

### Lambertian snow emissivities

#### Summary of recent work and application to Antarctica

R. Chawn Harlow, CONCORDIASI Workshop, 30 March 2010

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#### **CLPX-II** Collaborators

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- Jared Entin of NASA Terrestrial Hydrology Program
- FAAM and Met Office aircraft crew taking part.



#### Introduction

- Outcomes of CLPX-II and POLEX campaigns
- Implications for efforts in Antarctica.



#### Introduction

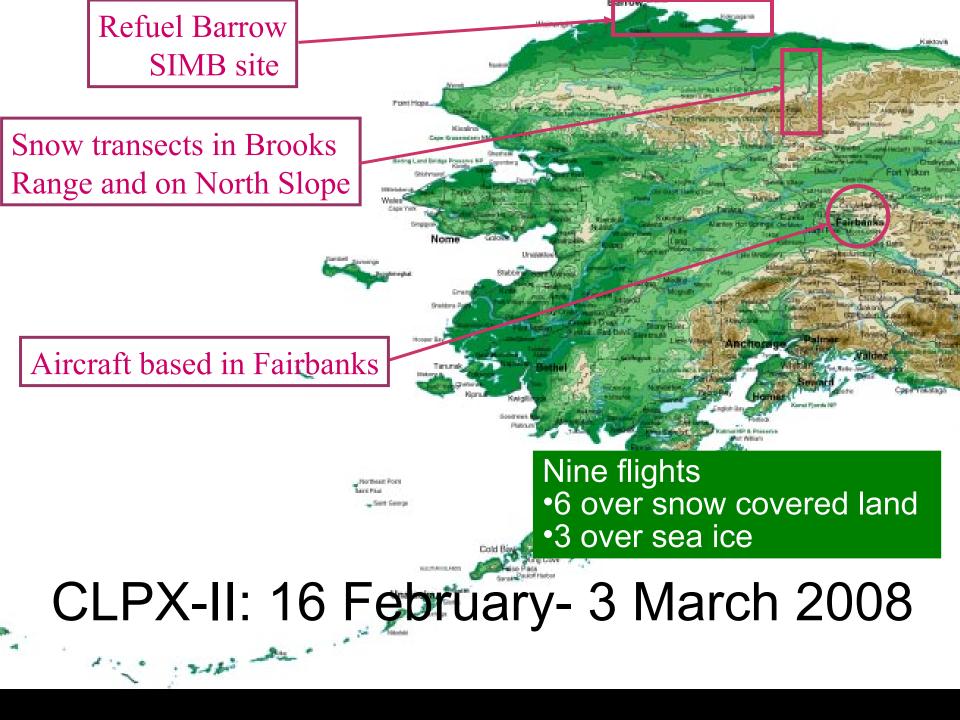
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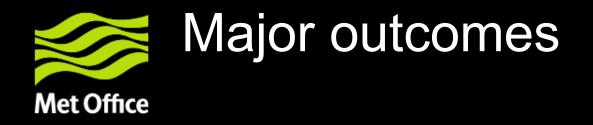
#### Motivation

- Hope to assimilate satellite microwave (AMSU and MHS) data over land and sea ice
  - Provide atmospheric temp and humidity profile information where there are holes in the conventional profile data.
- Need surface emissivity information derived from models and satellite observations.
- New snow module to be included in future Met Office land surface component JULES. Future data assimilation opportunities?
- First need to validate snow radiative transfer models in 20 to 200 GHz range.

