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***The hybrid ensemble variational data
assimilation system in HIRLAM***
(sensitivity to the ensemble generation approach)

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Thanks to Kai Sattler, Åke Johansson and Ole Vignes

Approximate solution to the sequential update problem



non-linear state space model

$$y_{t_i} = \mathcal{H}_{t_i}(X_{t_i}) + \epsilon_{t_i}$$
$$X_{t_i} = \mathcal{M}(t_i, t_{i-1})(X_{t_{i-1}}) + T_i \xi_{t_i}$$

tangent-linear approximation

$$X_{t_i} = X_{t_i}^f + \delta x_{t_i}, \quad X_{t_i}^f = \mathcal{M}(X_{t_{i-1}}^f)$$
$$y_{t_i} = \mathcal{H}(X_{t_i}^f) + H \delta x_{t_i} + \epsilon_{t_i}$$
$$\delta x_{t_i} = M_{(t_i, t_{i-1})} \delta x_{t_{i-1}} + T_i \xi_i$$

extended Kalman filter solution

$$\delta x_\tau = K_\tau (y_\tau - \mathcal{H}(X_\tau^f) - H_\tau \delta x_\tau)$$
$$K_\tau = B_\tau^f H_\tau^T (R_\tau + H_\tau B_\tau^f H_\tau^T)^{-1}$$
$$B_\tau = (I - K_\tau H_\tau) B_\tau^f$$

3D-Variational solution

$$\delta x_\tau = \operatorname{argmin} L(\delta x_\tau | X_\tau^f, y_\tau)$$
$$L(\delta x_\tau | X_\tau^f, y_\tau) = 0.5(\delta x_\tau)^T B^{-1} \delta x_\tau$$
$$+ 0.5(y_\tau - \mathcal{H}(X_\tau^f) - H_\tau \delta x_\tau)^T R_\tau^{-1} (y_\tau - \mathcal{H}(X_\tau^f) - H_\tau \delta x_\tau)$$

Ensemble Kalman filter

$$B_\tau^f = Z_\tau^f (Z_\tau^f)^T$$
$$B_\tau = Z_\tau^a (Z_\tau^a)^T$$
$$Z_\tau^a = Z_\tau^f T_\tau$$

Hybrid Variational Ensemble Kalman filter



Different approaches for using ensembles in variational data assimilation

- **Covariance modelling with parameters of the covariance model determined from an ensemble.** Use for example a wavelet-based covariance model (Alex Deckmyn; Loik Berre et al. Meteo-France)
- **Use the ensemble-based covariances in a hybrid variational ensemble data assimilation** (Barker et al. WRF, UK Met.Office, HIRLAM)
- **Ensembles can also be used to determine static background error statistics**

Toward data assimilation using flow-dependent structures in HIRLAM



Rapid update cycle

(a short range forecasting of events of shorter temporal and spatial scales of variability)

Meso-scale data assimilation

(extraction from observations information about atmospheric state related to processes with shorter temporal and spatial scales of variability)

Require

- Structures of background error covariance dependent of the observation network
- Structures of the background error covariance for meso-scale processes, which are dependent on the large scale forcing and are difficult to derive analytically

Different approaches for the generation of ensembles of analyses perturbations



dynamics

The EuroTEPS perturbations (based on global singular vector perturbations targeted to Europe with 24h optimization time; both the initial and the evolved singular vectors are used to sample EuroTEPS)

dynamics+observation network

The ETKF rescaling perturbations (LAM perturbations with EuroTEPS boundaries based on the Generalized Breeding technique; both the dynamical instabilities and the density and the quality of the observation network are used to construct transformation; multiplicative and additive inflation are employed)

observation network

The EnsDA perturbations (LAM perturbations with EuroTEPS boundaries; the ensemble of data assimilation runs with stochastically perturbed observations; are not tuned well enough yet)



HIRLAM approach to use ensembles in 3D-Var

- **Sample (or Construct)**

perturbations which reflect structures of the analysis error (*EuroTEPS*, *ETKF* or *EnsDA*)

- **Grow** flow-dependent structures by integrating analysis ensemble forward in time to obtain the 6h forecast perturbations.

- **Perform** the variational data assimilation blending the structures of the full-rank statically and analytically deduced B_{3D-Var} and the flow- and observation-network dependent structures of the rank-deficient B_{ens}^f .

- **Repeat** Steps 1-3



Lorenc (2003) augmentation of the control vector space:

spectral space

$$J(\delta x_{3D-Var}, \alpha) = \beta_{3D-Var} J_{3D-Var}(\delta x_{3D-Var}) + \beta_{ens} J_{ens}(\alpha) + J_o$$

$$\frac{1}{\beta_{3D-Var}} + \frac{1}{\beta_{ens}} = 1 \quad J_{ens} = \frac{1}{2} \alpha^T A^{-1} \alpha$$

Spatial mean of $\alpha_k = 0$;
Spatial variance of $\alpha_k = 1/K$ is constant and controls amplitude;
Horizontal auto-correlation controls smoothness of α_k fields

The same α_k fields for all vertical levels and all types of model state components

grid-point space

$$\delta x = \delta x_{3D-Var} + \sum_{k=1}^K (\alpha_k \circ \delta x_k^{ens})$$

Empirical matrix A contains spectral density of the horizontal auto-correlation of α_k fields

Spatial averaging is applied on vorticity, divergence, temperature, specific humidity and log of surface pressure in order to preserve a geostrophic balance.

Extraction of the flow-dependent information from the ensemble of forecast perturbations



HIRLAM approach

raw ensemble forecast-error variances +
Schur filtering of raw ensemble forecast error correlations in grid-point space +
mix with the climatological forecast error covariances

Meteo-France approach

filtering of ensemble forecast error variances and possibly covariances by spatial averaging (in spectral and/or wavelets space)

with the precision of Loik Berre!



What makes the hybrid ensemble variational approach attractive

- ◆ Ensemble of forecast error perturbations contains flow-dependent structures;
- ◆ Ensemble of the forecast error perturbations contains anisotropic structures (both large- and small-scales are represented);
- ◆ Ensemble of the forecast error perturbations reflects relationship between large-scale forcing and meso-scale developments.

Flow situation – 21 August 2007



Data assimilation experiment (12.08.2007-24.08.2007)

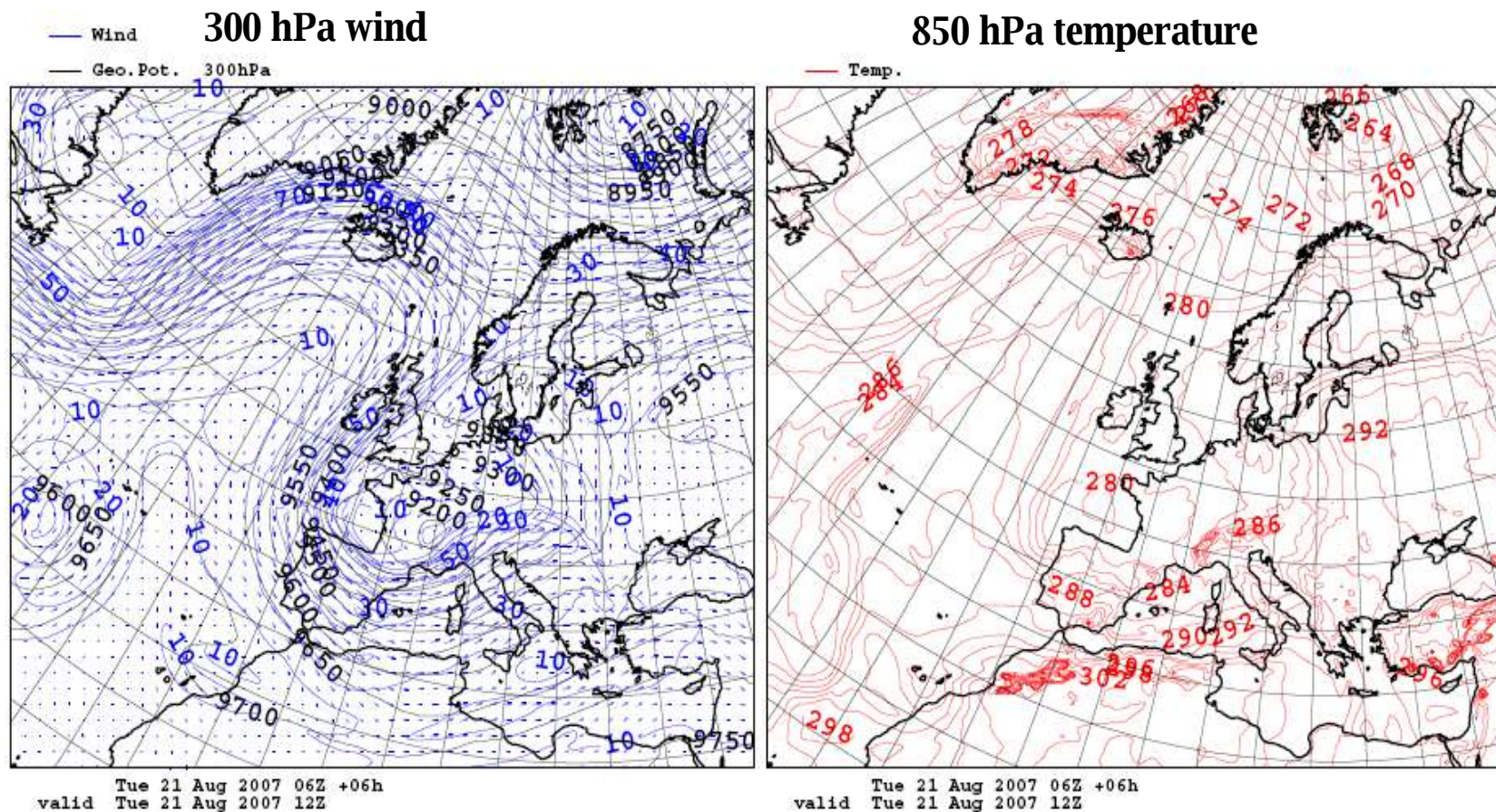
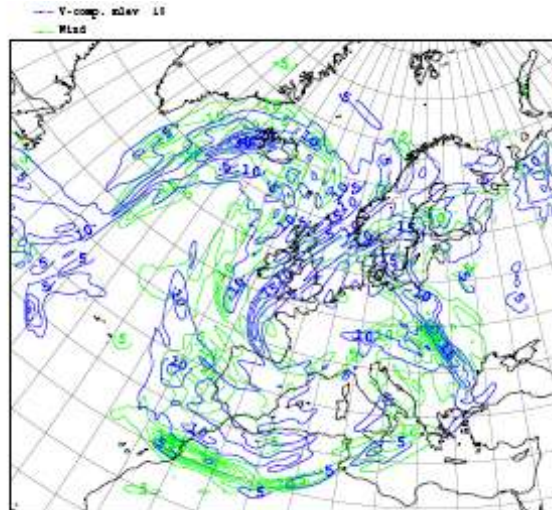


