

# ***Developments in large scale physics***

Presented by Yves Bouteloup

14<sup>th</sup> ALADIN Workshop  
Innsbruck, 1-4 June 2004

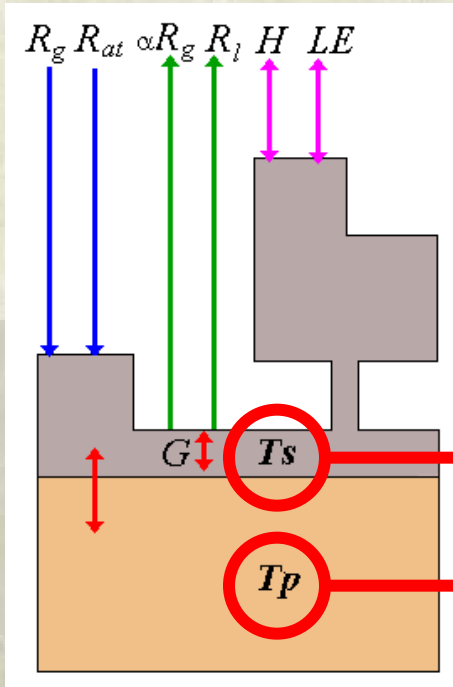
# Modifications made in the operational suite

- **Change in the analyse of the soil moisture (over land)**
  - **Change in the cloud scheme**
  - **New radiation scheme : FMR15**
- } Presented by  
F. Bouyssel  
(Prague 2003)

- **Pre-conditioning of second minimization**
- **Modifications without impact on forecasts :**
  - **New surface fields in DDH**
  - **Developments in order to use the new format of NESDIS SST files**
  - **Modifications in the codes for futur use**

# Operational surface scheme (ISBA)

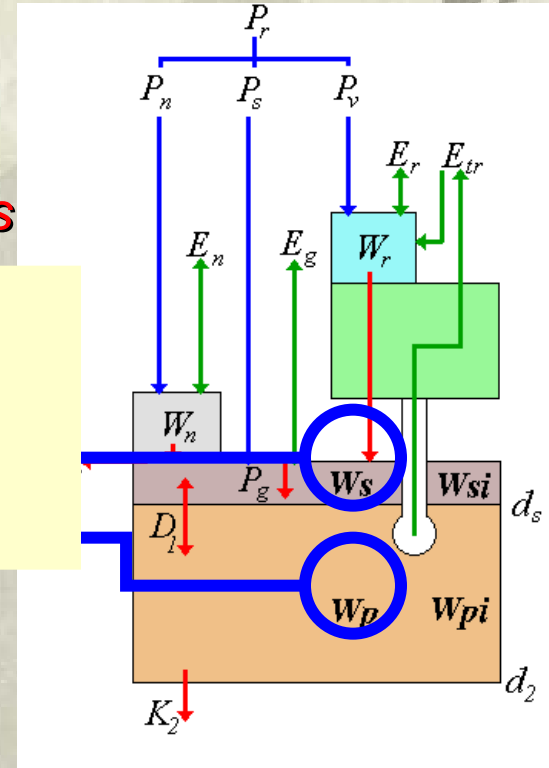
## Thermal exchange



## Analysed parameters

- Surface temperature
- "Deep" temperature
- Surface water contains
- Total water contains

## Hydrous exchange



- A description of the soil and vegetation's characteristic at any grid point is given by a database (percentage of sand, clay, albedo, fraction of vegetation, foliar index ...)

# It's very important to have a good initialisation of soil moisture over land

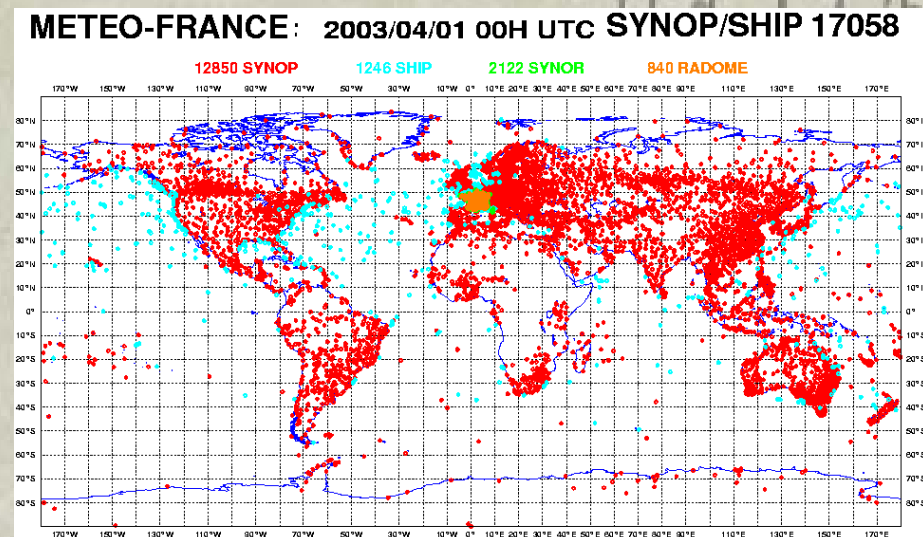
- ❖ The soil moisture has a very important impact in the distribution of the net heat surface solar flux between latent heat and sensible heat
- ❖ Has a large impact on the evolution of the PBL
- ❖ The variation of soil moisture has a time scale of several weeks.

## Difficulties of soil moisture analyse :

- ❖ No direct observation
- ❖ Large spatial variability

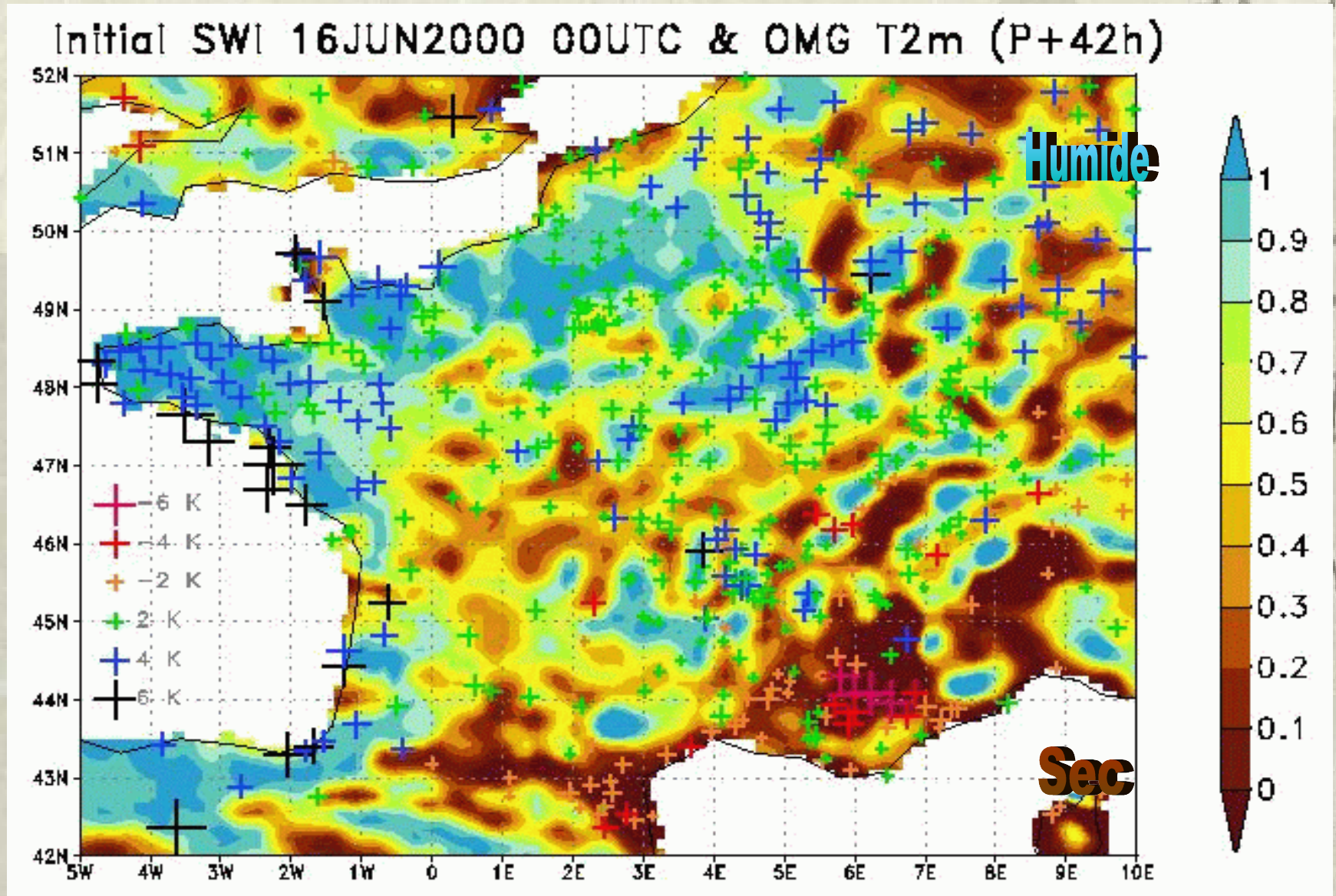
## Useful observations :

