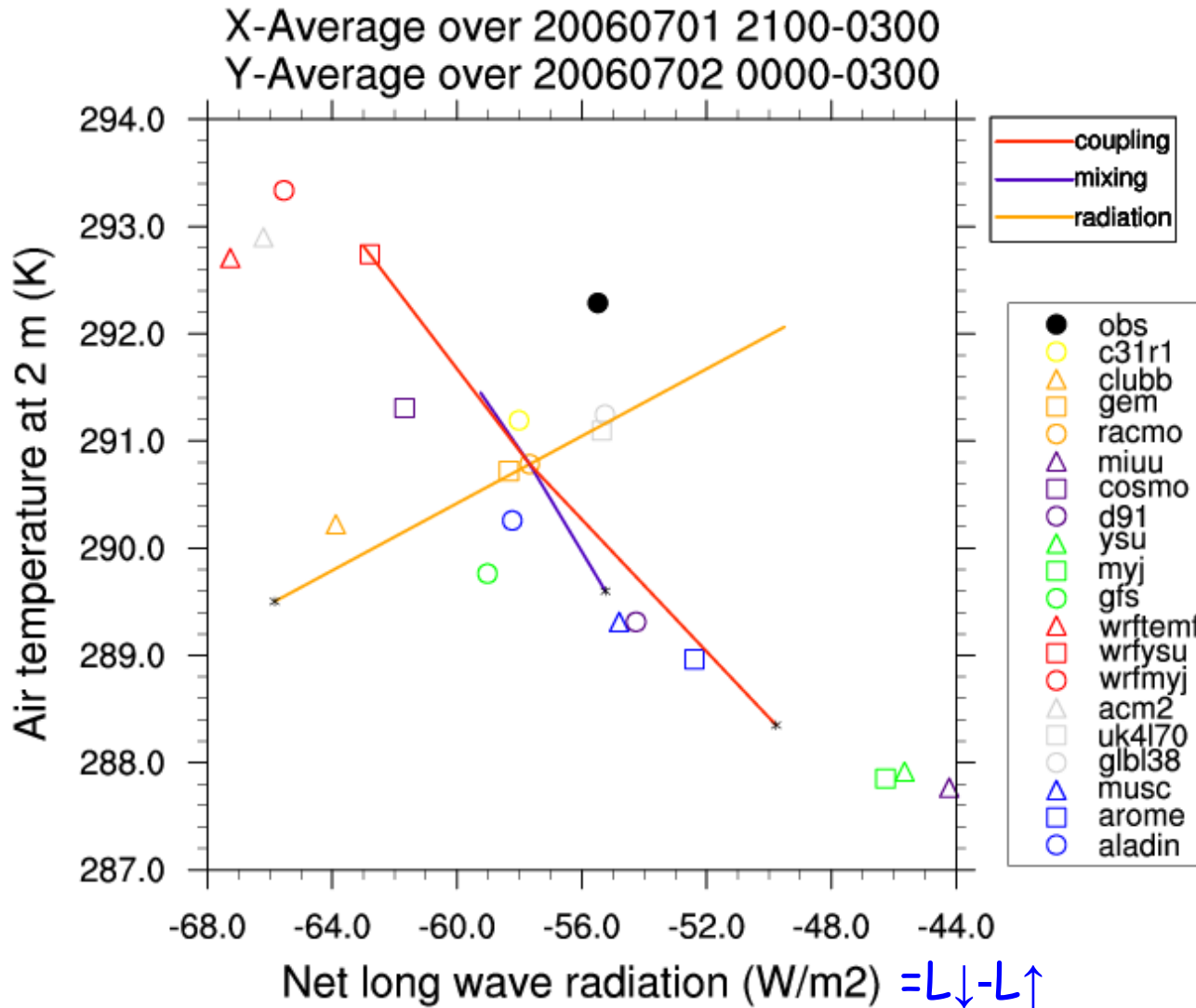
# Models show large range of results for 2 m temperature

GABLS3



(Bosveld et al., 2014)

# Process diagram: Influence of surface radiation



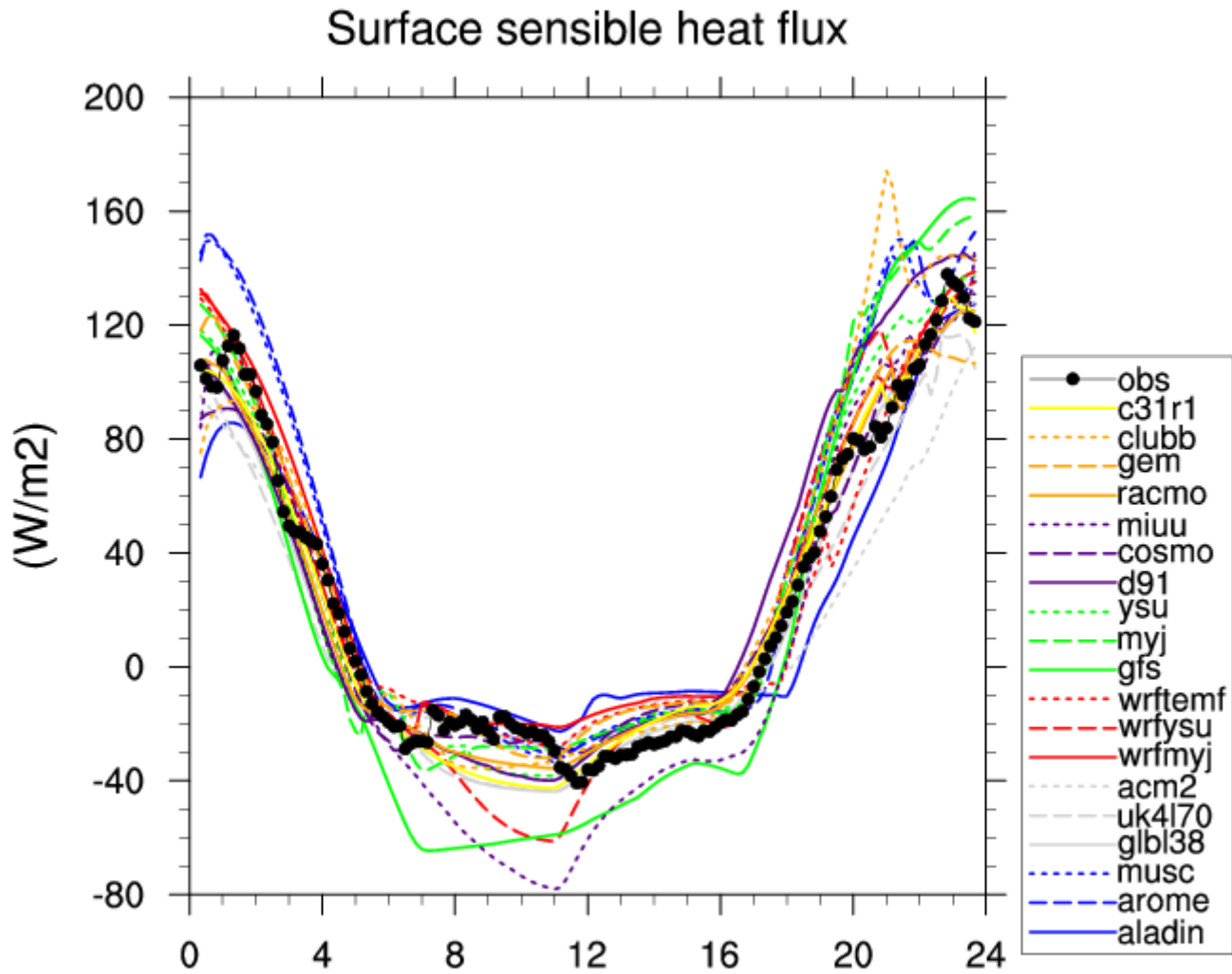
Note:  
 $L_{\uparrow}$  is strong  
function of  
surface  
temperature