Figure 4. 300 hPa wind and geopotential (left) and 850 hPa temperature (right) taken from the background model state at 21 August 2007 06UTC + 6h; experiment based on equal static and ensemble contributions to the background error variance

Example of ensemble (ETKF) variance fields (12 members)

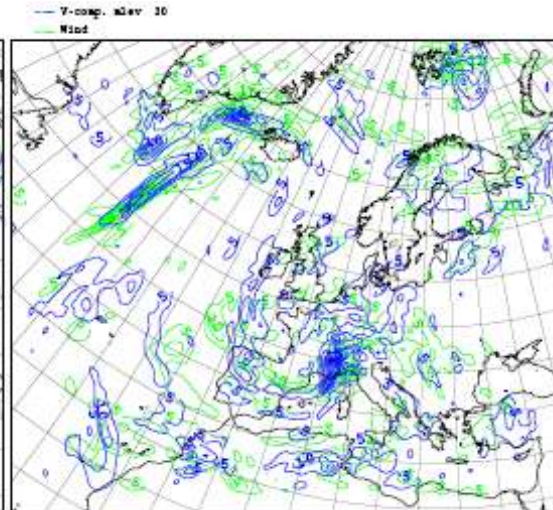


Wind
components
model level 15



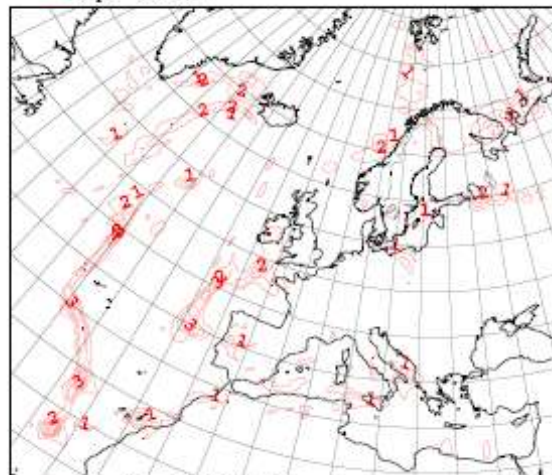
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Tue 21 Aug 2007 00Z

Wind
components
model level 20



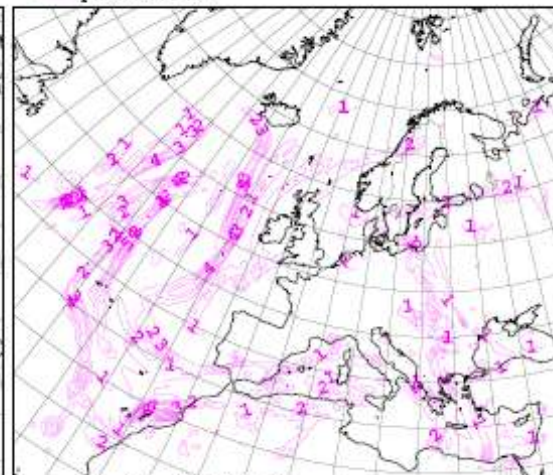
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Temperature
model level 30



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Specific
humidity
model level 30



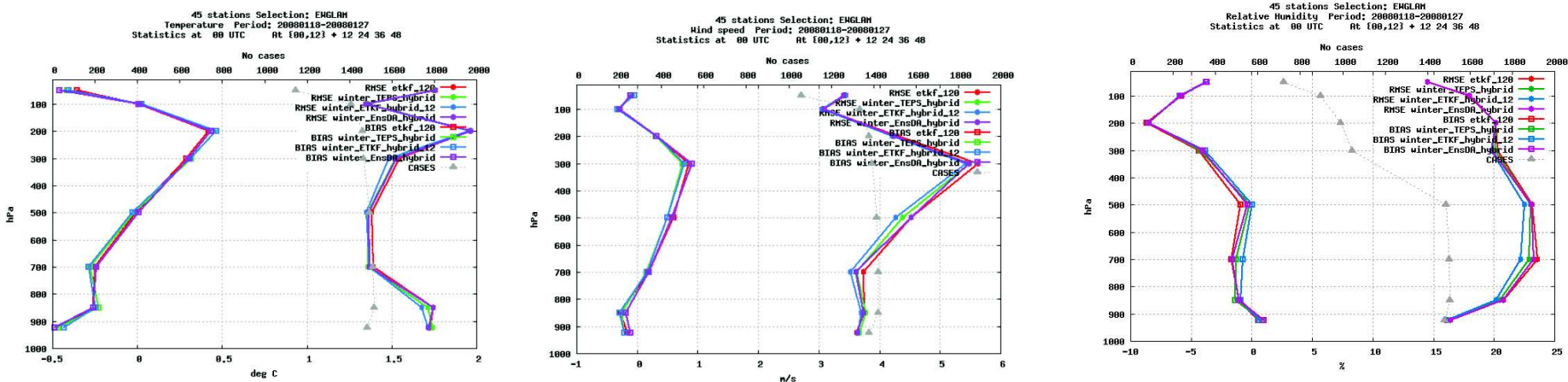
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Tue 21 Aug 2007 00Z

Figure 5. Examples of estimated background error variances based on the ensemble of +6h forecast valid at 21 August 2007 12UTC from the hybrid data assimilation experiment hybrid6_diag. Wind components at model level 10 (upper left), wind components at model level 20 (upper right), temperature at model level 30 (lower left) and specific humidity at model level 30 (lower right).

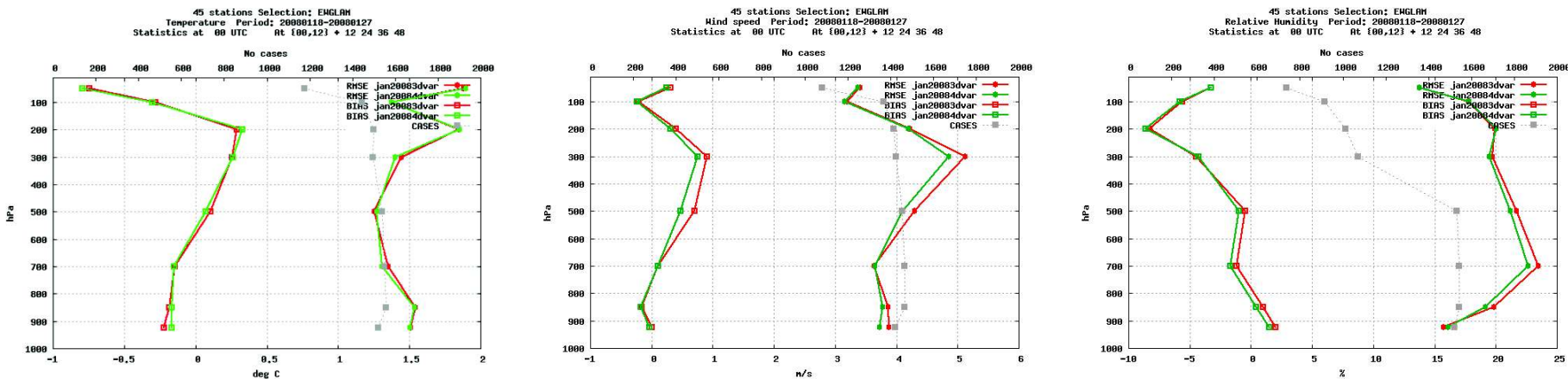
Verification scores : “winter case” 17Jan2008-27Jan2008



3D-Var versus hybrid approach (*ETKF*, *EnsDA*, *TEPS*)



3D-Var versus 4DVAR approach



Temp.

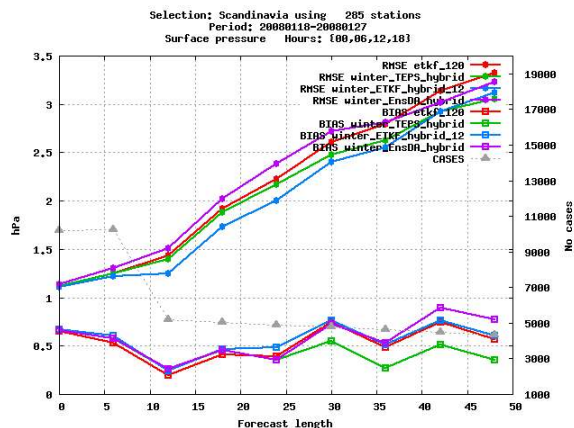
Wind speed

Rel. humid.