- ❖ Precipitations (pluvio, radars)
- ❖ SYNOP (T2m,H2m)
- ❖ Satellite observations (MW,IR)





# Correlation between T2m errors and soil moisture for a 42 hours range forecast 17 June 2000 at 18h



# Operational surface analyse

1) First an optimal analyse of T<sub>2m</sub> et H<sub>2m</sub>

$$\Delta T_{2m} = T_{2m}^a - T_{2m}^f \quad \Delta RH_{2m} = RH_{2m}^a - RH_{2m}^f$$

2) Then correction of the surface variables (T<sub>s</sub>, T<sub>p</sub>, W<sub>s</sub>, W<sub>p</sub>) using 2m difference between analyse values and observations.

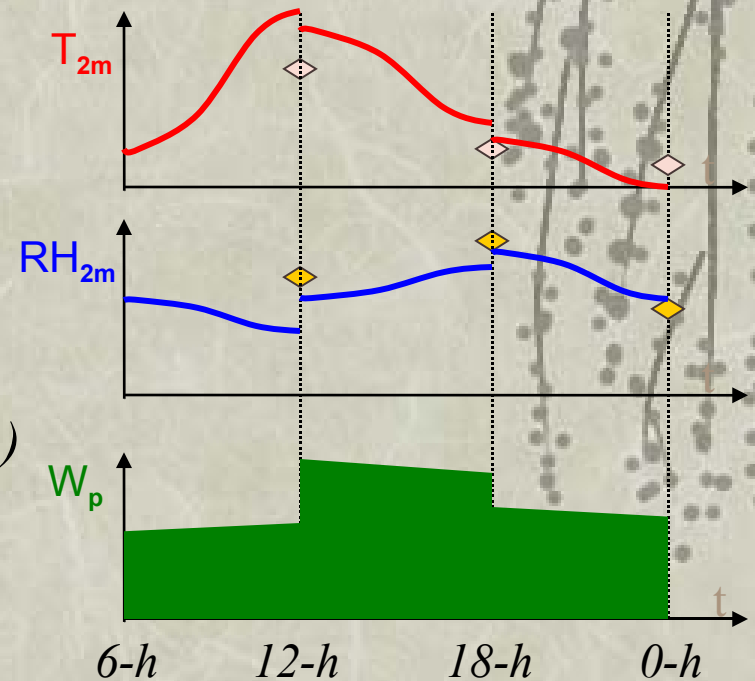
$$T_s^a - T_s^f = \Delta T_{2m}$$

$$T_p^a - T_p^f = \Delta T_{2m} / \beta$$

$$W_s^a - W_s^f = \alpha_{W_s T} \Delta T_{2m} + \alpha_{W_s RH} \Delta RH_{2m}$$

$$W_p^a - W_p^f = \alpha_{W_p T} \Delta T_{2m} + \alpha_{W_p RH} \Delta RH_{2m}$$

$$\alpha_{W_p/sT/RH} = f(t, \text{veg}, LAI/Rs_{min}, \text{texture}, \text{atm.cds.})$$



# Modifications introduced in the surface analyse

- spatial smoothing of soil wetness index
- new statistical scheme for background error (for 2m fields)
- reduced corrections for deep soil moisture, dividing by 2 the coefficients
- removal of time smoothing
- taken into account of the solar zenithal angle into the optimal coefficients

These modifications were proposed along last years by Stjepan Ivatek-Sahdan, Agnesz Mika and François Bouyssel (ALADIN Newsletter 21 and 22) and they were evaluated by Adam Dziejczic and François Bouyssel in a cheap assimilation suite along June 2003 (Newsletter 24)

The suite is based on a blending between the modified surface and low-level fields and the large-scale upperair analysis increments from the operational 4d-var assimilation suite.



# Soil wetness index for 2 may 2004

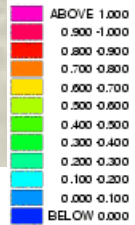
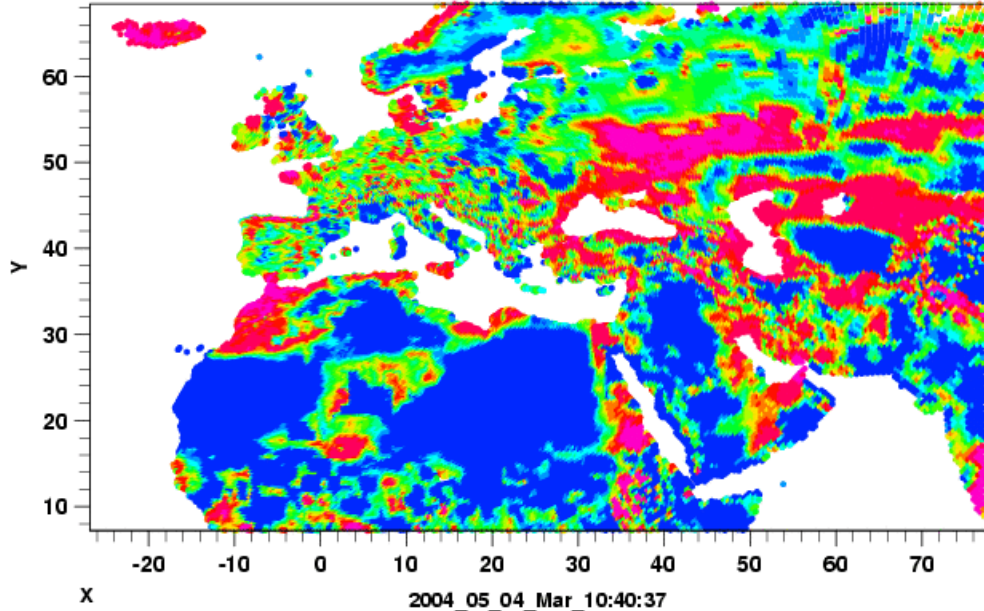
Min=-2.216  
Max=3.213  
Moy=0.3390  
Rcm=0.7318

**OPER**



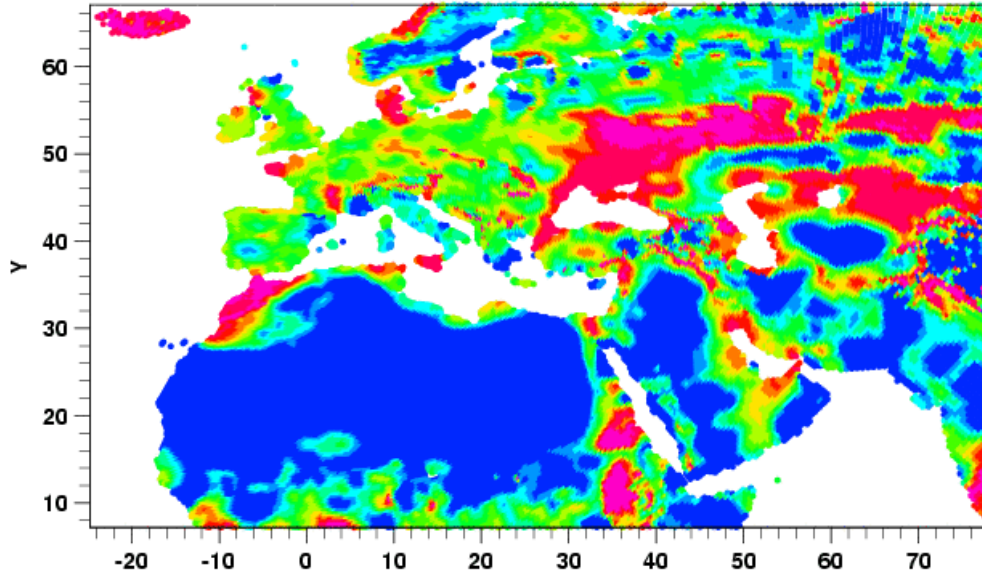
oper\_preswi\_2004050218.b

ARPEGE/ALADIN

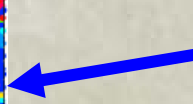


Min=-1.971  
Max=3.212  
Moy=0.2934  
Rcm=0.7233

dbi\_preswi\_  
ARPEGI



**DBL**





# Modifications of the soil moisture analyse : impact on the scores

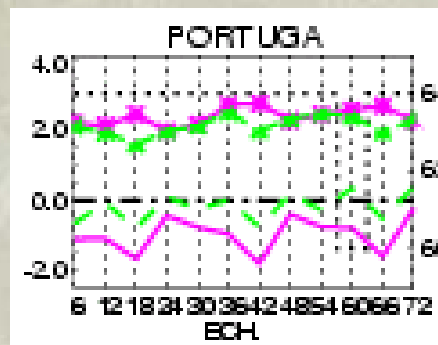
Experiments with a “simplified” assimilation cycle :