(Bosveld et al, 2014)

Warmer surface

Colder Surface

GABLS3



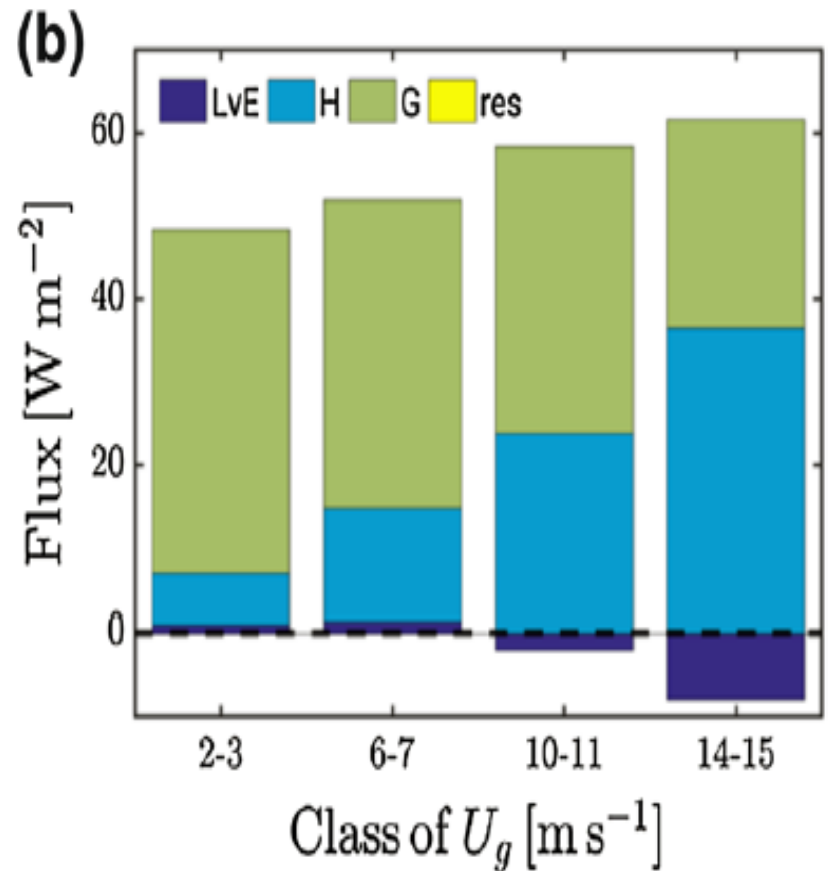
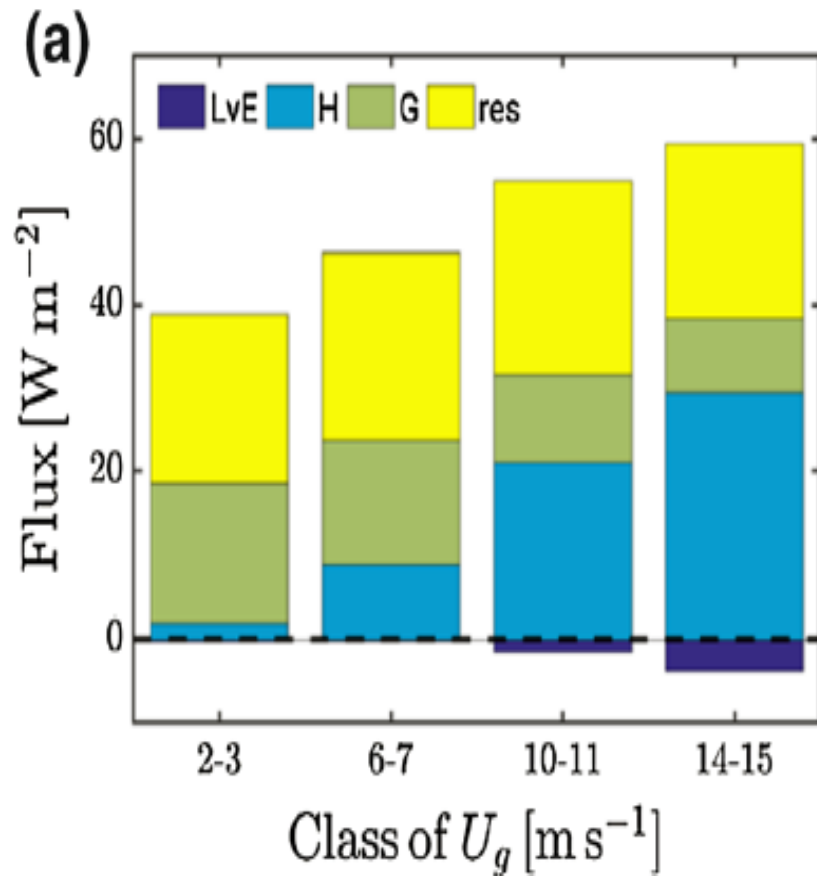
noon

midnight

noon

(Bosveld et al., 2014)