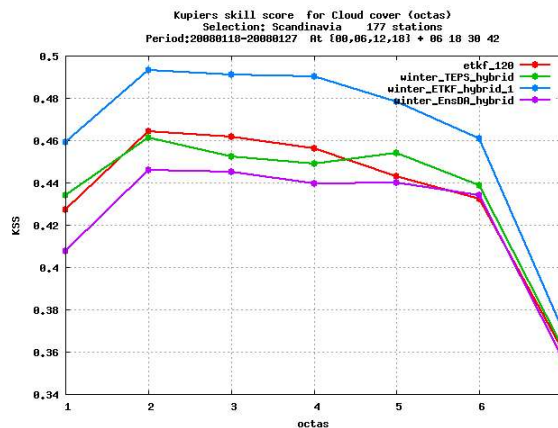
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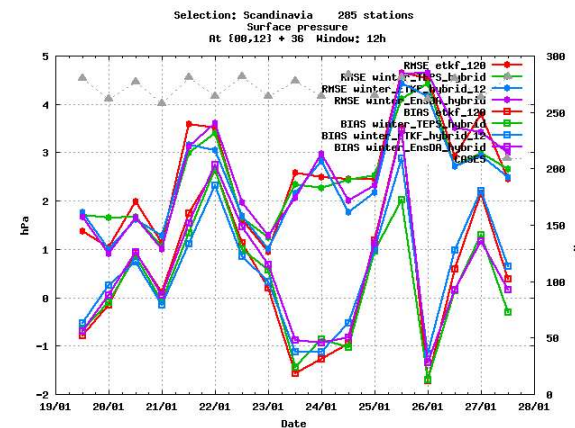
3D-Var versus hybrid approach (*ETKF*, *EnsDA*, *TEPS*)



Psm1



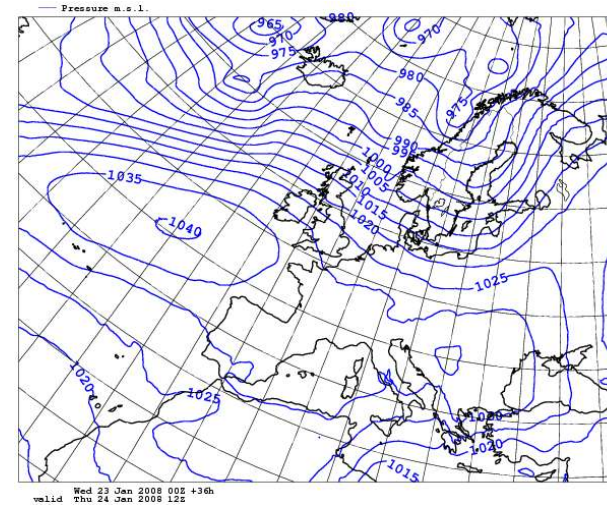
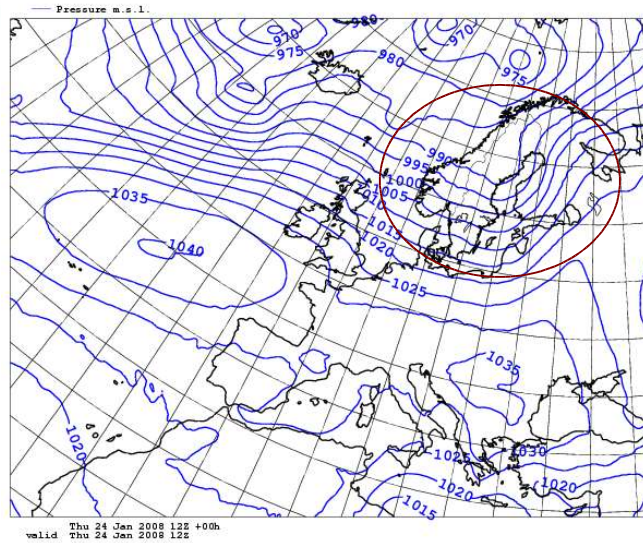
Cloud. Kuiper score



Surface pressure TS

3D-Variational data assimilation runs in non-optimal setup in these experiments (severe cold bias of SST). Experiments should be redone with optimal settings!!!

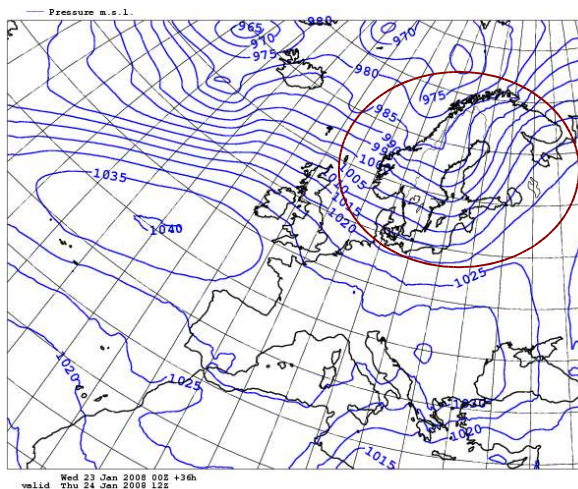
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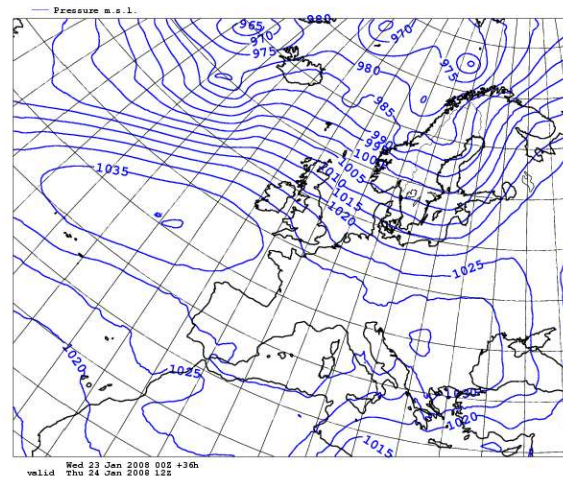
3DVAR

3d-Var analysis

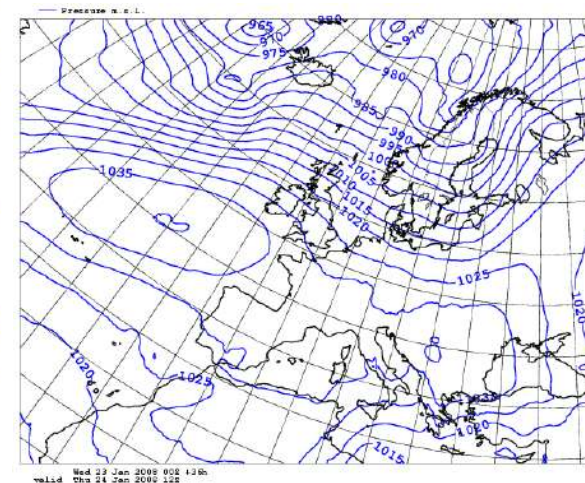
versus +36 h forecasts



3DVAR+ETKF



3DVAR+EnsDA



3DVAR+TEPS

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22 Jan 00 UTC (front), mlev 28



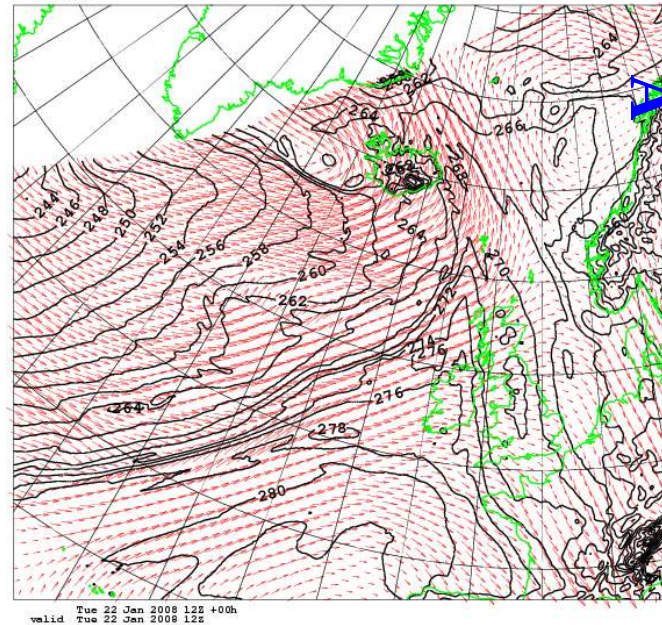
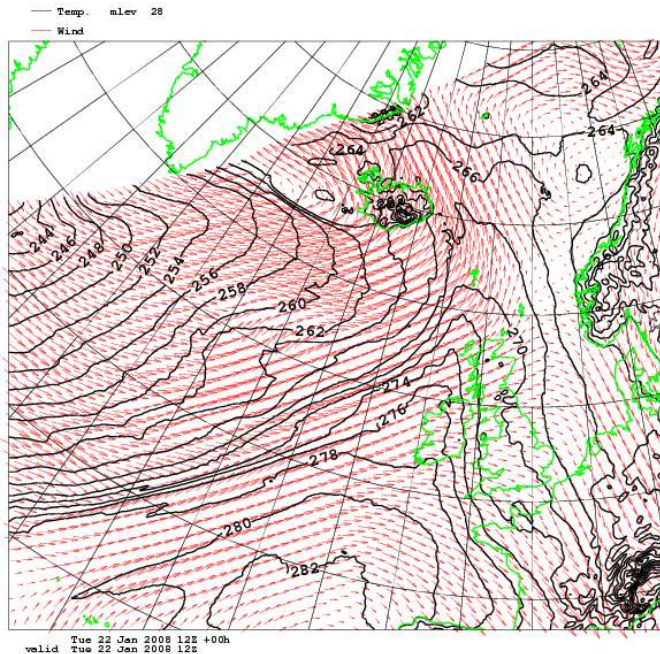
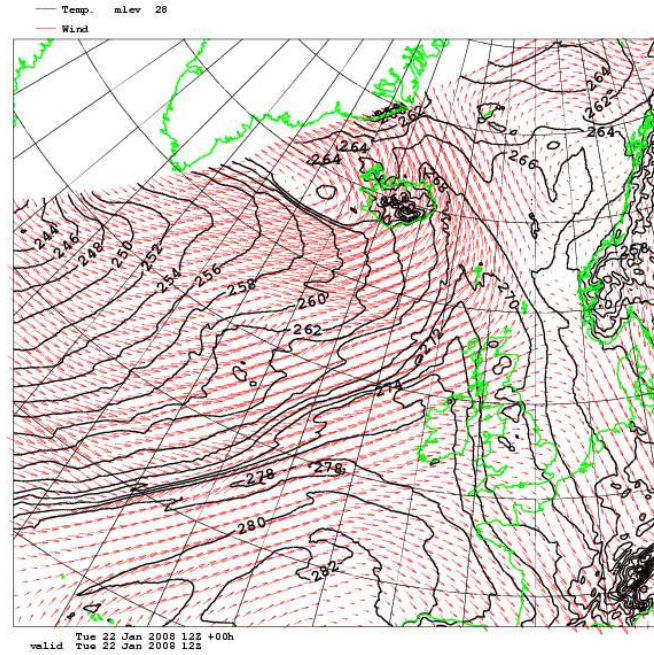
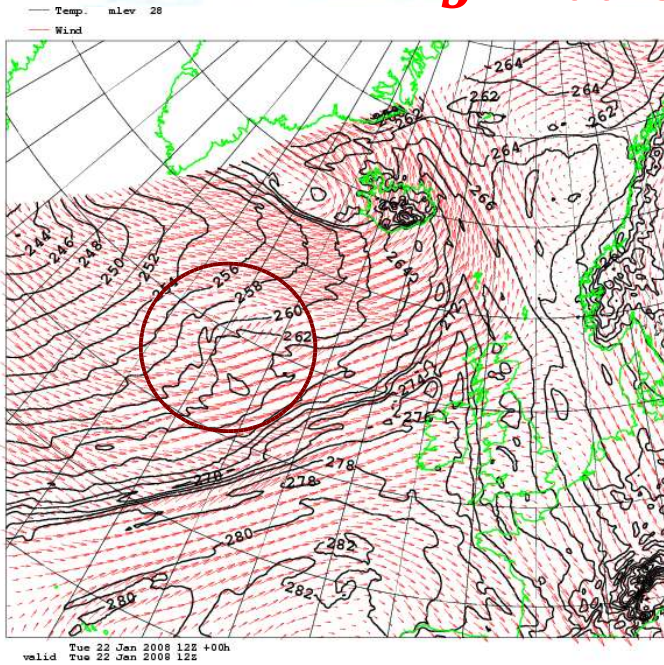
3DVAR + TEPs