- \* 2 month in summer : 1/5/2003 to 3/7/2003
- \* 2 month in winter : 1/12/2003 to 3/2/2004

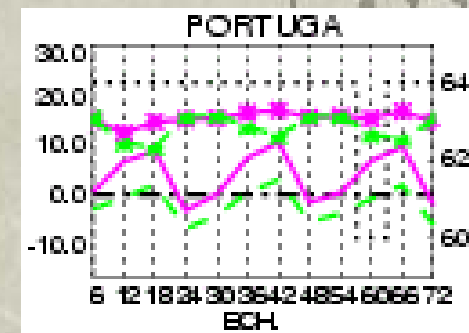
Scores of the surface parameters forecast :

- \* neutral on 10m wind and surface pressure
- \* neutral on T2m and H2m in winter
- \* improvement on T2m and H2m in summer over Europe, neutral elsewhere

T2m



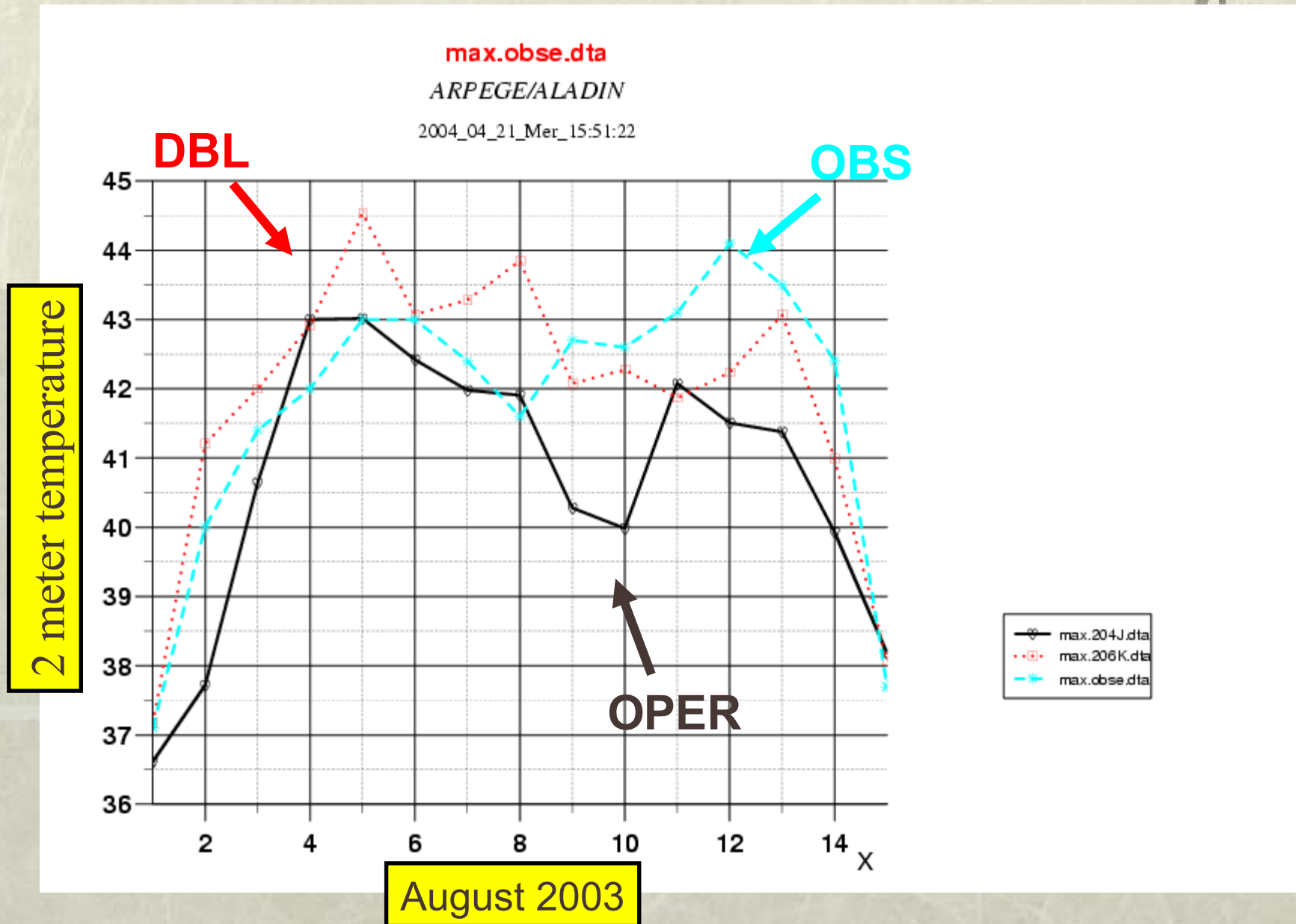
H2m



Scores SYNOP  
16 forecast 96h  
15/06/03 to  
3/07/03

DBL  
OPER

# Behaviour of the double suite during 2003 midsummer heat : Maximum forecast temperature over France



# Behaviour of the double suite during 2003 midsummer heat

## Number of grid point with $T_{max} > 35^\circ$ or $40^\circ$

Number of grid point

nbpts.sup40.206K.dta

ARPEGE/ALADIN

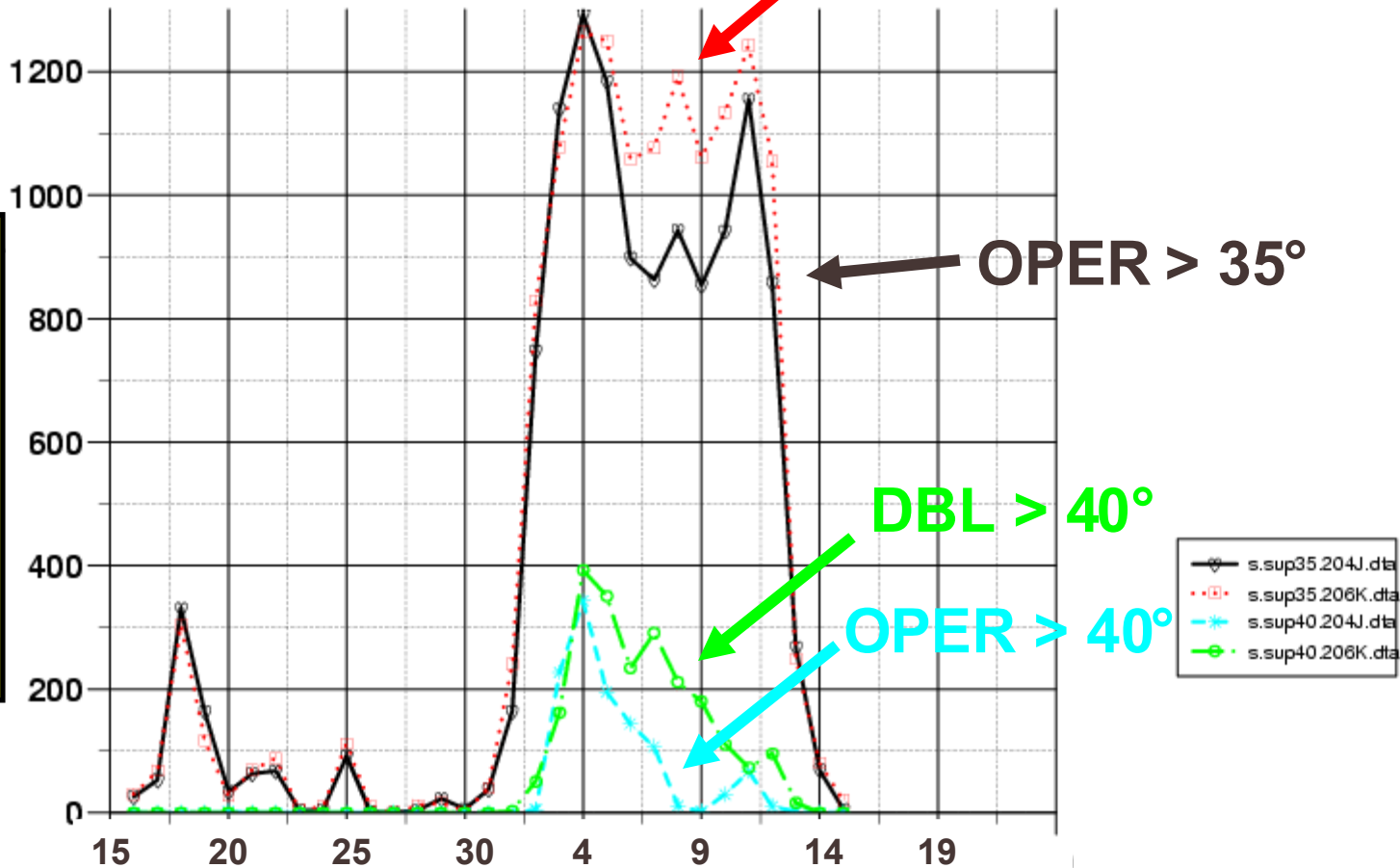