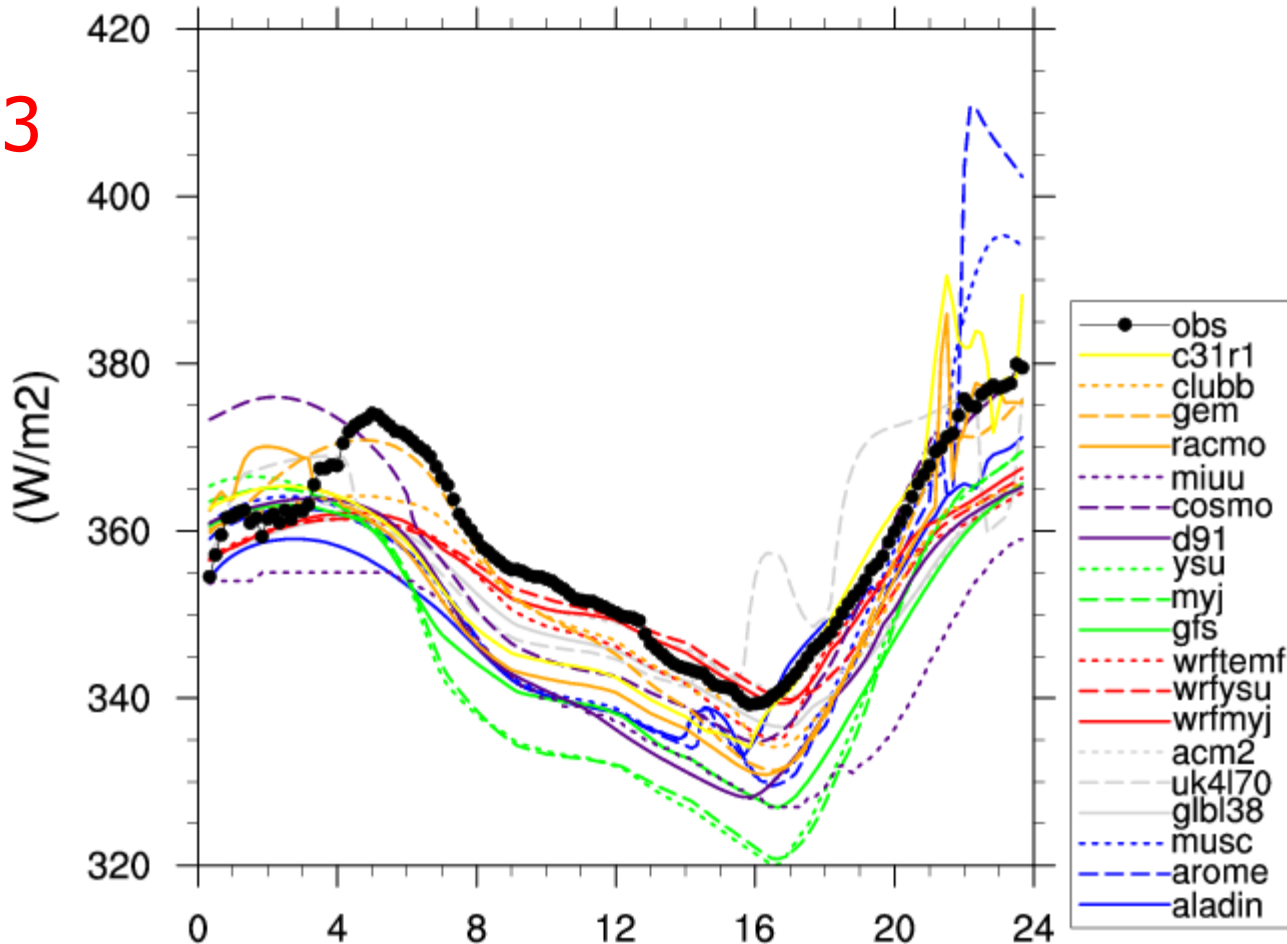
Observed (a) and modelled (b) components of the surface energy balance for four classes of geostrophic wind speed  $U_g$  using 11 years (2005–2015) of daily SCM simulations vs Cabauw observations (for  $Q_{net} < -30 \text{ W/m}^2$ ). The residual term is indicated by 'res' (after Baas et al., BLM, 2018)





# Models underestimate downward longwave radiation at surface (positive feedback with temperature)

GABLS3



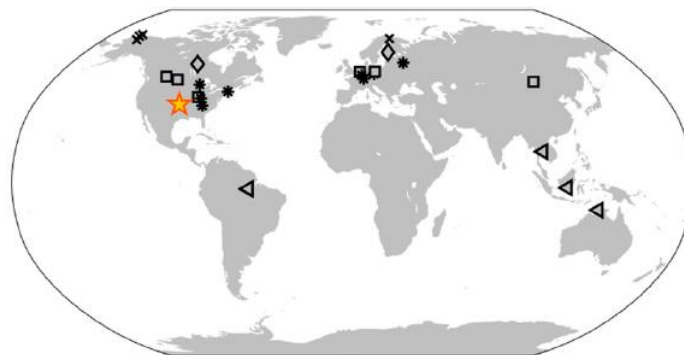
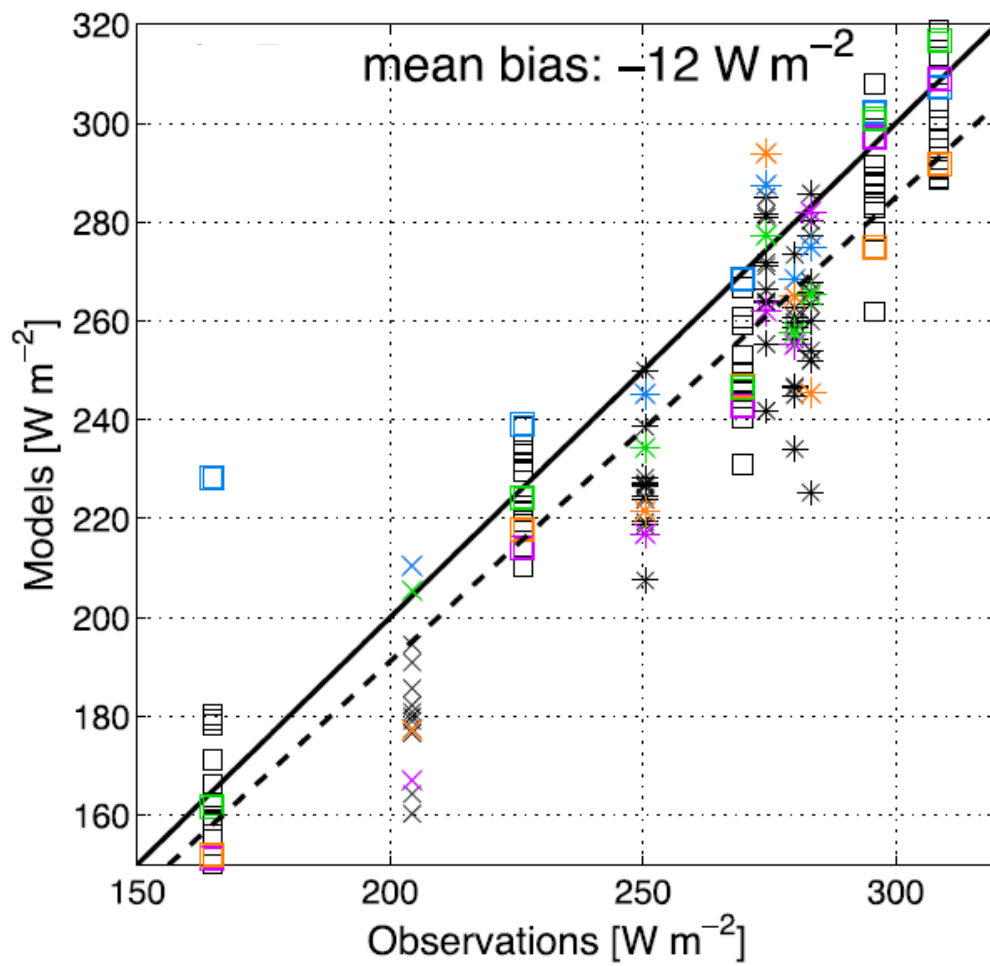
noon