#### POLEX Campaign March 2001

- Five flights over sea ice
  - Three over pack ice
  - Two over Marginal Ice Zone
- One flight over snow covered land

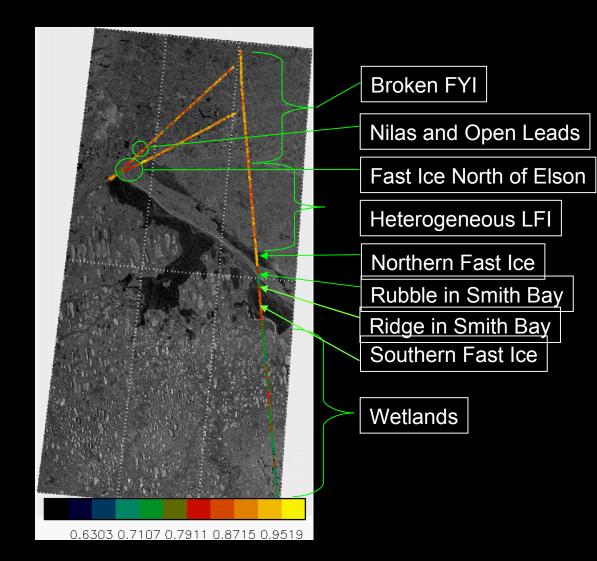




- AMSU-B emissivities at 89, 157 and 183 GHz.
  - mm-waves
- Snow surfaces are Lambertian at 183 GHz (Harlow, 2009).
- Sea ice mm-wave Lambertian emissivities can be related to sea ice type through typical snow cover stratigraphic sequences (Harlow, 2010a).
  - Coupled atmosphere-snow-ice-ocean models may lead to better analysis of snow water content, ice type, and atmospheric temperature and humidity via assimilation of microwave emissivities on the 20-200GHz range.
- mm-wave Lambertian snow emissivities may be modeled with MEMLS if mm-scale surface roughness is taken into account (Harlow, 2010b).
  - Extension of MEMLS to ~200GHz.
  - To be confirmed in future campaigns.

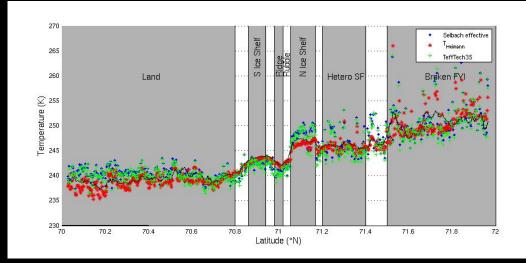


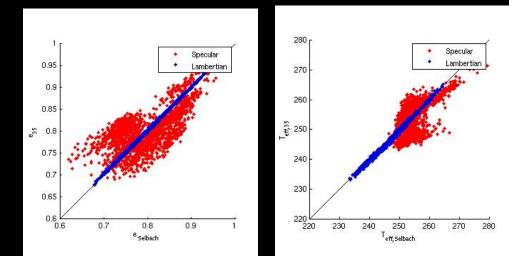
#### Snow surfaces are Lambertian at 183 GHz





- Two estimates of 183 GHz effective temperature and emissivity based on microwave T<sub>B</sub>'s
- Close agreement under Lambertian scattering
  - RMSD  $\cong 0.5~K$
  - RMSD = 4.4 K for specular.
- Confirmed with data from 12 flights