2004\_04\_21\_Mer\_15:53:39

DBL > 35°

OPER > 35°

DBL > 40°

OPER > 40°



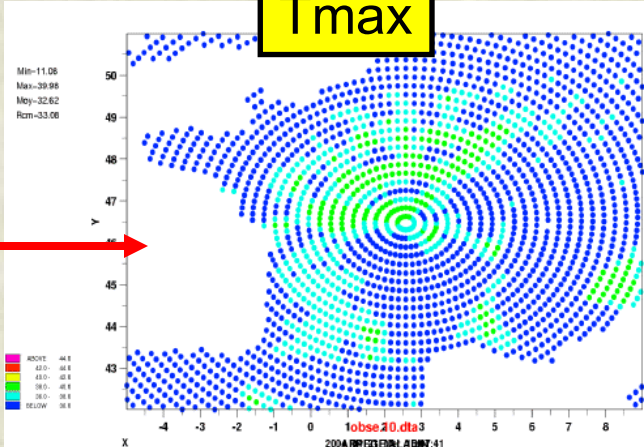
July and August 2003

# Behaviour of the double suite during 2003 midsummer heat

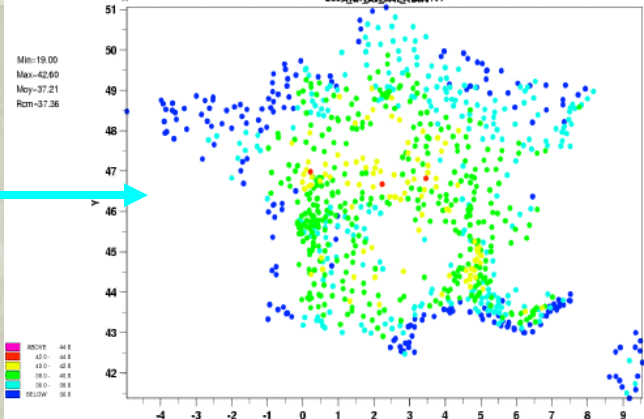
Tmax

Convective precipitation

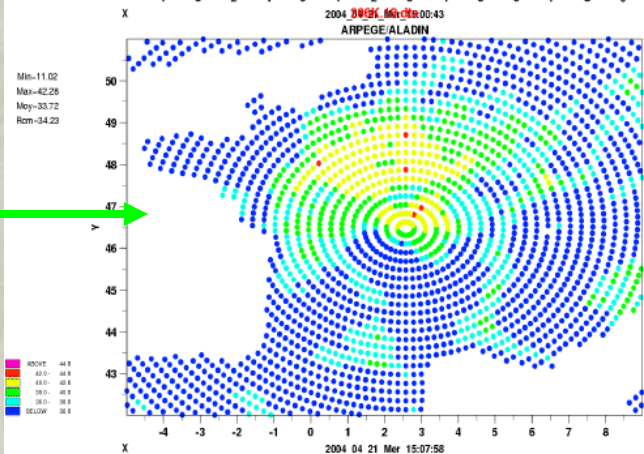
OPER



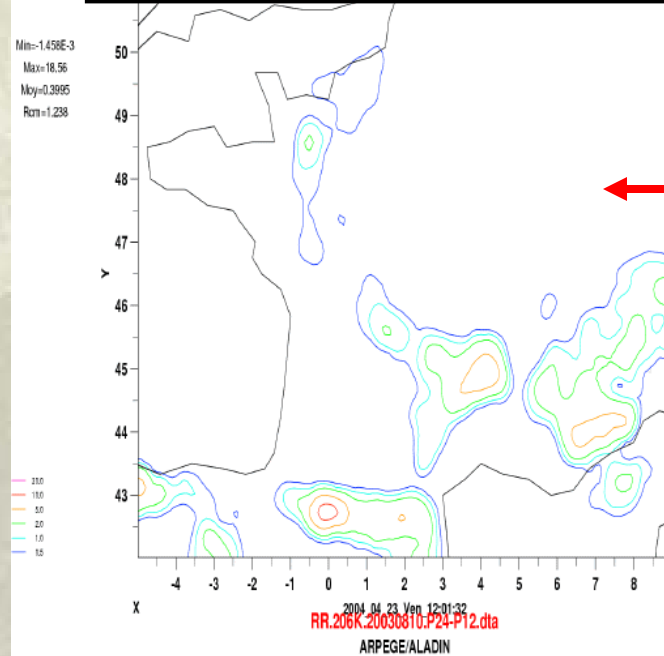
OBS



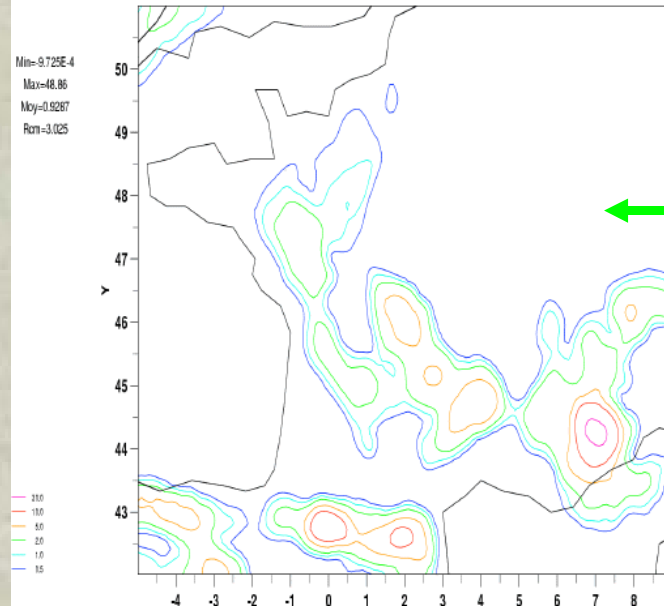
DBL



OPER



DBL





# Scores of the double suite (AC)

## Geopotential

## Temperature

GEOPOTENTIEL : PA.F.0/AC-PAD.F.0/AC

(/1.00m) Chaîne 2004\_02, Rayonnement et Nebulosite, bis  
47 cas, 15/03/2004\_00UTC -> 05/05/2004\_00UTC

TEMPERATURE : PA.F.0/AC-PAD.F.0/AC

(/0.05K) Chaîne 2004\_02, Rayonnement et Nebulosite, bis  
47 cas, 15/03/2004\_00UTC -> 05/05/2004\_00UTC