midnight

noon

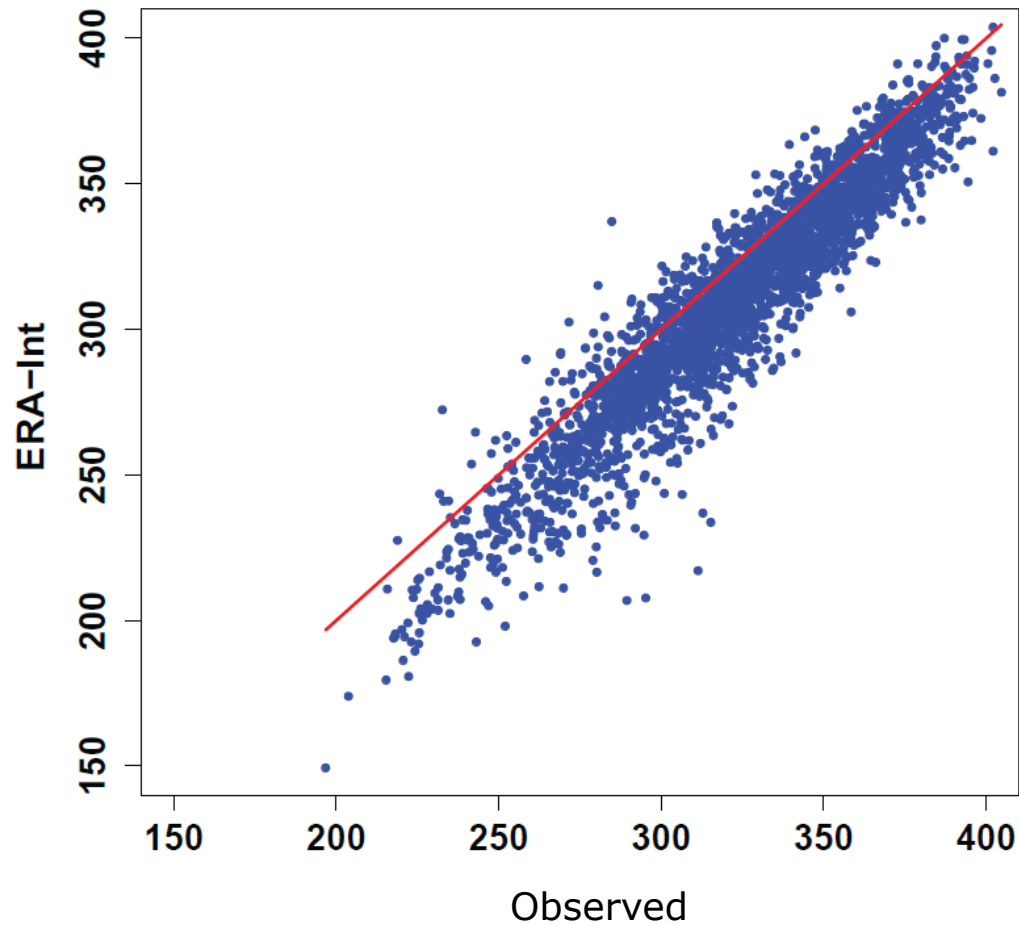
(Bosveld et al., 2014)

# Downward longwave radiation typically underestimated by climate models on various locations (10 year averages)



- \* BCC-CSM1.1
  - \* CESM1
  - \* HadGEM2-A
  - \* IPSL-CM5A-MR
  - \* Various models
- 
- Midlatitude Grassland
  - \* Midlatitude Forest
  - × Arctic Tundra/Wetland

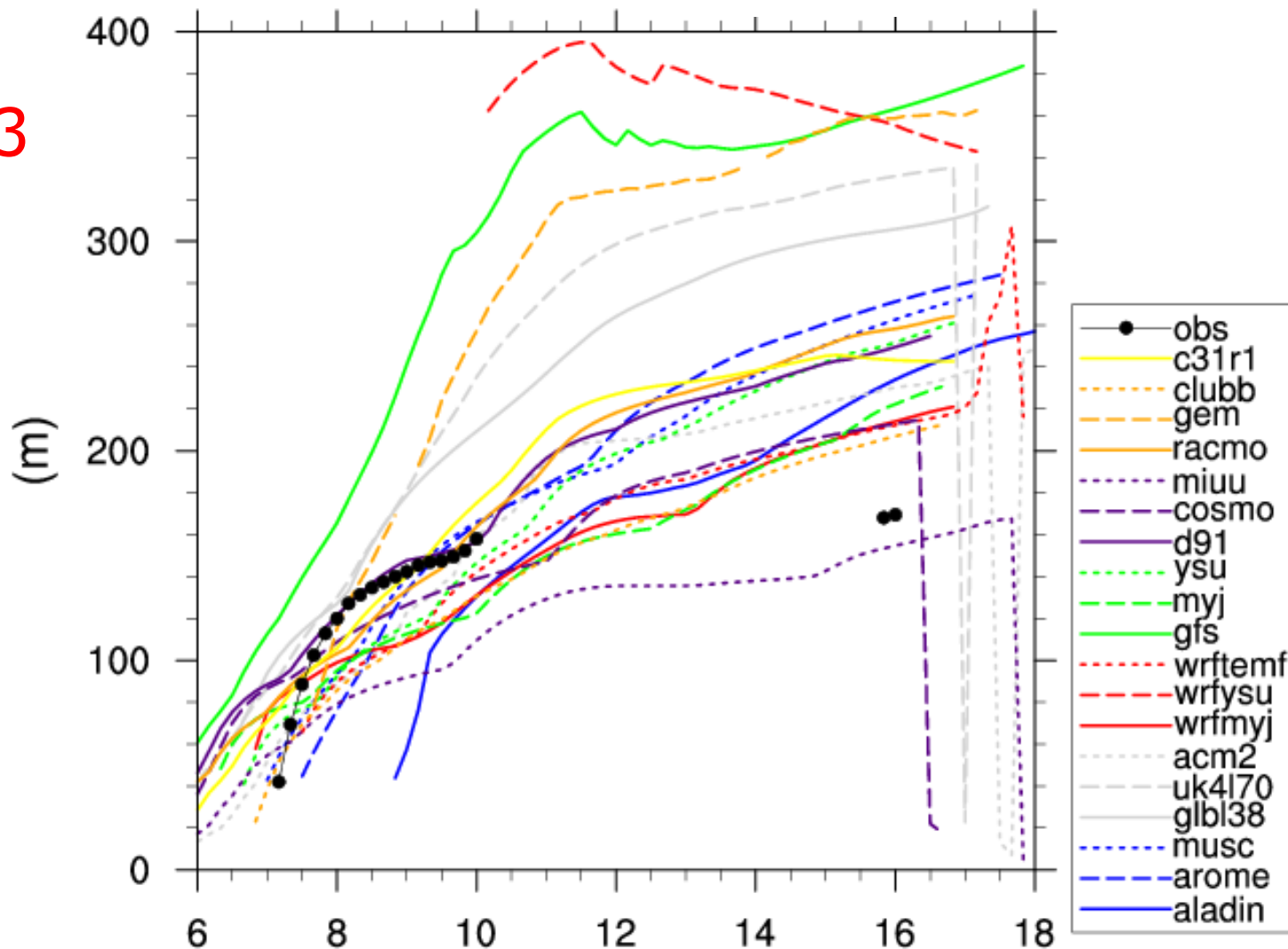
*Average values of incoming long-wave radiation measured in Wageningen versus ERA-Interim (period 2001-2012)*