3DVAR + EnsD

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3DVAR

3DVAR + ETKF



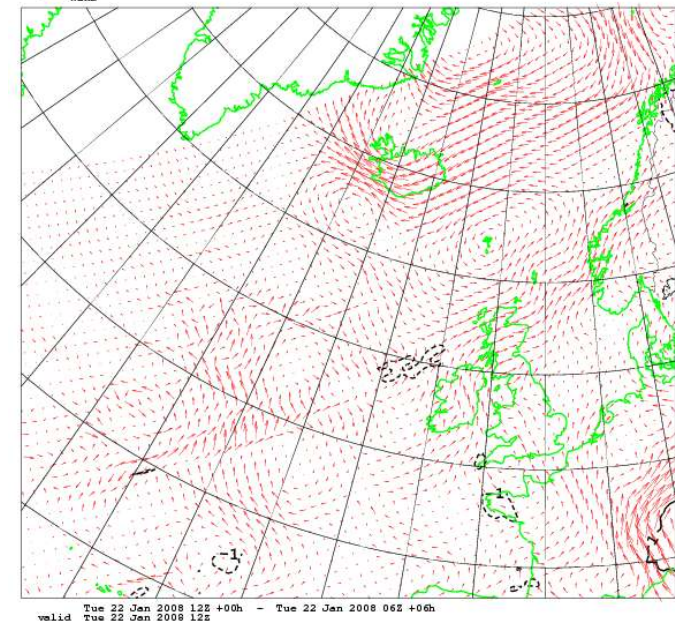
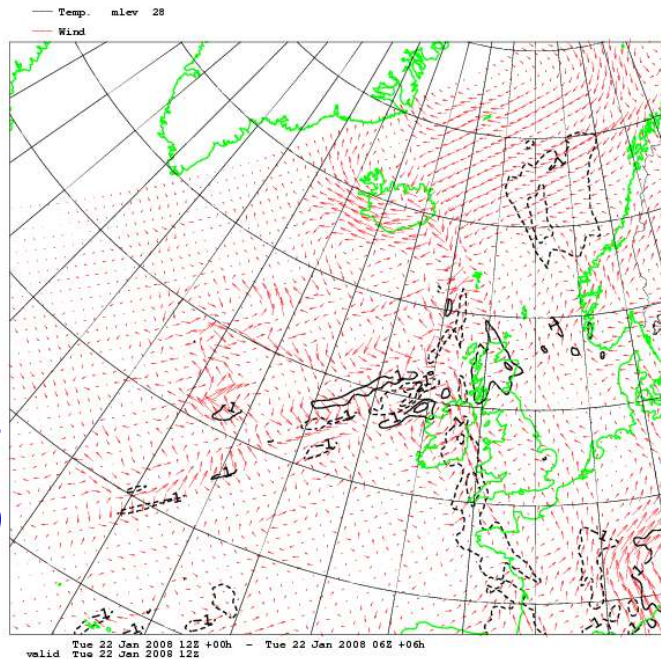
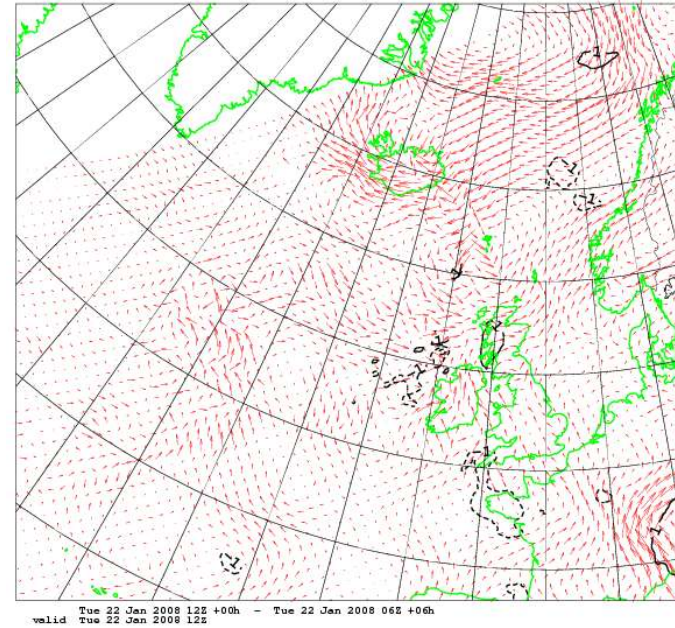
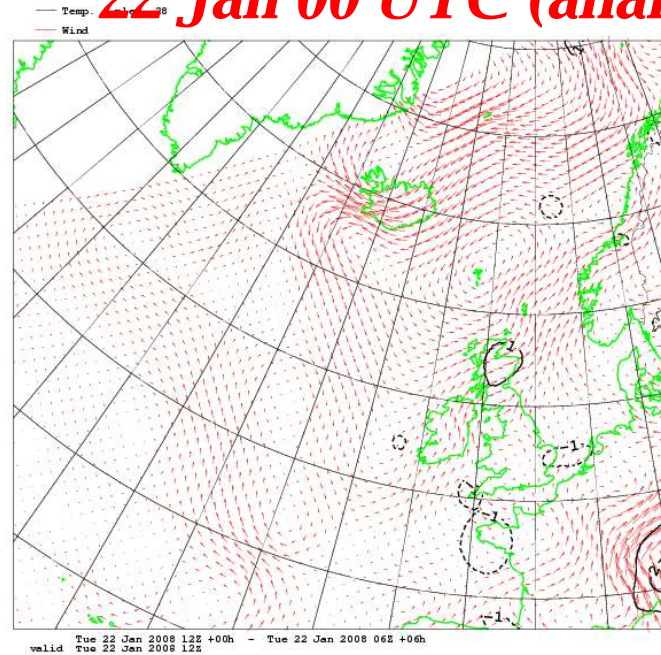
22 Jan 00 UTC (analysis incr. along front), mlev 28