Eqm

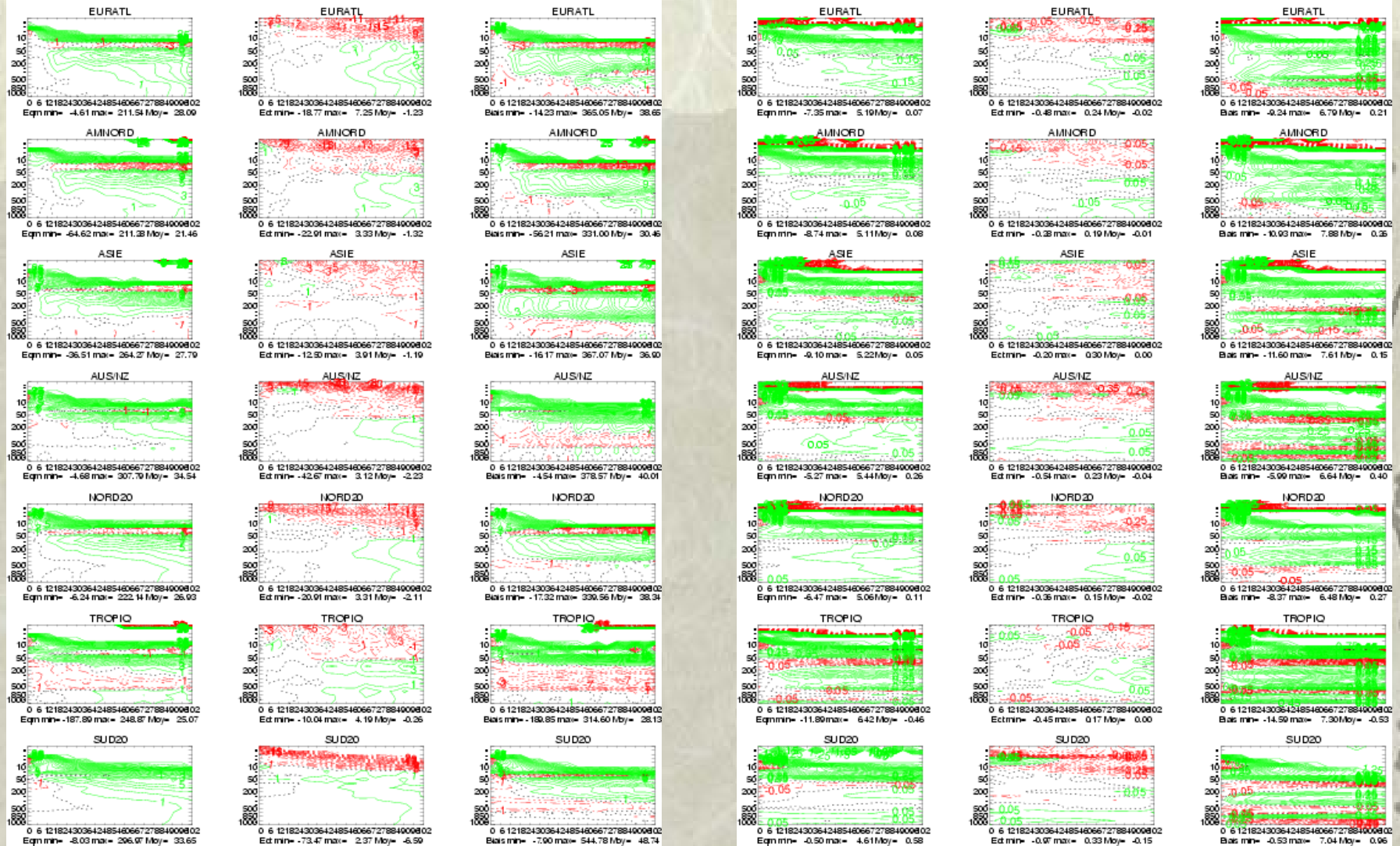
Ect

[Biais]

Eqm

Ect

[Biais]



# Scores of the double suite (AC)

Wind

Moisture

VENT : PA.F 0/AC-PAD.r 0/AC

(/0.20m/s) Chaîne 2004\_02, Rayonnement et Nebulosite, bis  
47 cas, 15/03/2004\_00UTC -> 05/05/2004\_00UTC

HUMIDITE : PA.F 0/AC-PAD.r 0/AC

(/1.00%) Chaîne 2004\_02, Rayonnement et Nebulosite, bis  
47 cas, 15/03/2004\_00UTC -> 05/05/2004\_00UTC

Eqm

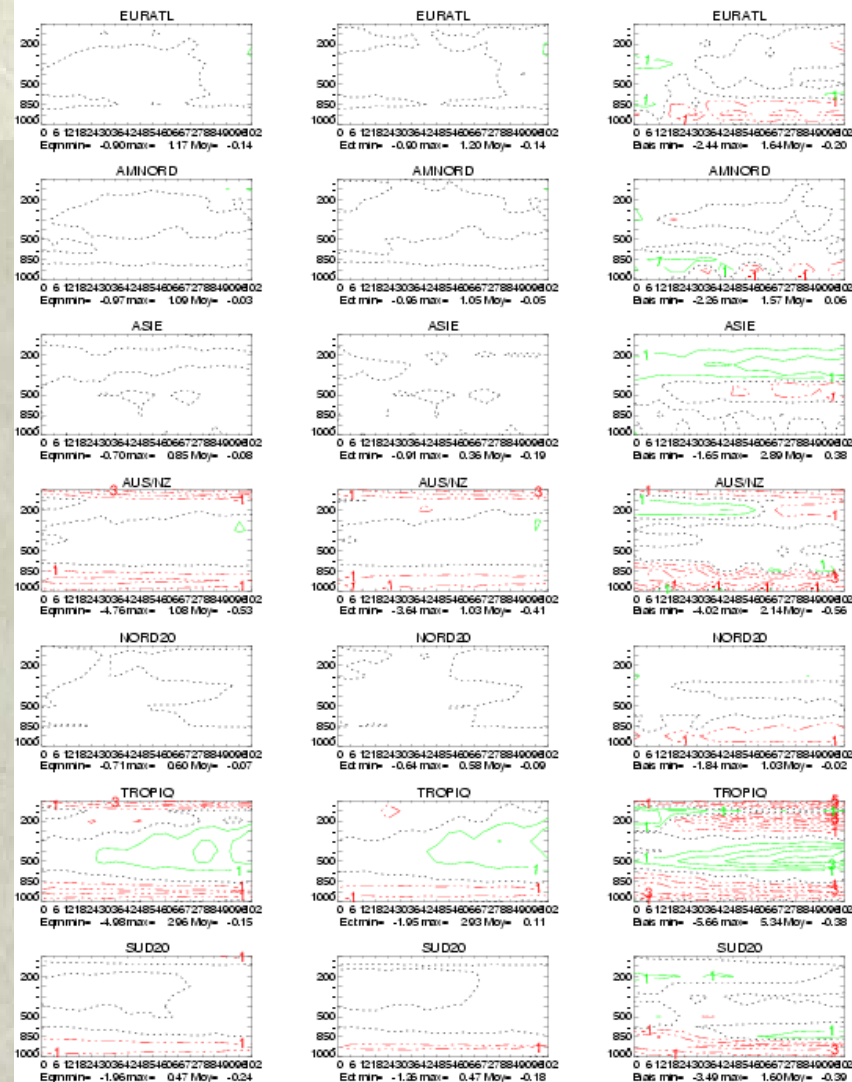
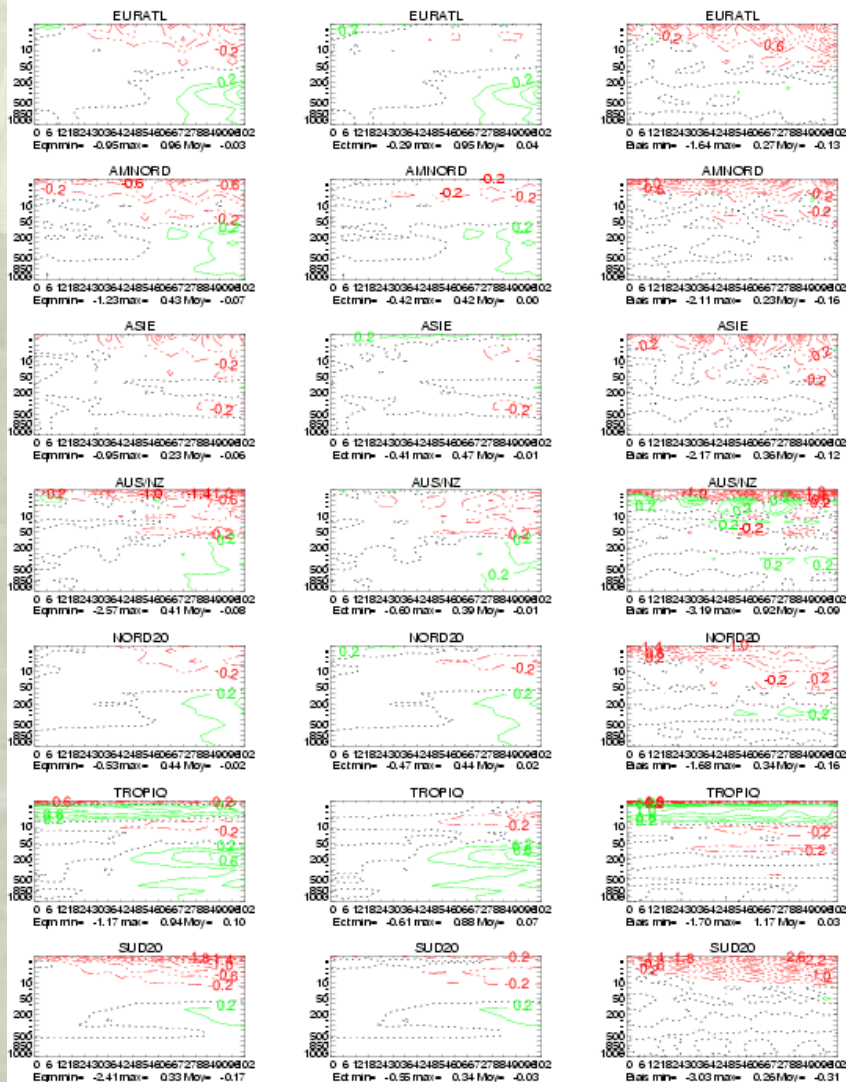
Ect