Courtesy Hartogensis and Anton Beljaars (De Bruin, 2016)

# Models show typically too deep boundary layers

GABLS3



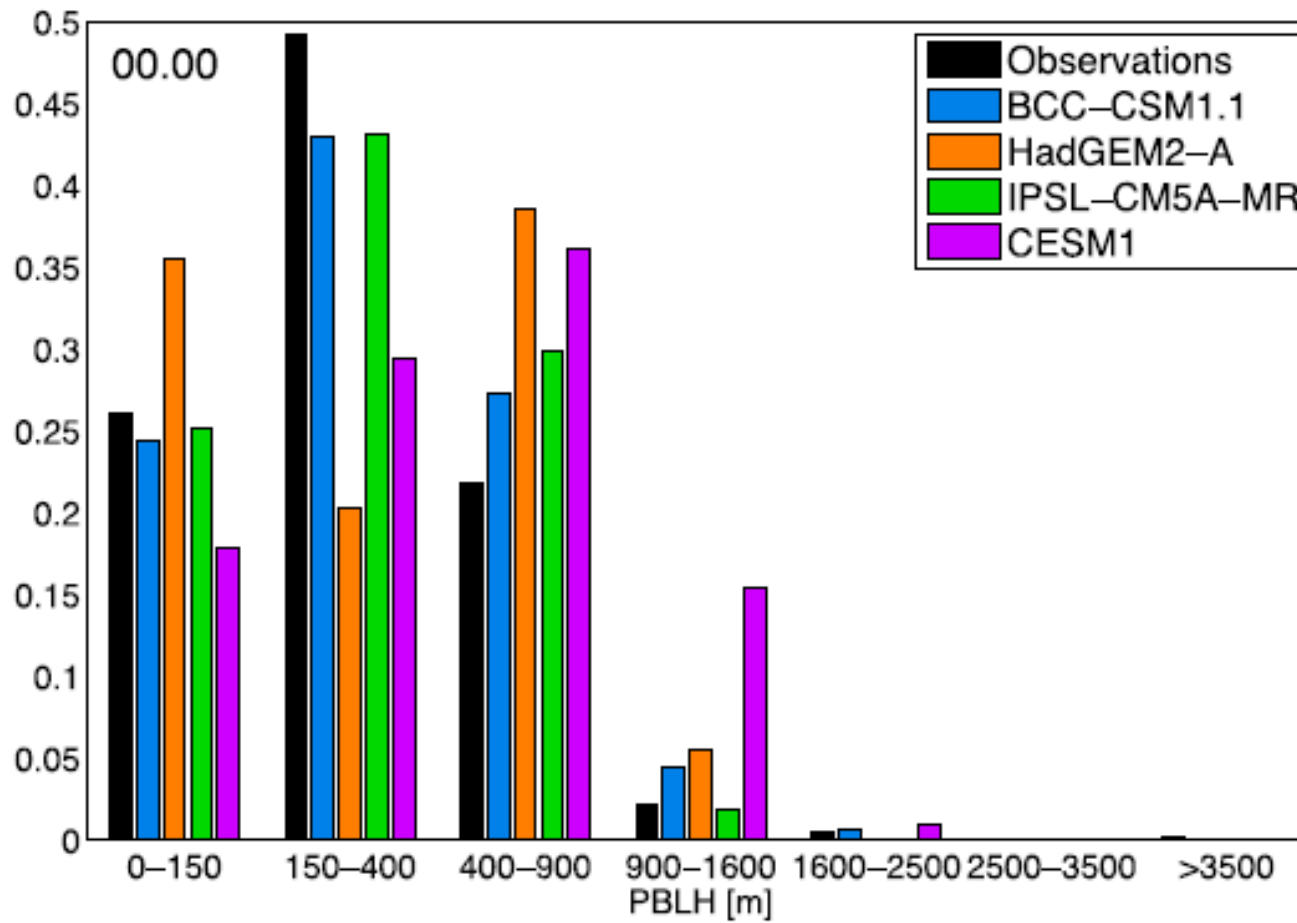
evening

midnight

morning

(Bosveld et al, 2014)

Also deeper stable boundary layer depths in climate models  
(10 year averages for ARM Southern Great Plains site, USA)



# Why so difficult to represent stable boundary layers?

Many small scale processes and non-linear interactions not fully understood

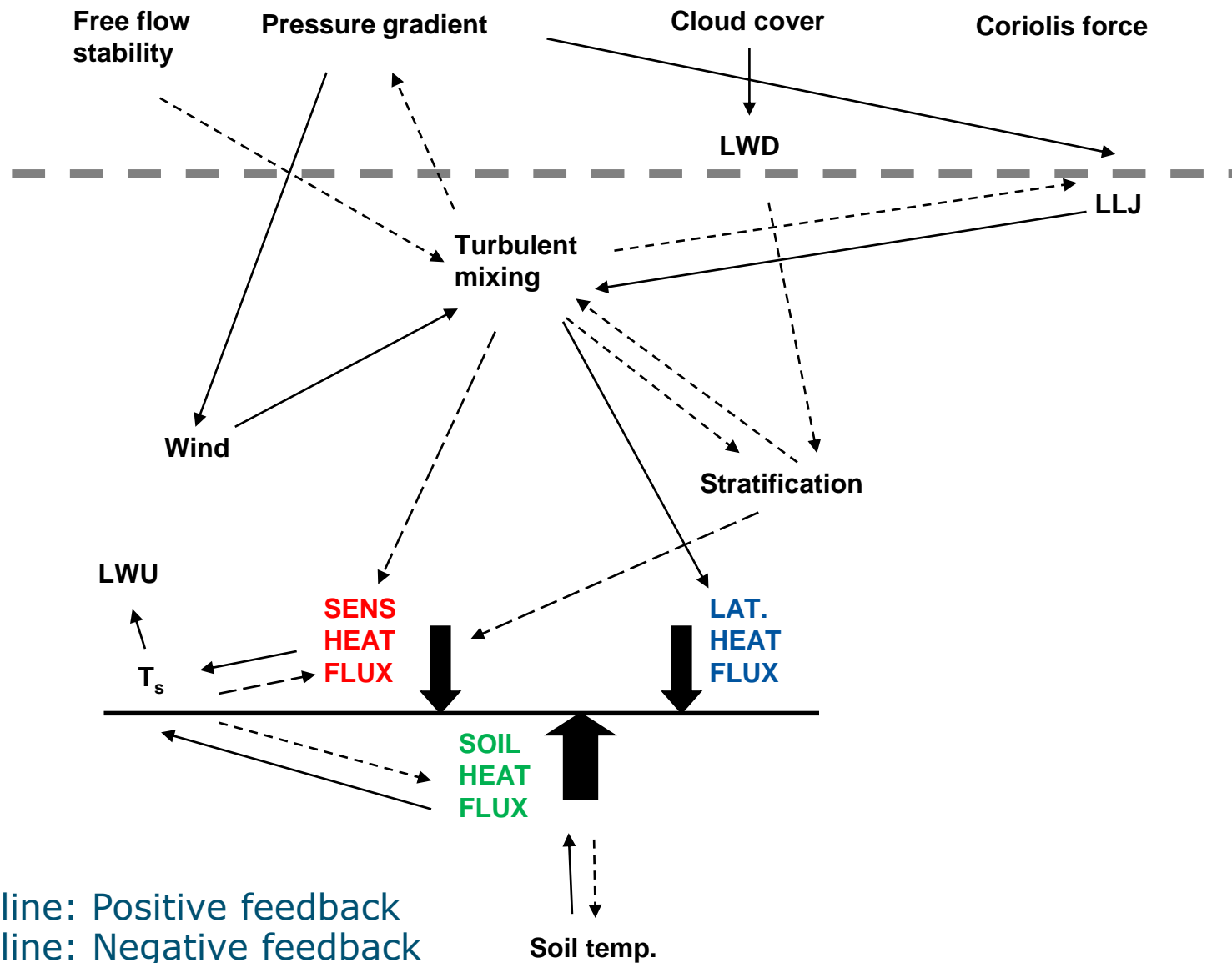
Realistic turbulence schemes improve the representation of low-level jet, wind turning, pbl depth, diurnal cycle...