3DVAR

3DVAR+TEPS

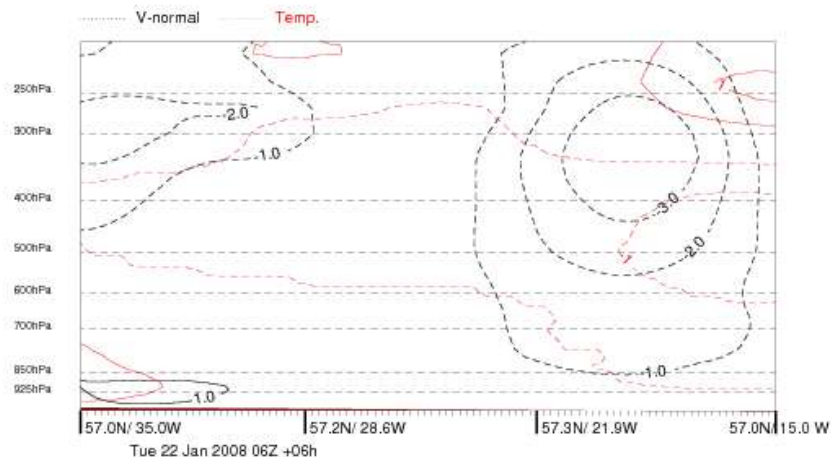
3DVAR+EnsDA



22 Jan 00 UTC (cross-section of across front of analysis incr.)

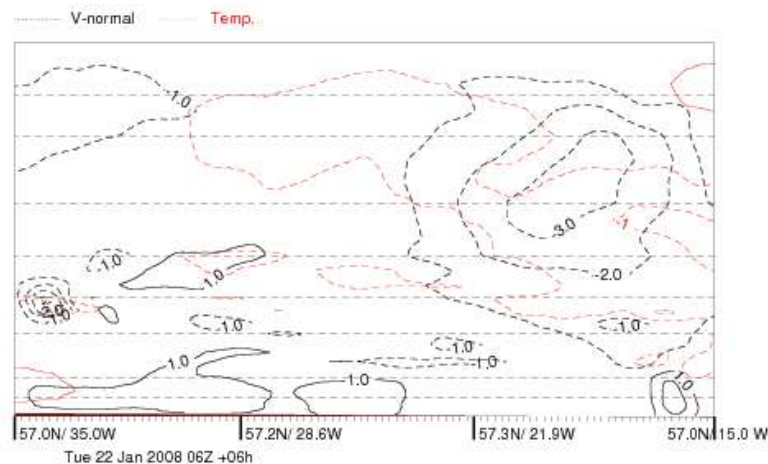


22 Jan. 2008 12UTC + 0h valid time: 22 Jan. 2008 12UTC



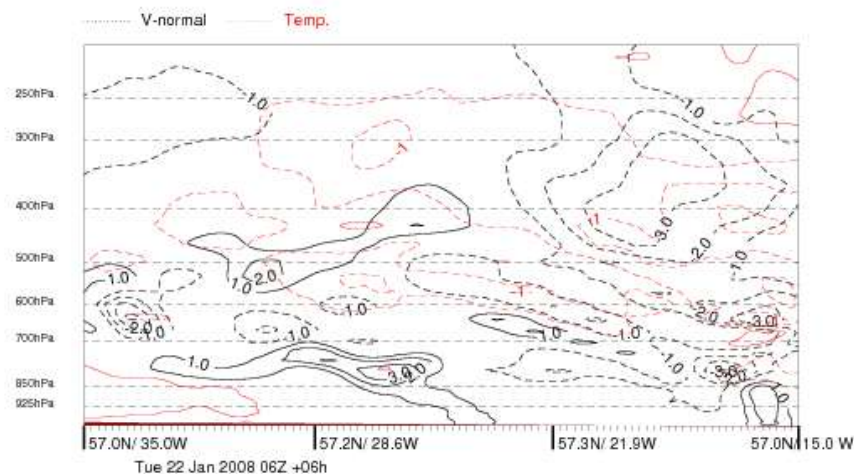
3DVAR

22 Jan. 2008 12UTC + 0h valid time: 22 Jan. 2008 12UTC



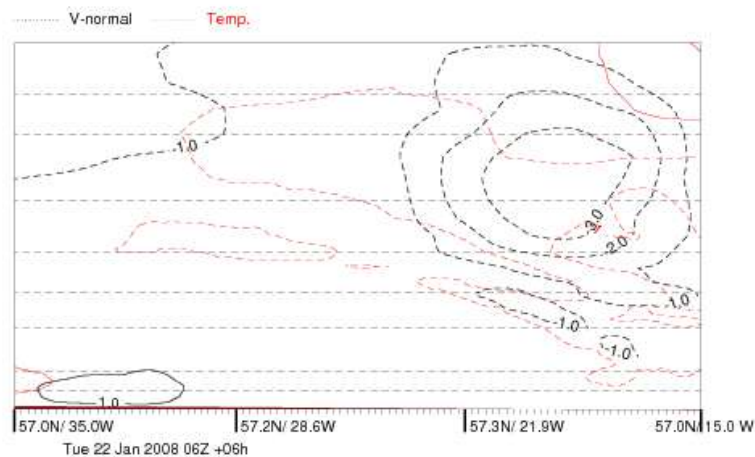
3DVAR+TEPS

22 Jan. 2008 12UTC + 0h valid time: 22 Jan. 2008 12UTC



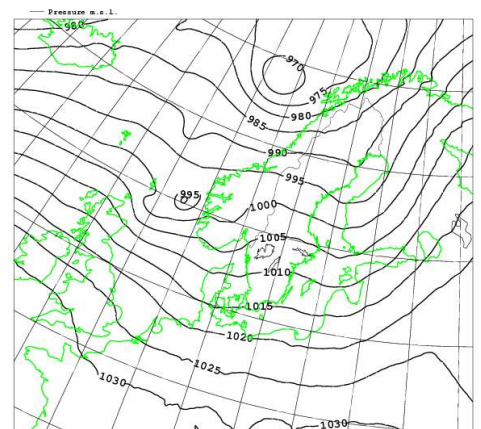
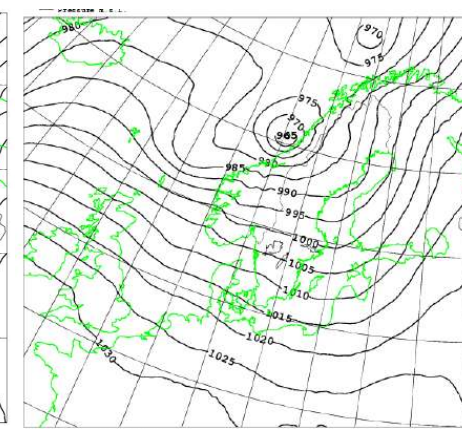
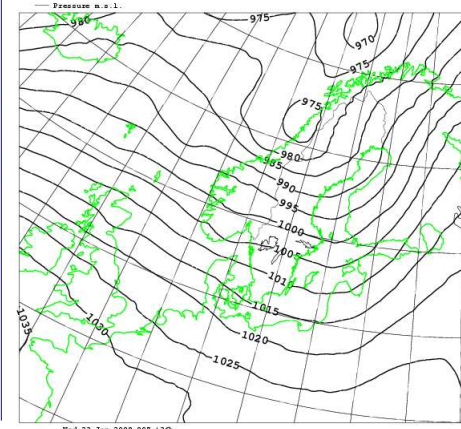
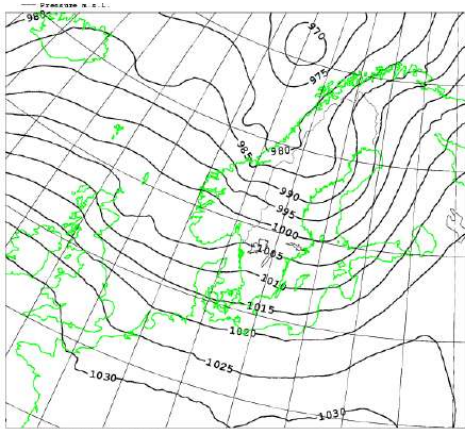
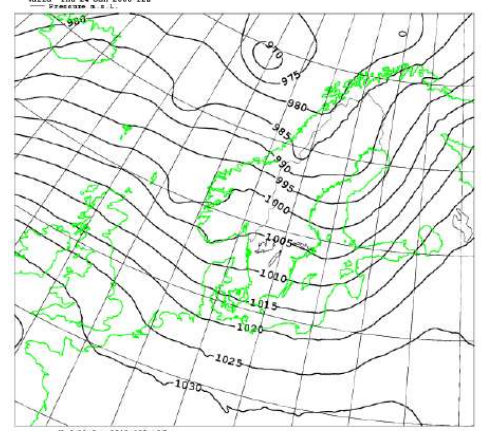
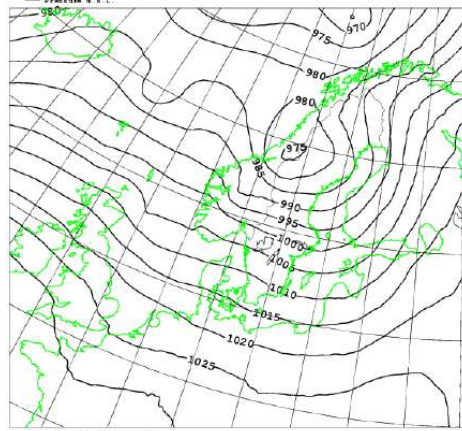
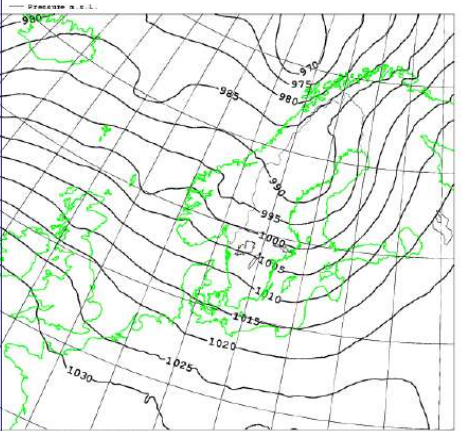
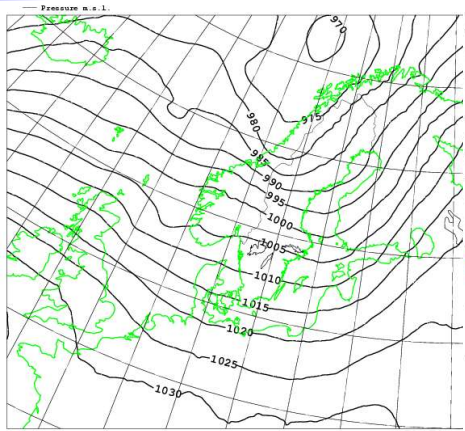
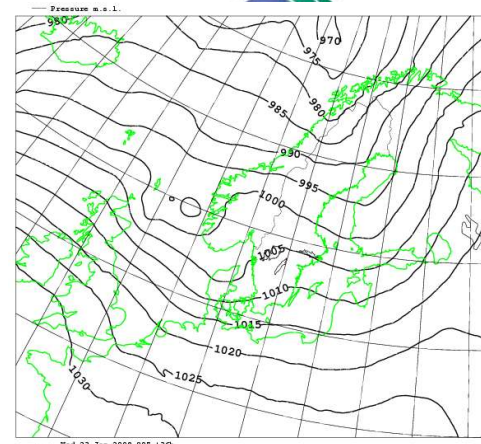
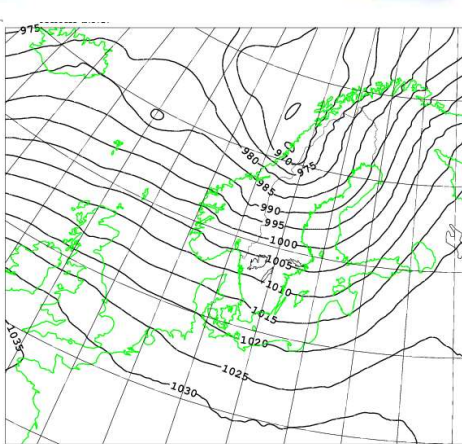
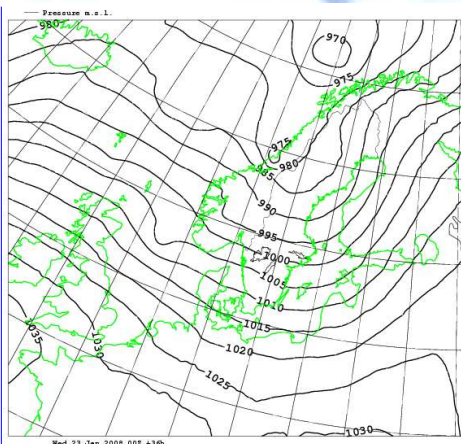
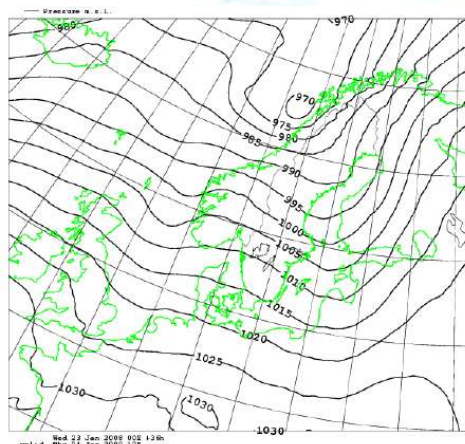
3DVAR+ETKF