[Biais]

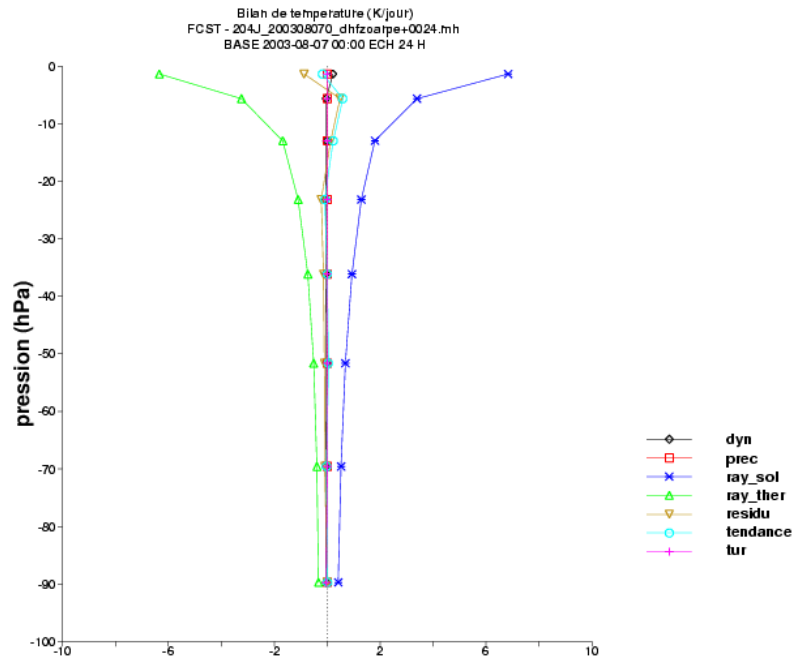
Eqm

Ect

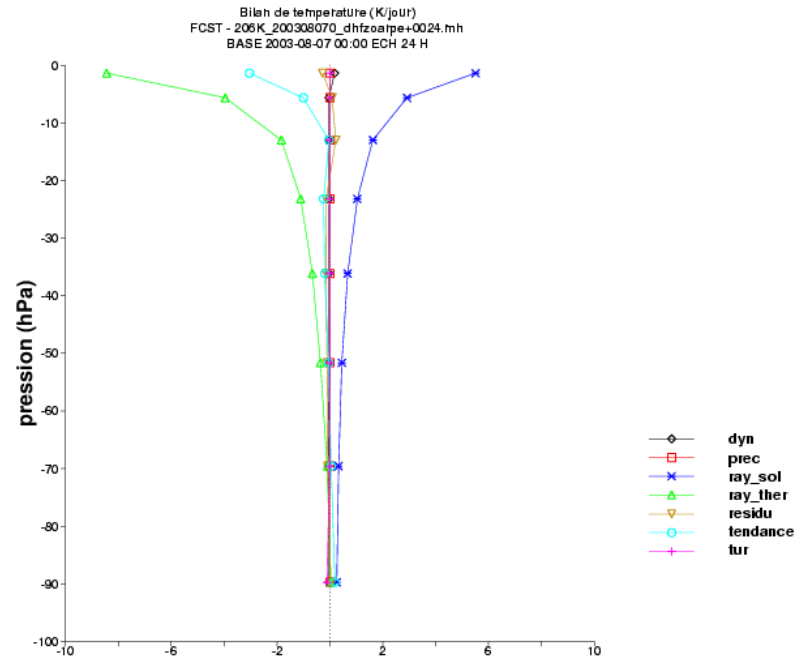
[Biais]



# Excessive cooling on the 2 first levels of the model



oper



double

# Prospects

1) In operation since 24 of may

2) Next experimental suite :

- Use of UGAMP climatology for ozone (excessive cooling)
- Use of aerosols climatology
- Variable mixing length for turbulence

3) Medium time prospects ... :

- Collaboration between GMGEC (Climat) and GMAP on new schemes for turbulence (TKE), microphysic (Lopez) and convection (Gérard, Gueremy ???)



## The UGAMP climatology

The UGAMP ozone climatology has been built up by Dingmin Li and Keith P. Shine at the department of meteorology of the University of Reading.

It's a 4-dimensional distribution of atmospheric ozone resulting from the combination of several observational data sets.

These data sets include satellite observations (SBUV, SAGE II, SME, TOMS) as well as ozone sonde data provided by the Atmospheric Environment Service of Canada. This global climatology covering five years (1985 to 1989),

For more details see :

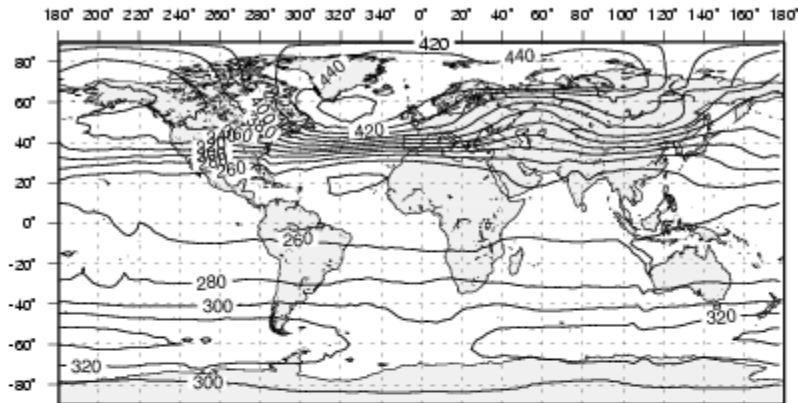
Li, D. and K. P. Shine : A 4-dimensional Ozone Climatology for UGAMP Model, UGAMP Internal Report 35, April 1995.

Or :

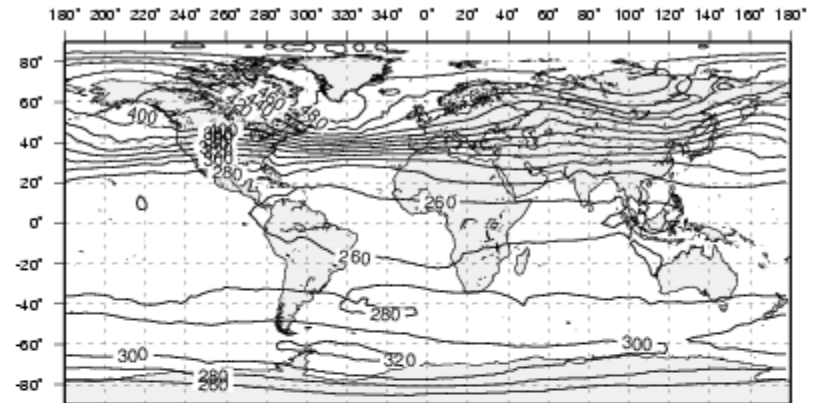
<http://badc.nerc.ac.uk/data/ugamp-o3-climatology>

UGAMP climatology : total ozone in dopson  
(ARPEGE/ALADIN value : 284 DU)