But, result in negative impact on larger scale model performance (e.g., filling of cyclones)

Connection to land-atmosphere coupling, sloping terrain, (small-scale) orographic wave drag, et cetera

(Louis, 1979; Mahrt, 1987; Steeneveld et al., 2008; Sandu et al., 2013; Holtslag et al., 2013 among others)

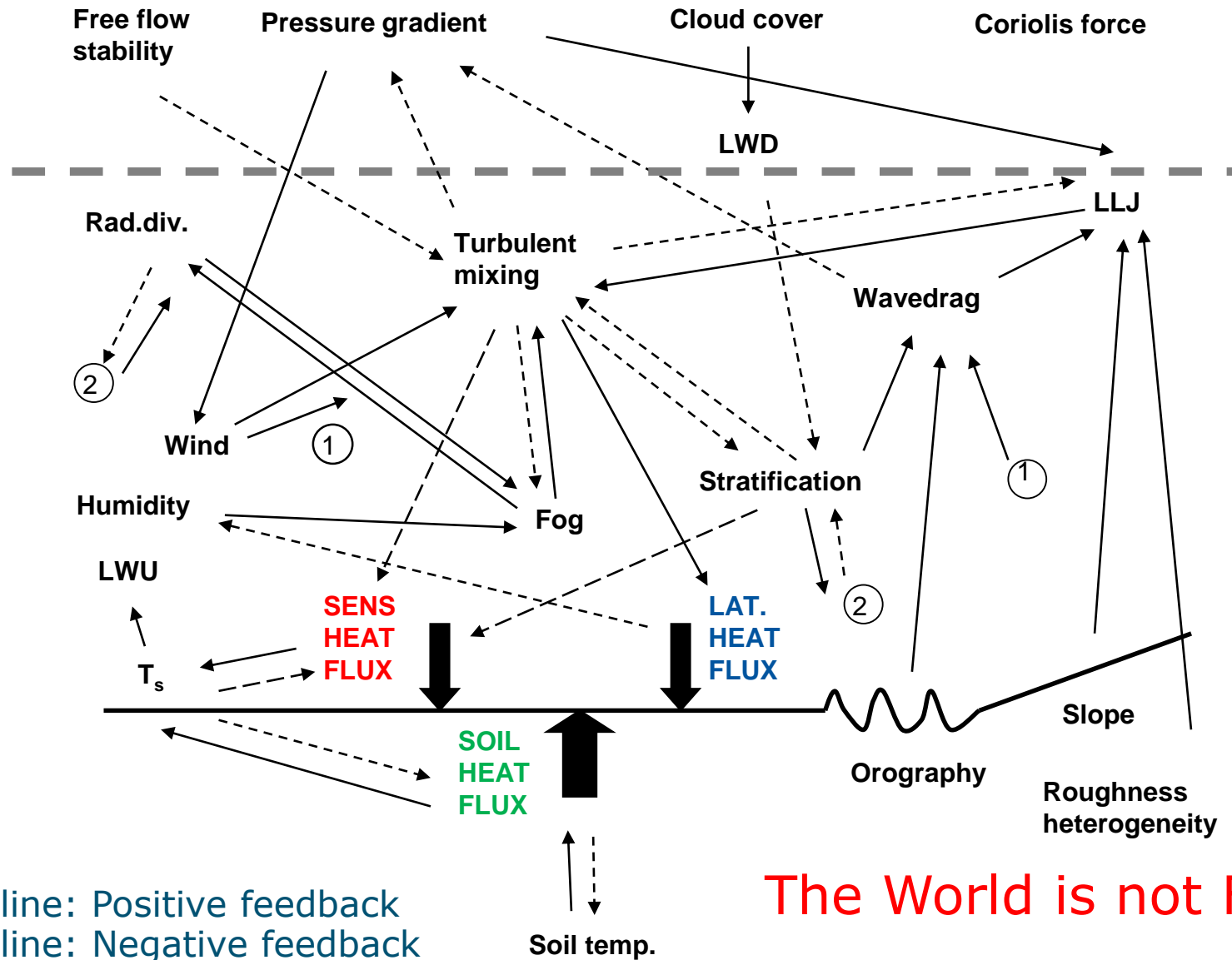
# Stable Atmospheric Boundary Layer Processes and Feedbacks