22 Jan. 2008 12UTC + 0h valid time: 22 Jan. 2008 12UTC

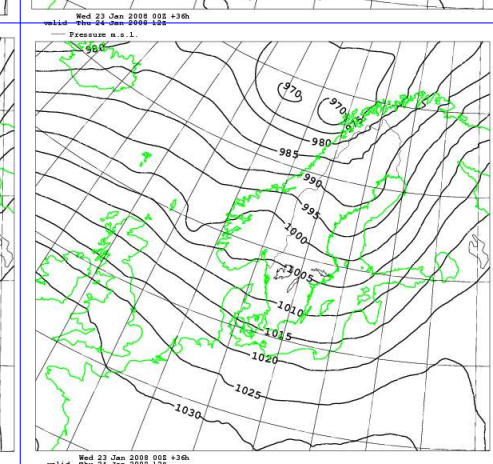
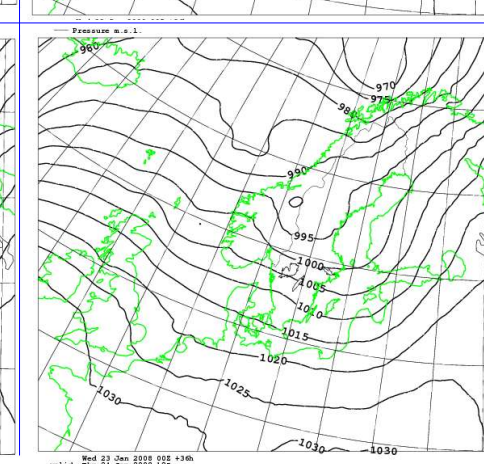
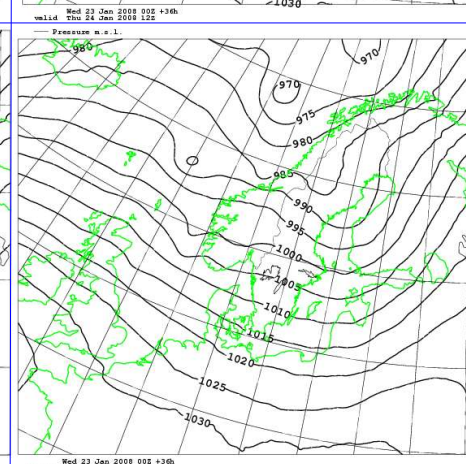
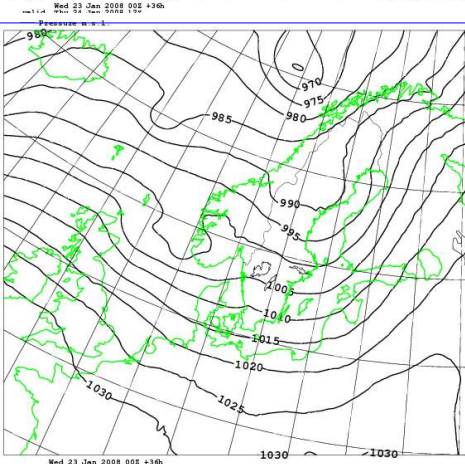
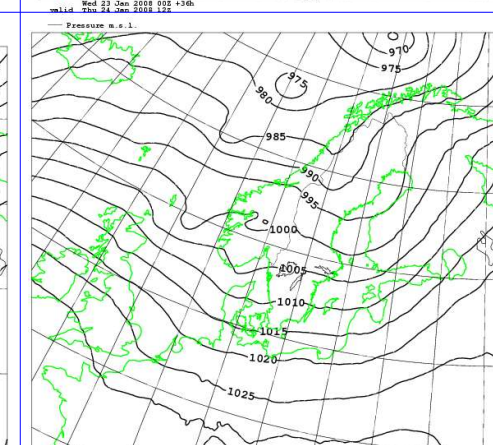
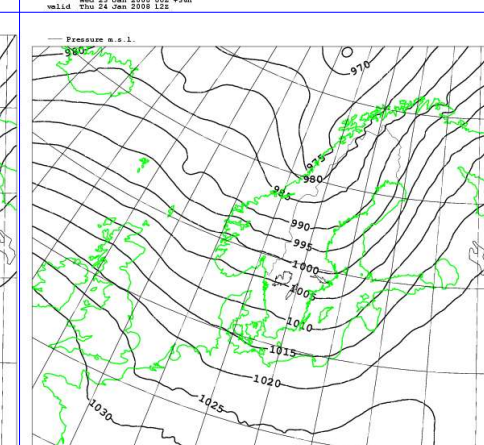
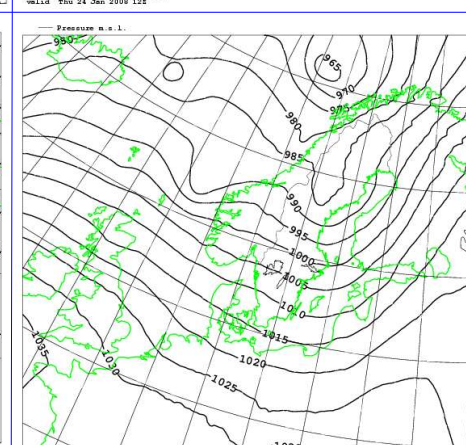
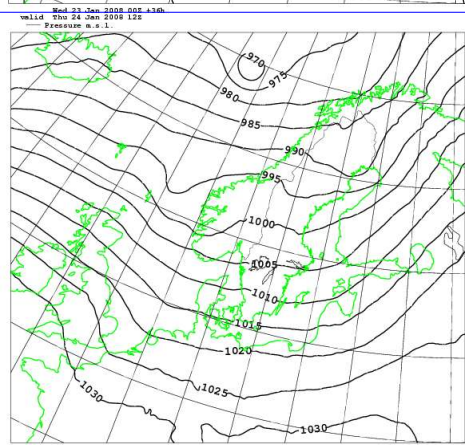
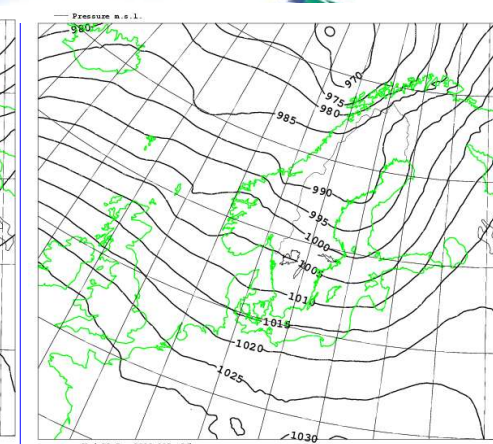
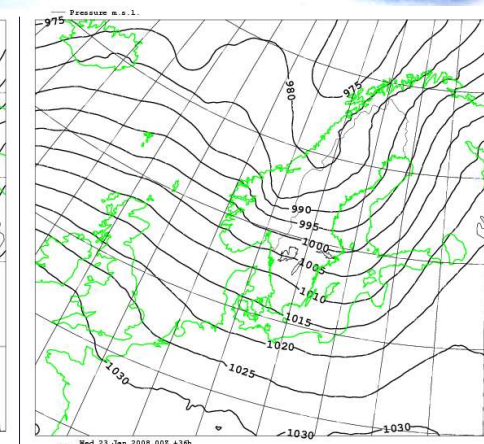
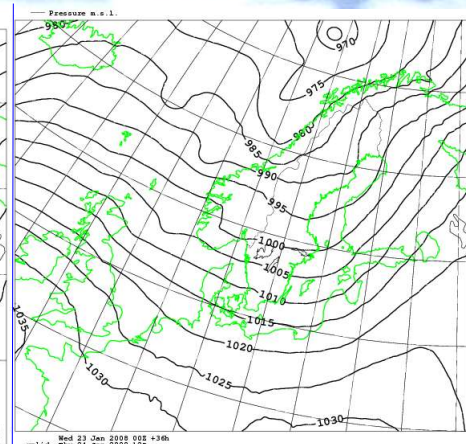
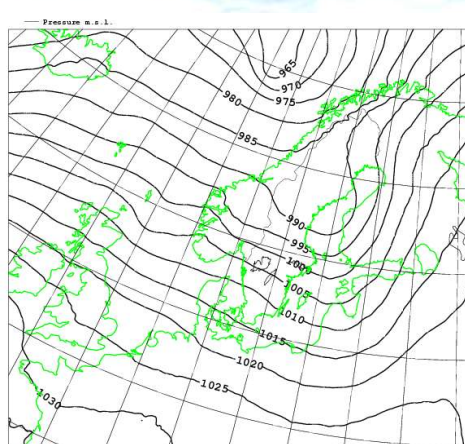