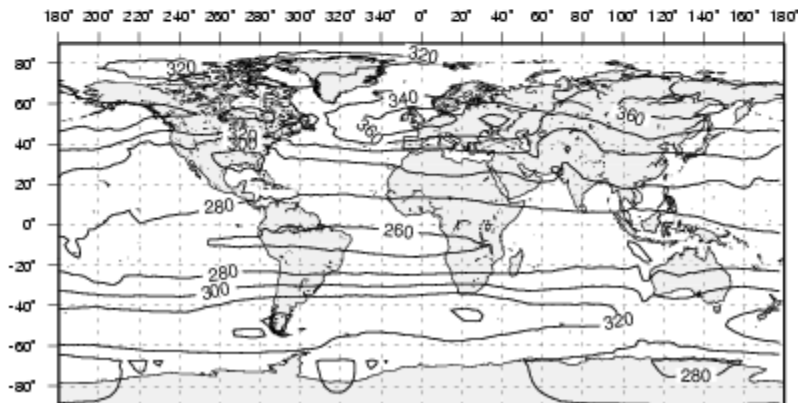
january



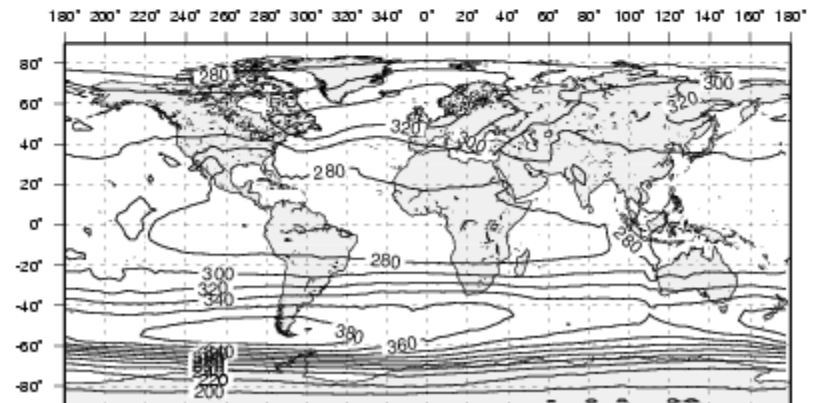
march



july



september



## Use of the UGAMP climatology

The same analytical formula is used : 
$$\int_0^p q_{o_3} dp = \frac{a}{1 + (b/p)^{c/2}}$$

In ARPEGE /ALADIN  $a=0.06012$ ,  $b=3166$  and  $c=3$   
This work is similar to the one made by Rada & al (2000)  
in Bucharest but with a climatology

The new system try to fit the formula to the climatology under 3 constraints :

- Total ozone is conserved
- Altitude of maximum
- Value of the maximum in DU/hpa

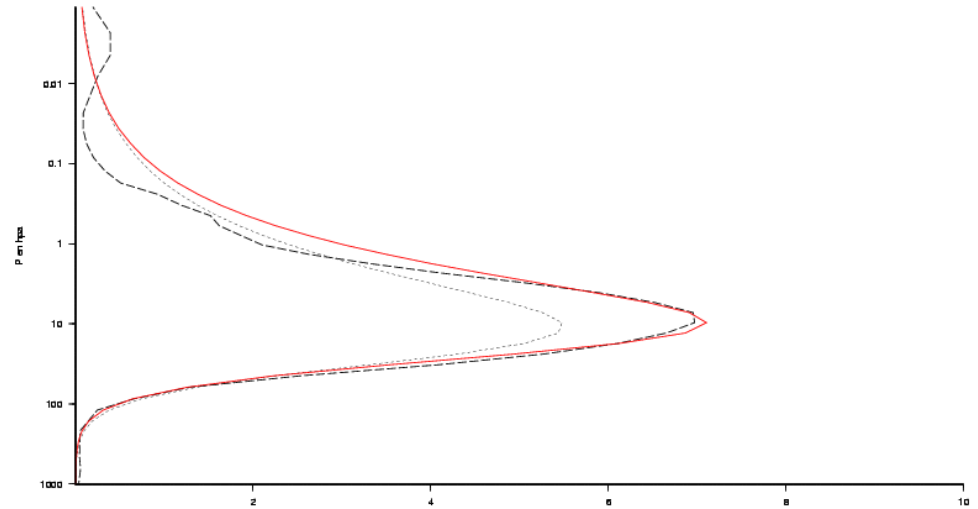
Rada C., A. Sima and M. Caian, 2000 : Ozone Profile Fitted to Bucharest Measured Data, ALADIN Newsletter 18, 51-57

Sometimes the result is very good:

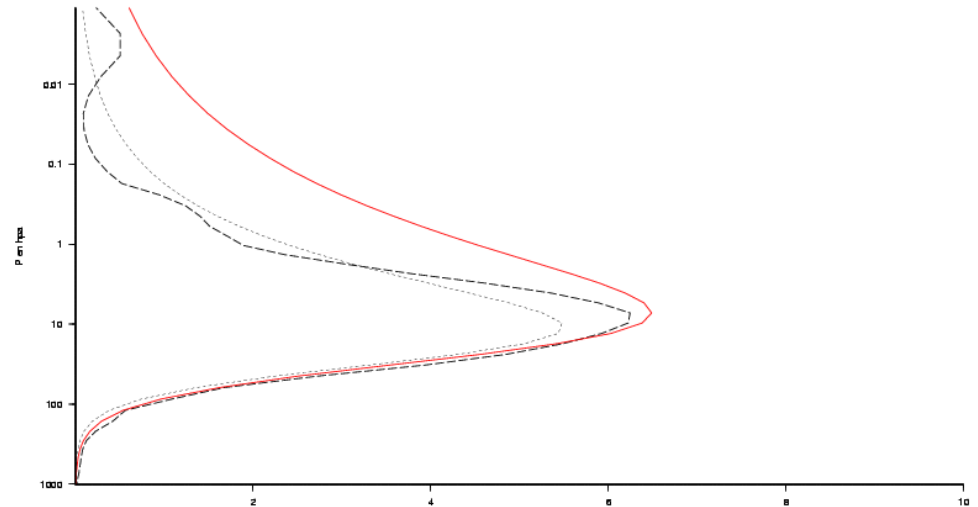
Arpege/aladin profile (dot)  
Climatology (dash)  
Fitted profile (red solid)

Sometimes it's not perfect !

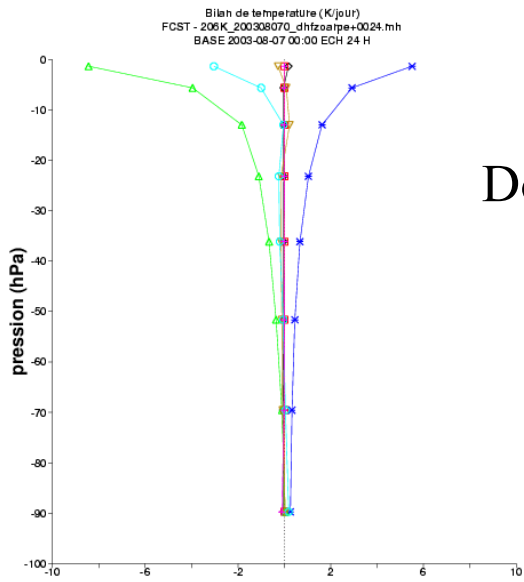
July 0°E 20°N



July 0°E 45°N

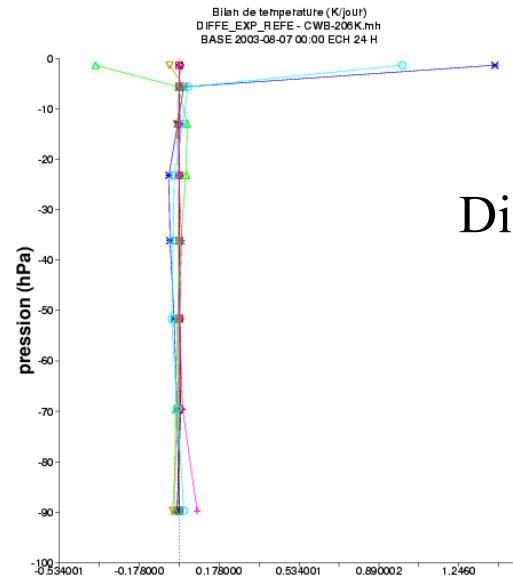




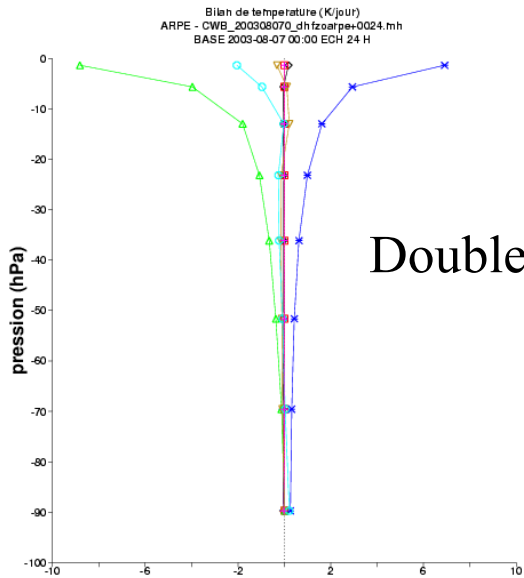


Double suite

Global impact of  
 the new ozone profiles



Difference



Double suite + UGAMP



# Modifications of the turbulence scheme

(work of Eric Bazile on a GABLS case)