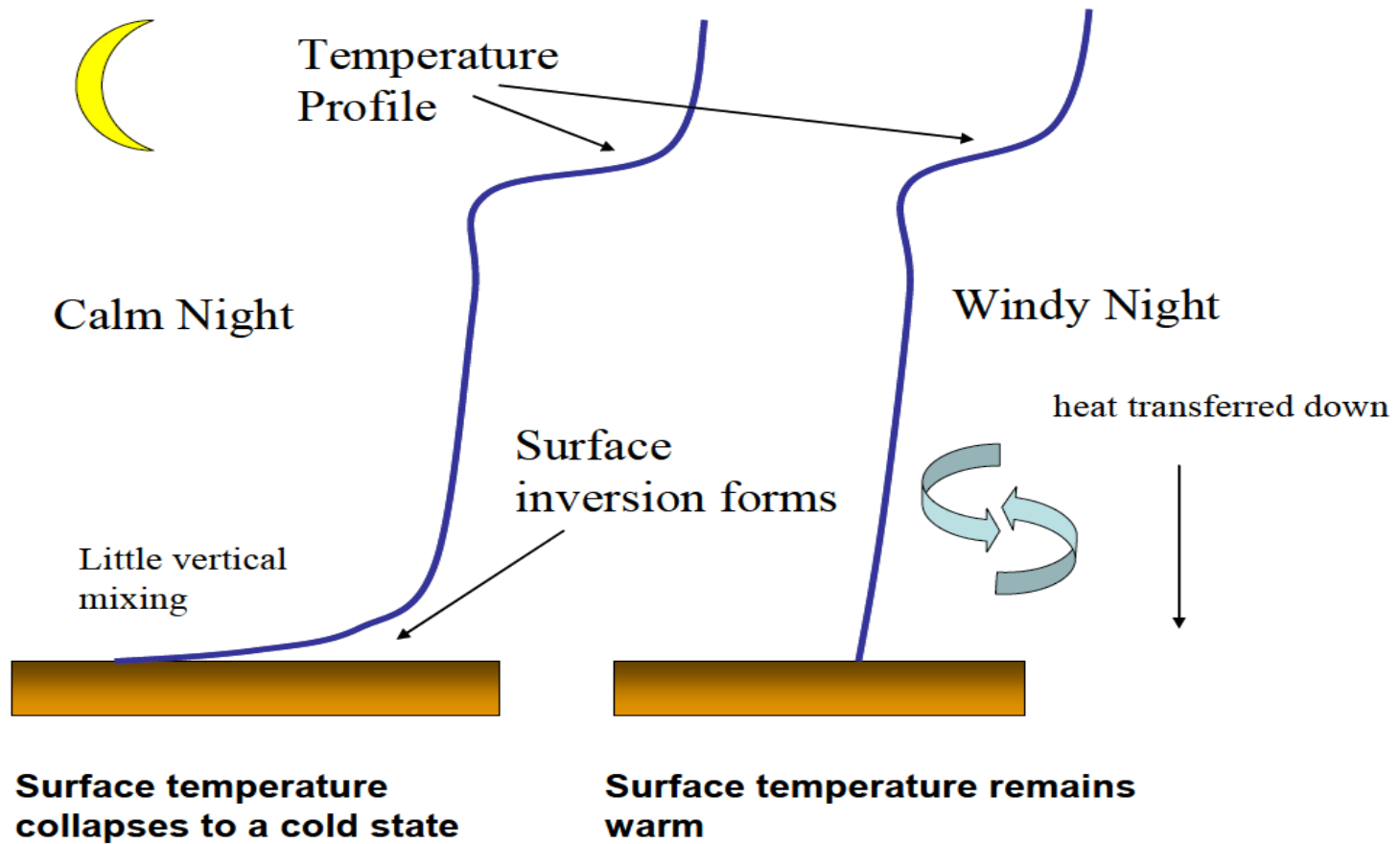


# Stable Atmospheric Boundary Layer Processes and Feedbacks



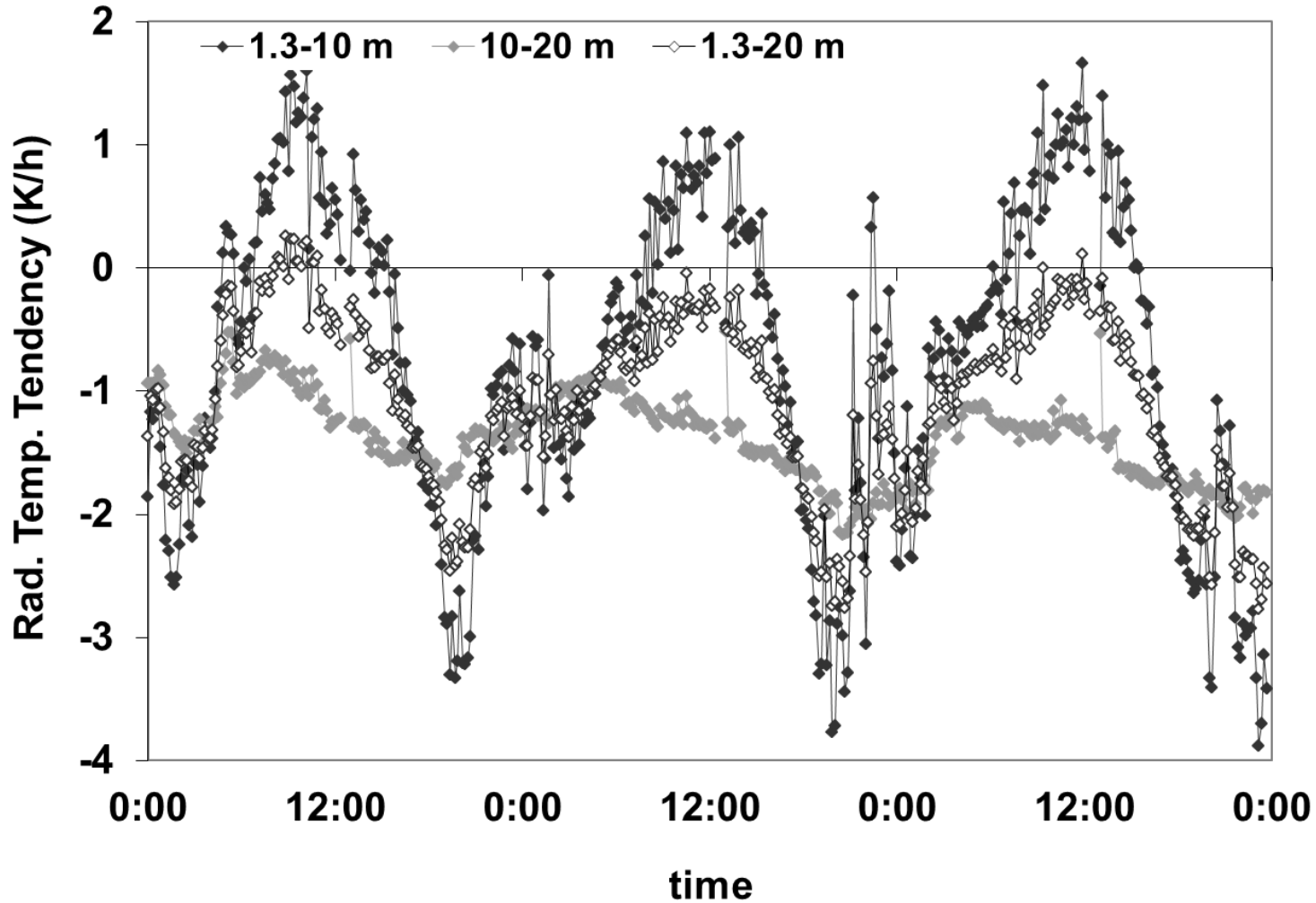
(after Steeneveld, 2014)

# Contrast between the temperature profiles in calm and windy nights under clear skies (adapted from McNider et al, 2010)