# Met Office Impact of diffuse scattering in the 20-200 GHz range

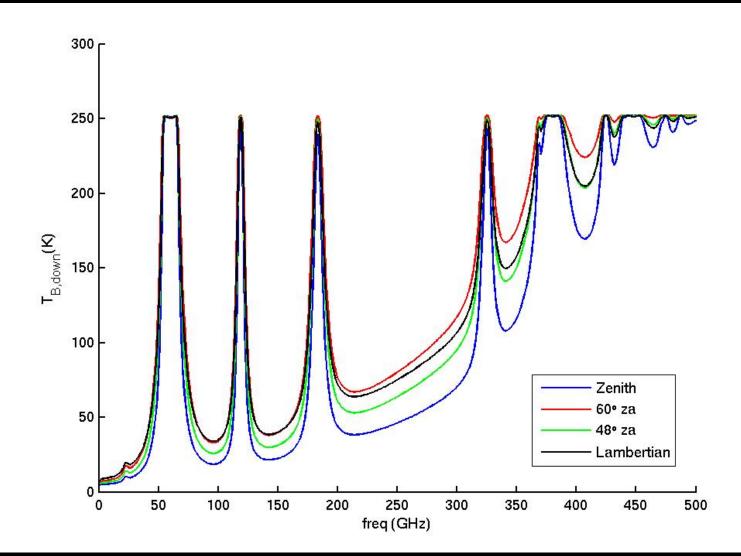
- The degree to which Antarctic surface is Lambertian has been studied at 37 and 50 GHz by Rosenkranz and Matzler (2008).
  - Scattering characterized as linear combination of Lambertian and specular scattering with weight for Lambertian given by  $\rm A_L$
  - $A_L \sim 0.7$  to 1.0 but spatially variable
- Guedj (2009) discussed emissivities retrieved with various degrees of Lambertivity (A<sub>L</sub>)
  - Suggested seasonal variation from Lambertian in the winter to semi-Lambertian or specular in the summer.
- Goal now is to quantify the impact of Lambertian scattering on the upwelling radiation at the surface.



- Use ARTS to generate downwelling  $T_B$ 's at 10 zenith angles ( $\theta$ ) between 0 and 90°.
- Frequencies 1 to 1000 GHz in 1 GHz steps
- B345 atmospheric conditions– Cold and dry.
- Integrate  $T_B$ 's over the Lambertian BSF.
- This approximates the effective downwelling seen by a Lambertian surface.
- Specular surface reflects like a mirror so nadir measurements see reflections of the zenith  $\rm T_{\rm B}$
- Compare Lambertian  $T_B$ 's to those from a specific direction  $\theta$  in the sky for a range of optical depths
- Compare upwelling Lambertian and specular  $T_{B}$ 's.
- Summarize differences in terms of acceptable uncertainties for sounding.