GABLS : GEWEX Atmospheric Boundary Layer Study  
GEWEX : Global Energy and Water Cycle Experiment

J. Cuxart, A.A.M. Holtslag, R.J. Beare, E. Bazile, A. Beljaars, A. Cheng, L. Conangla, M. Ek, F. Freedman, R. Hamdi, A. Kerstein, H. Kitagawa, G. Lenderink, D. Lewellen, J. Mailhot, T. Mauritsen, V. Perov, G. Schayes, G-J. Steeneveld, G. Svenson, P. Taylor, S. Wunsch, W. Weng, and K-M. Xu (2004). Single-column model intercomparison for a stably stratified atmospheric boundary layer (in internal review)

Louis formulation for turbulence :

$$F_{\psi} = \frac{\rho K g \Delta \psi}{\Delta \phi}$$

with :

$$\begin{cases} K_m & = & L_m L_m \left| \frac{\partial V}{\partial z} \right| F_m (R'_i) \\ K_h & = & L_m L_h \left| \frac{\partial V}{\partial z} \right| F_h (R'_i) \end{cases}$$

Where  $R'_i$  is a function of the Richardson number  $R_i$ , of the mixing length  $L_h$  and  $z$

## Modifications of the turbulence scheme

- The mixing length becomes a function of the PBL height
- The PBL height is computed following Troen and Mahrt (1986)

→ Slight improvement of wind and temperature profiles in stable cases (1d)

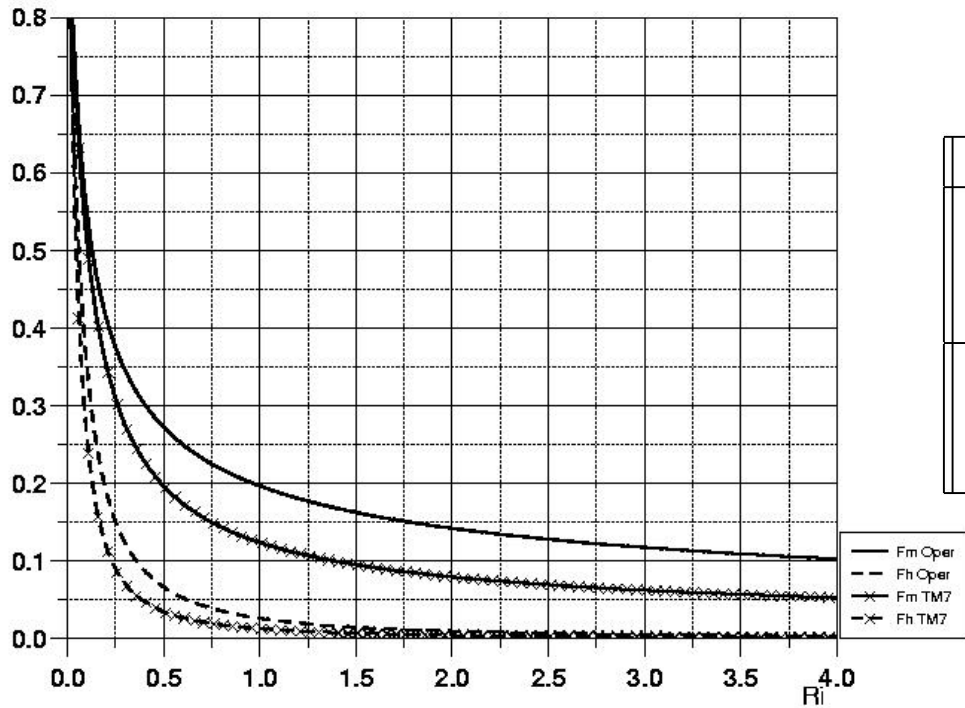
- The function  $F_m$  and  $F_h$  has been modified

→ Improvement on the GABLS case (better low level jet), and in 3d simulations

Troen I. And L. Mahrt, (1986). A simple model of the atmosphere boundary layer; sensitivity to surface evaporation. *Bound.-Layer Meteor.*, **37**:129-148.

# Modifications of the function $F_m$ and $F_h$

Fh(Ri) and Fm(Ri)

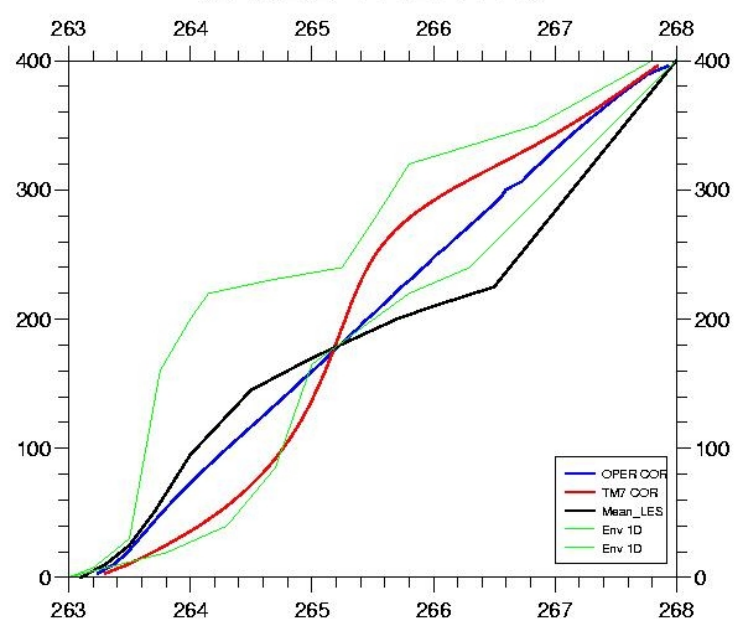


	<i>Oper</i> $b = d = 5$	<i>Test</i> $d = 1$ $b = 5$ $k = 5$
$\frac{1}{F_m}$	$1 + \frac{2bR_i}{\sqrt{1+dR_i}}$	$1 + \frac{2bR_i}{\sqrt{1+dR_i}}$
$\frac{1}{F_h}$	$1 + 3bR_i\sqrt{1+dR_i}$	$1 + 3kbR_i\sqrt{1+dR_i}$



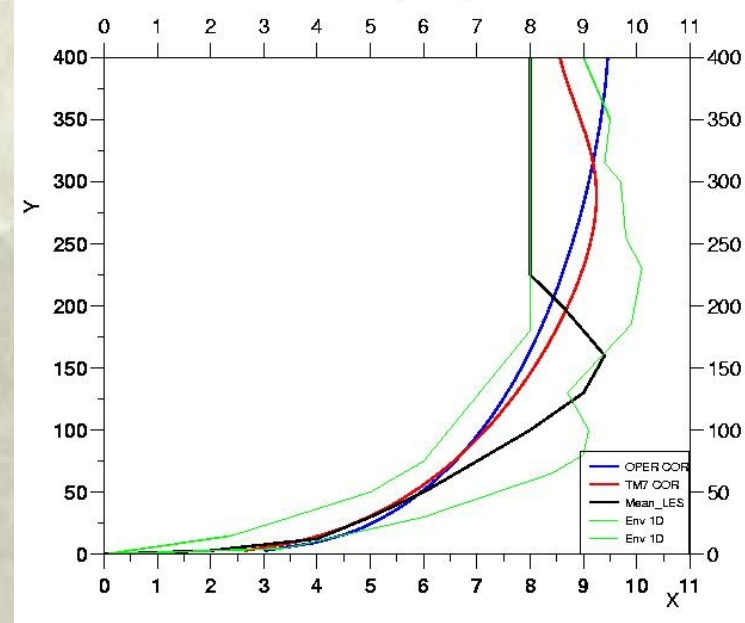
# Impact of the modifications of the turbulence scheme

## GABLS THETA H9



Gabls case :  
Wind →  
← theta

## WIND MODULE (m/s) GABLS H9



3d test 24 forecasts in january 2004 :  
Improvement of 10m wind direction →

## NORD20