3DVAR+EnsDA

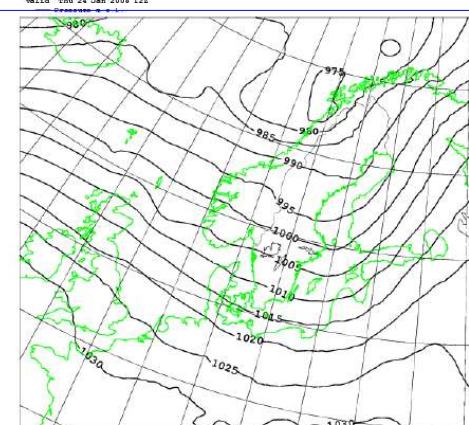
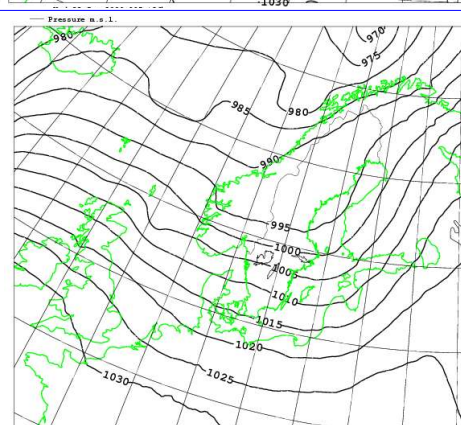
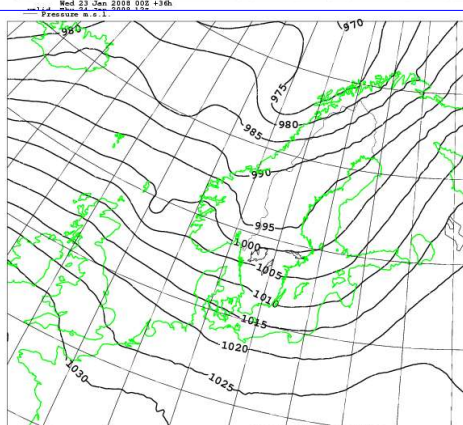
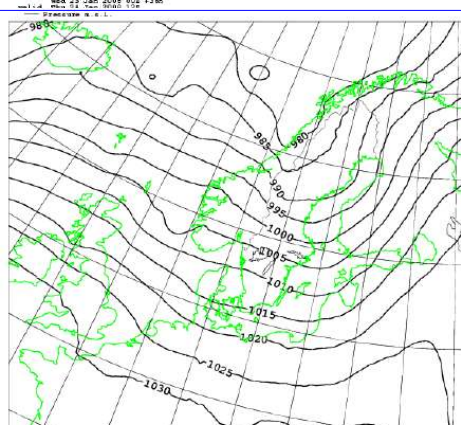
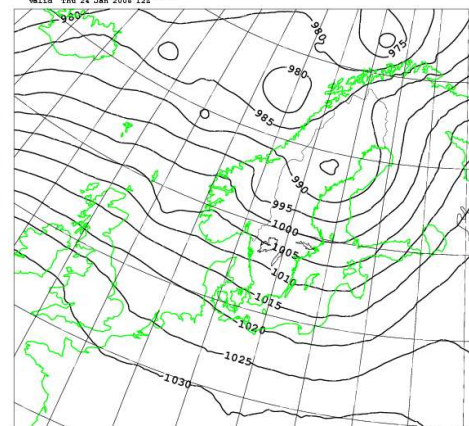
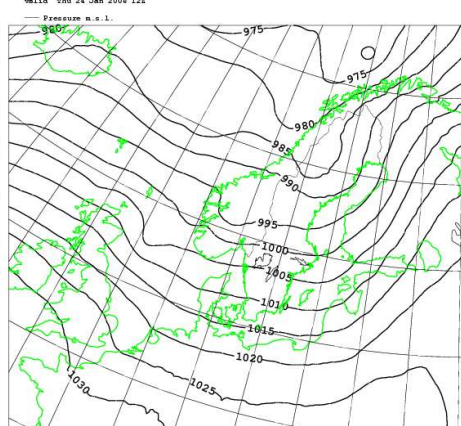
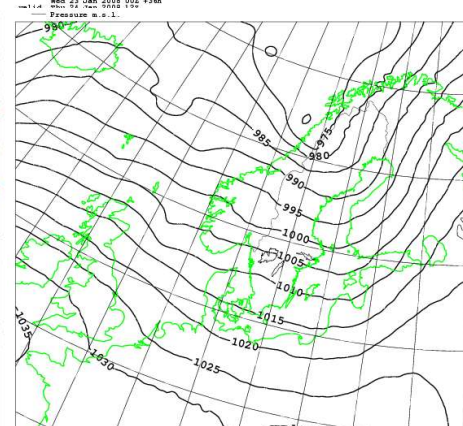
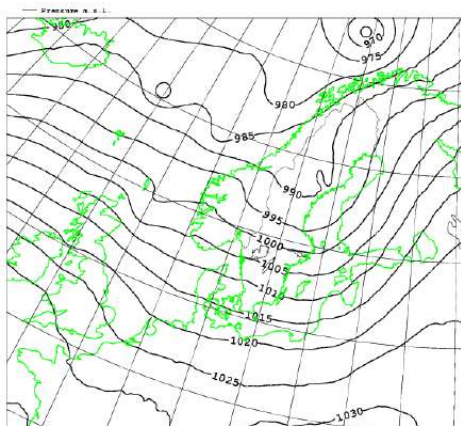
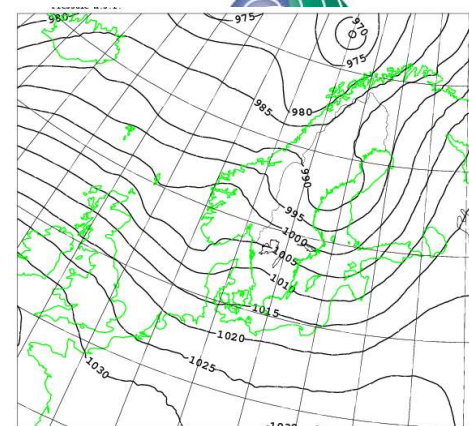
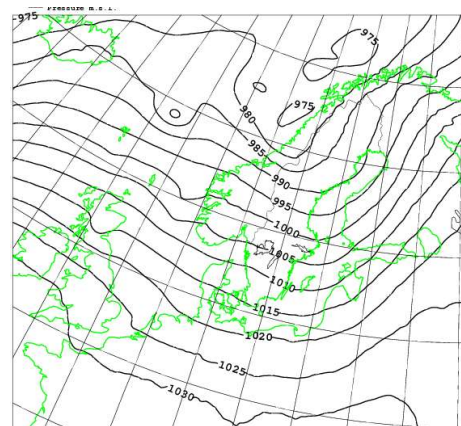
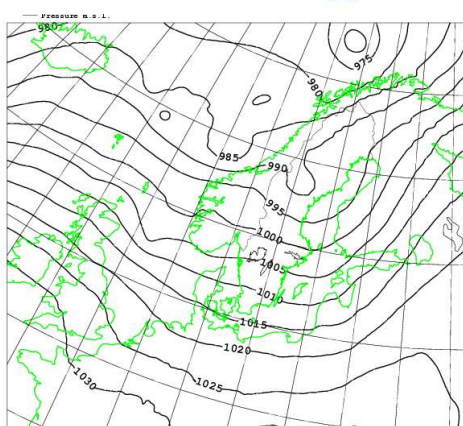
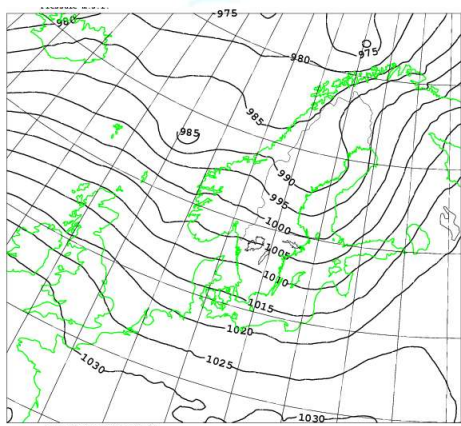
+36 h EuroTEPS valid at 24 Jan 12 UTC