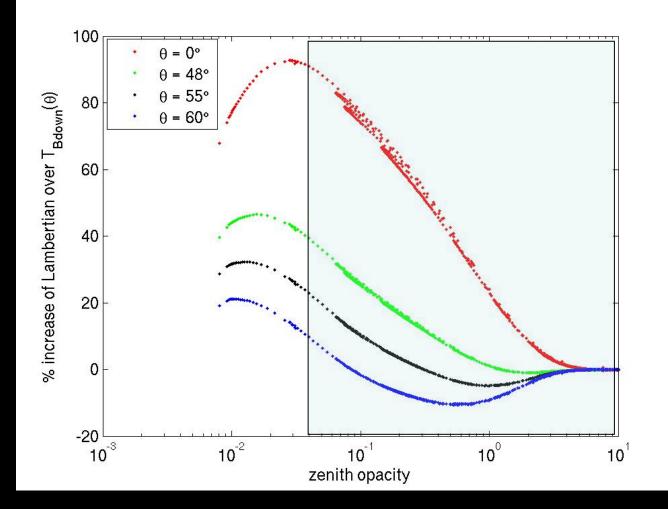


#### Lambertian effective downwelling



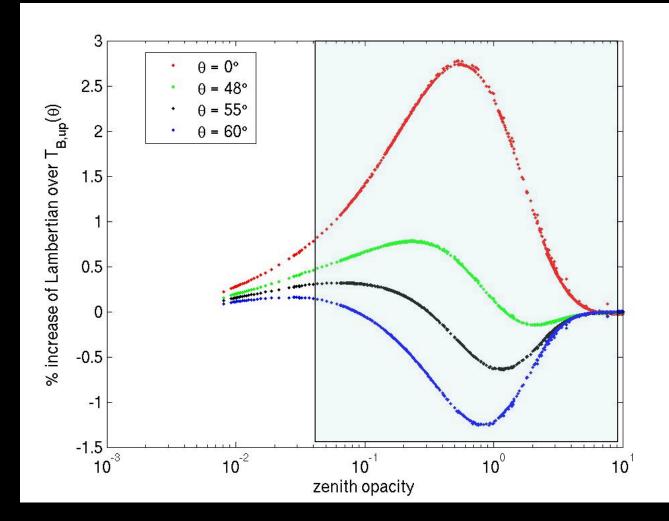


## Impact of Lambertian scattering a function of zenith opacity



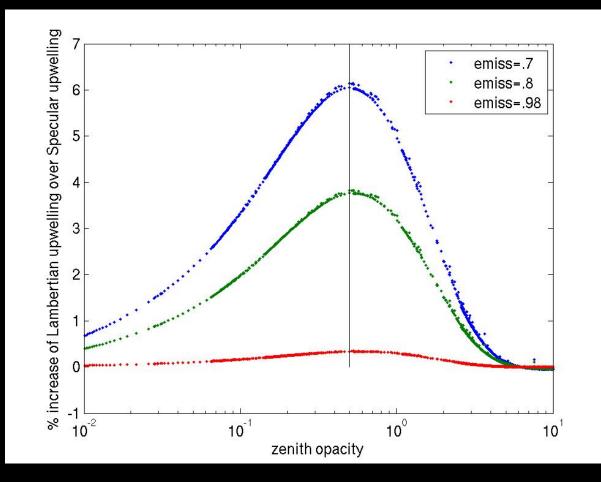


### Difference between Lambertian and Specular upwelling when emiss=.85

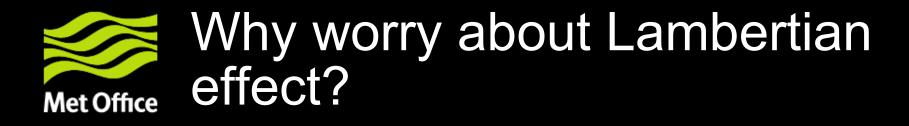




## Effect of Lambertian scattering as a function of zenith opacity



Max effect at Zen op of 0.5



- Lambertian effect is a function of zenith opacity: Change in spectral shape of retrieved surface emissivity.
  - Retrieval of geophysical variables must ingest any error in emissivity.
- These differences in emissivity introduce ambiguity of ~5-10 K in upwelling TB which may contaminate any retrieved profile using radiance assimilation at TOA.



# Influence of Lambertian scattering on RT

- Where emissivities are relatively small there are large impacts on the radiances and thus retrieved emissivities particularly at opacities of 0.5.
- Lambertian scattering should have a strong impact for
  - snow covered surfaces in the microwave
  - desert surfaces in the IR.



- Microwave emissivities are likely to vary with frequency and elevation in Antarctica from ~0.7 to near unity (Picard et al, 2009; Guedj et al, 2009; etc.).
- For near nadir instruments (AMSU-A/B & MHS) the degree of diffuse scattering needs to be taken into account.
  - Up to 15 K error in upwelling at surface otherwise.
  - Up to 3 K for observations at  $55^{\circ}$  incidence (SSMI, SSMI/S, etc).
  - Both are larger than acceptable for sounding purposes (English, 2008).
- The seasonal variation needs to be investigated as well as the Lambertivity at frequencies lower than 183 GHz.



- Added capabilities from ISMAR
  - Suites of channels on 118 and 325 GHz water vapour lines will be useful in evaluating A<sub>L</sub> at other frequencies.
- TAIGASEE in Finland
  - Nov 2012, Jan and Mar 2013
  - Evaluate performance of MEMLS over a season.
- Sea ice campaign 2013/14 in Alaska(?)



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- [5] S. Guedj et al, "Toward a better modeling of surface emissivity to improve AMSU data assimilation over Antarctica," *presented at 2<sup>nd</sup> Workshop on Remote Sensing and Modeling of Surface Properties, Meteo France, Toulouse, France, 9-11 June 2009.*
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