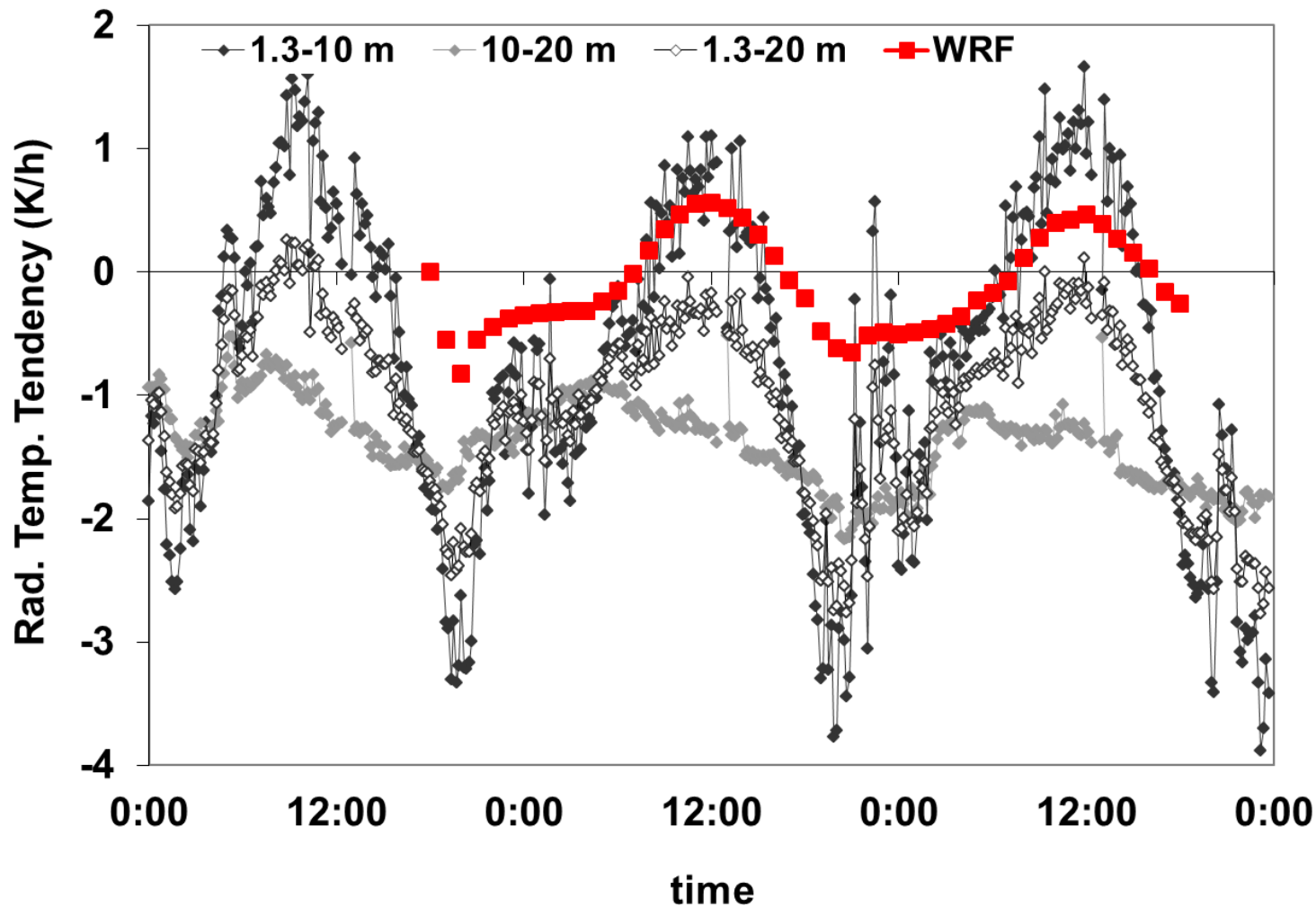
# Near surface longwave heating rates

(Three clear calm days, June 2006, Wageningen, NL)

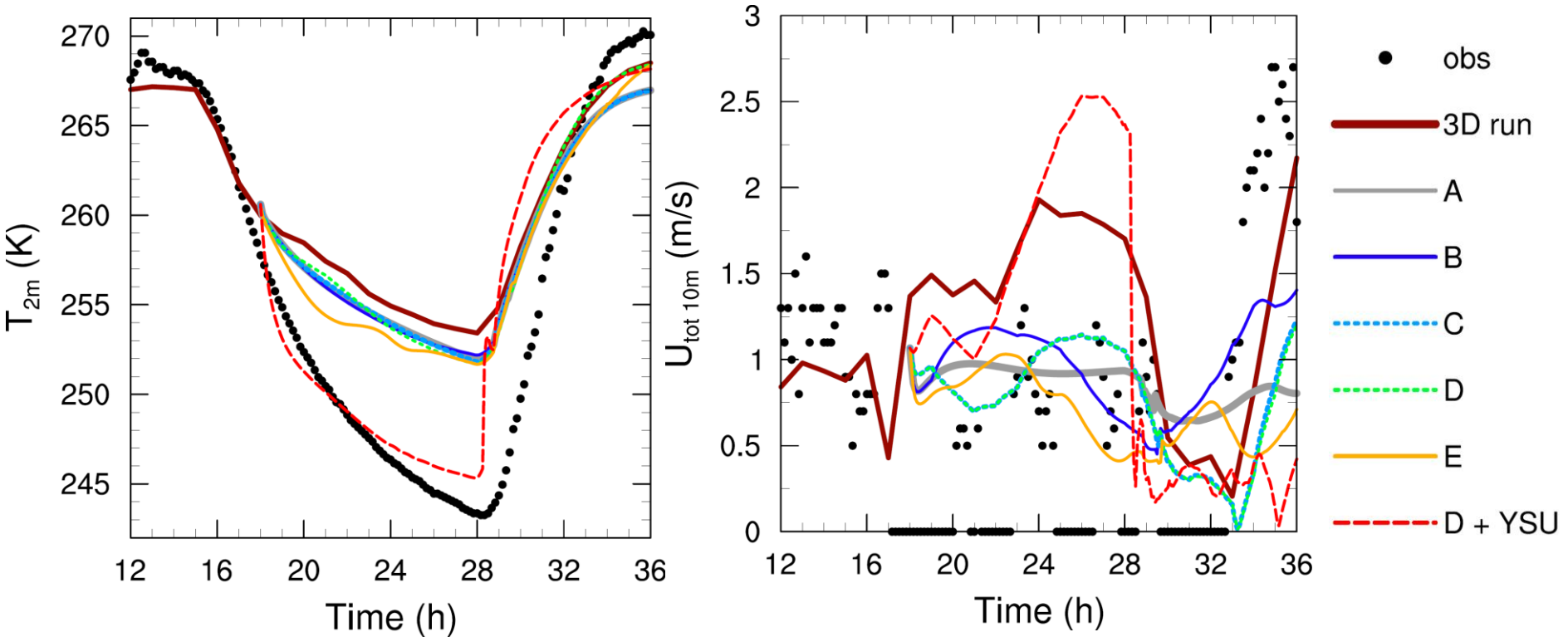


## Near surface longwave heating rates:

Typically underestimated in models (like in WRF, here at  $\sim 10$  m)



# Modelling stable boundary layers over snow and ice are still very challenging as well!



Example for Södankyla (Finland) for various WRF single column model set ups in comparison with obs and 3D

# Stable Boundary Layers: Still a challenge!

We know a lot already, but still need to better understand the many small scale processes and their non-linear interactions

Develop and test parameterizations for the very stable boundary layer when turbulence is not the dominant driver

Need for further model benchmark studies over land, sea, and ice under various atmospheric conditions  
(including radiation, advection, surface heterogeneity and coupling, fog,...)

# Stable boundary layers, more than turbulence