+36 h ETKF EPS valid at 24Jan 12 UTC



+36h EnsDA EPS valid at 24 Jan 12 UTC





Conclusions

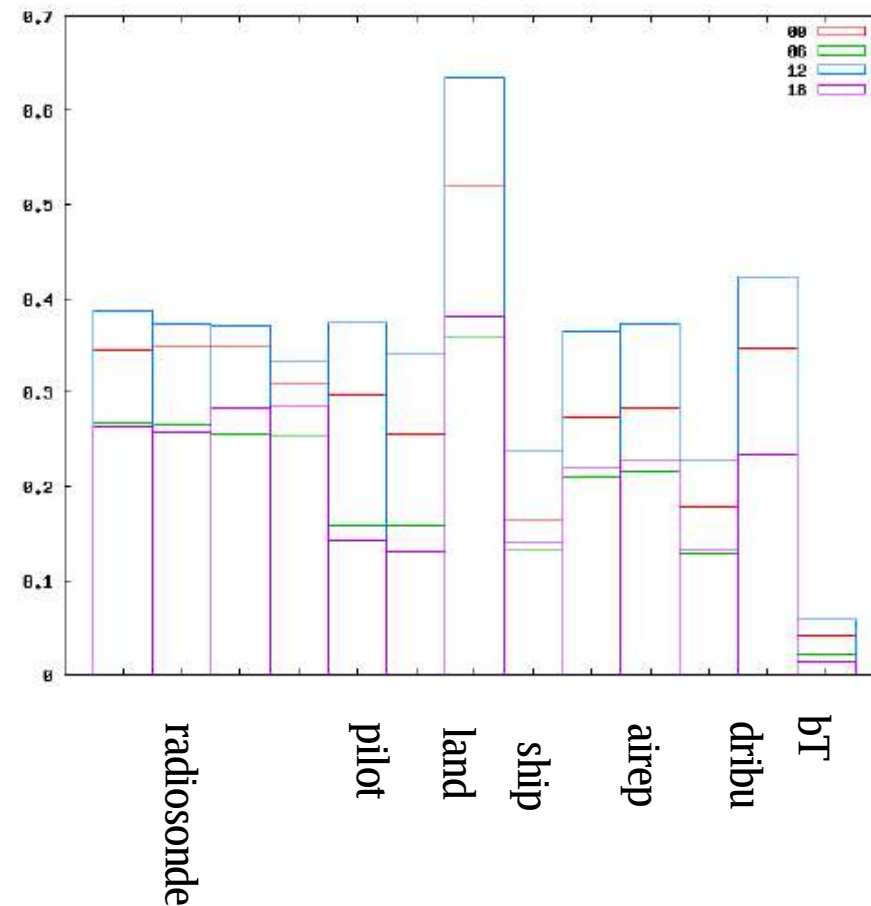
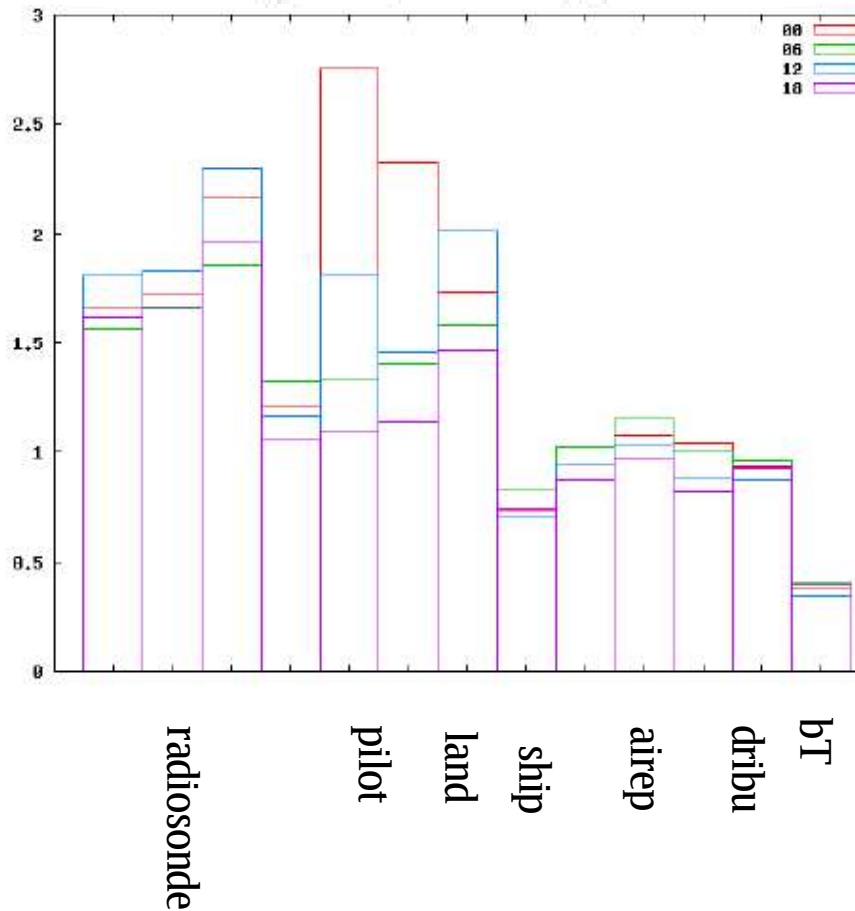
- The flow-dependent structures reflected in ensemble based forecast error covariances improves 3D-Var performance.
- All three ensemble generation strategies (EuroTEPS, ETKF and EnsDA) are useful for this purpose.
- Improvement of the standard verification scores for the hybrid ensemble variational data assimilation scheme over pure 3D-Var is of the same magnitude as of the 4D-Var over the 3D-Var.
- Further tuning is needed to improve performance of EuroTEPS and in particular EnsDA based hybrid scheme.
- The impact of the increased ensemble size on the EPS performance and the usefulness of the hybrid 4D-Var ensemble data assimilation scheme will be investigated.

Relative squared innovation variance versus relative spread



$$d_i^2 := \sum_{j=1}^P \left(\frac{y_j - \mathcal{H}_j^T x_i^b}{\sigma_{o,j}} \right)^2 \approx \alpha_i \sum_{n=1}^{N-1} \lambda_n^{N-1} + p$$

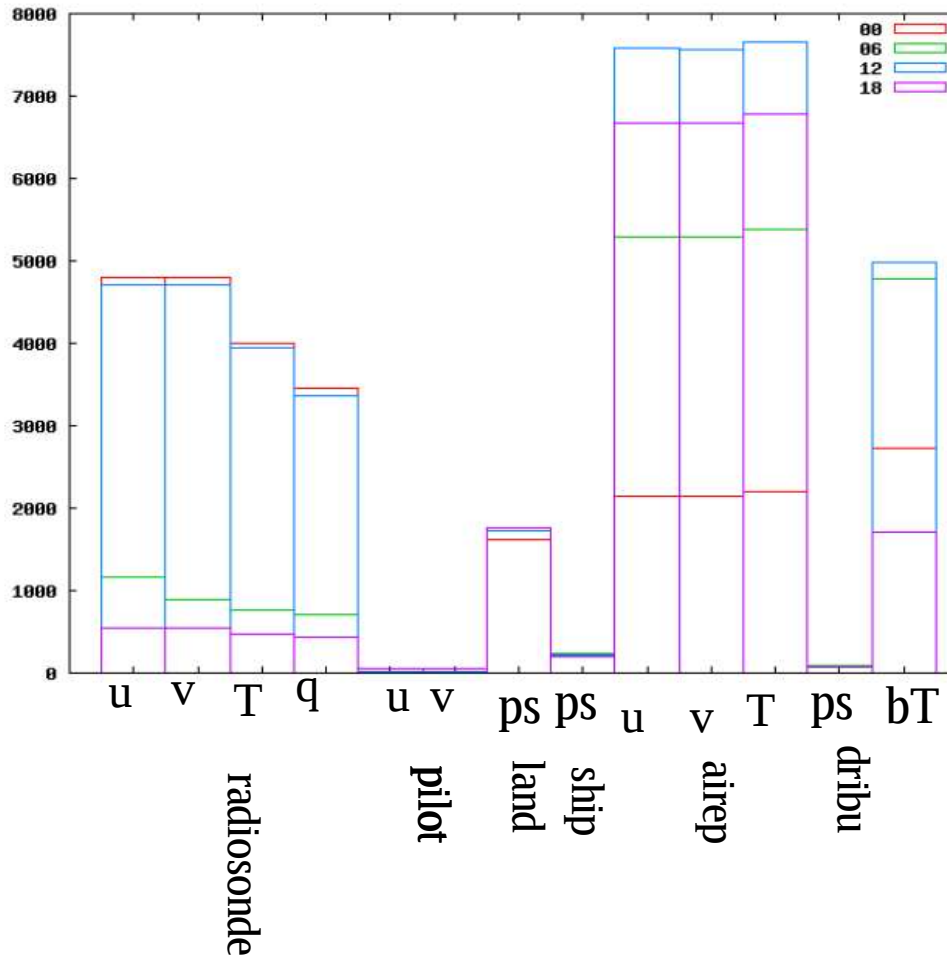
$$\text{trace}(B_{\text{ens}}^f) = \frac{1}{P} \sum_{j=1}^P \frac{\text{Var}(\mathcal{H}_j(d\mathcal{X}^f))}{\sigma_{o,j}^2}$$



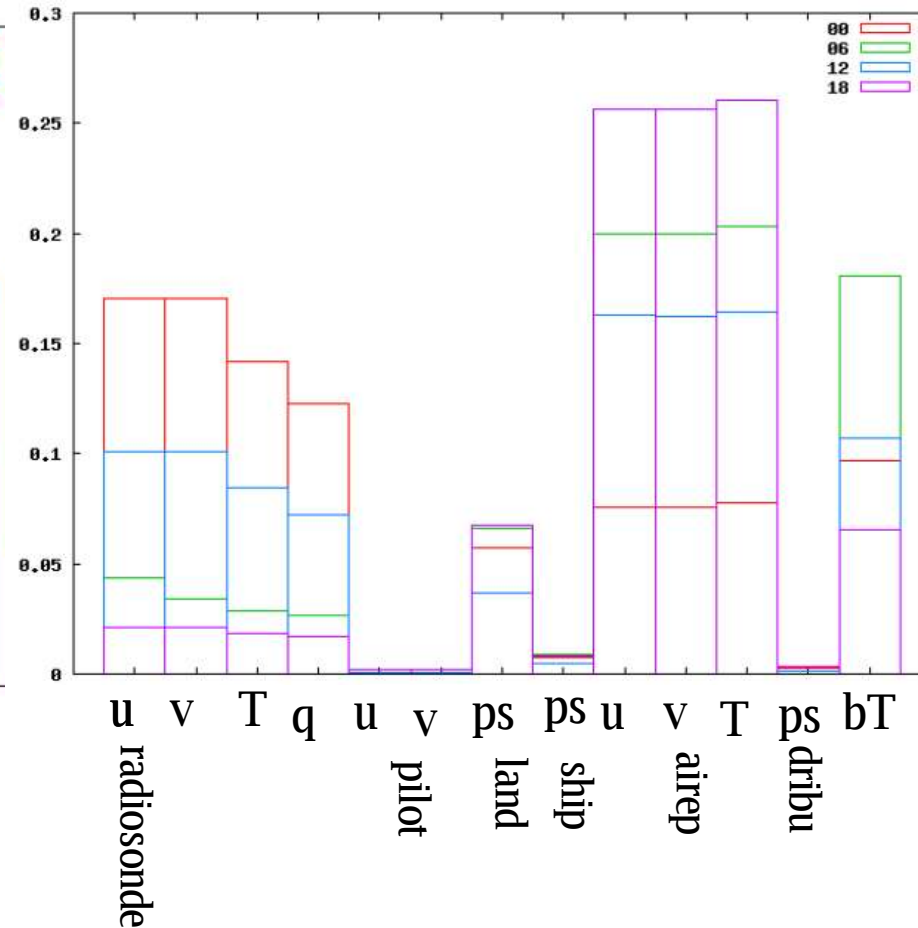
Earth observing system



absolute amount of observations



relative amount of observations





Thank You

for

Attention !..