

Reconnecting cloud representations in Alaro-1

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Xu-Randall's 1996 cloud fraction

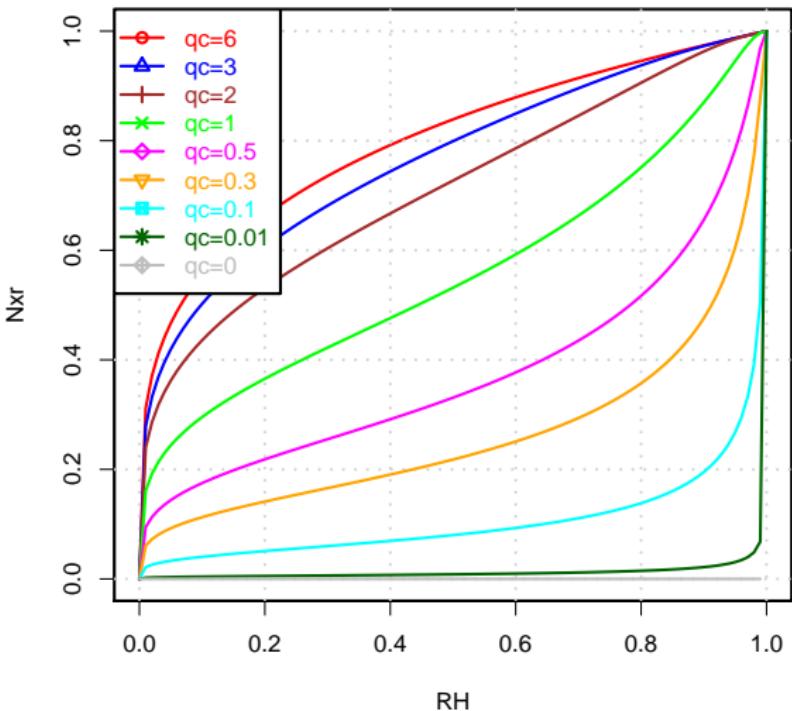
$$N = (RH)^r \left[1 - \exp \left\{ - \frac{\alpha \overline{q}_c}{(\overline{q}_s(1 - RH))^{\delta}} \right\} \right]$$

$$\alpha \sim 100; r \sim 0.25; \delta \sim 0.49$$

- mean \overline{q}_c etc. over large grid boxes (64km)
- $RH = \frac{\overline{q}_v}{\overline{q}_s}$, saturation moisture $\overline{q}_s(\overline{T}, p)$
- $N \rightarrow 0$ for $\overline{q}_c \rightarrow 0$ (or = 0 for $RH = 1$)

Xu-Randall's 1996 cloud fraction

XR full formula with given qc [g/kg], $qw=20$ [g/kg]



$$\left. \frac{\alpha \bar{q}_c}{1 - RH})^{\delta} \right\}$$

0.49

Xu-Randall's 1996 cloud fraction

STEADY STATE DIAGNOSTIC

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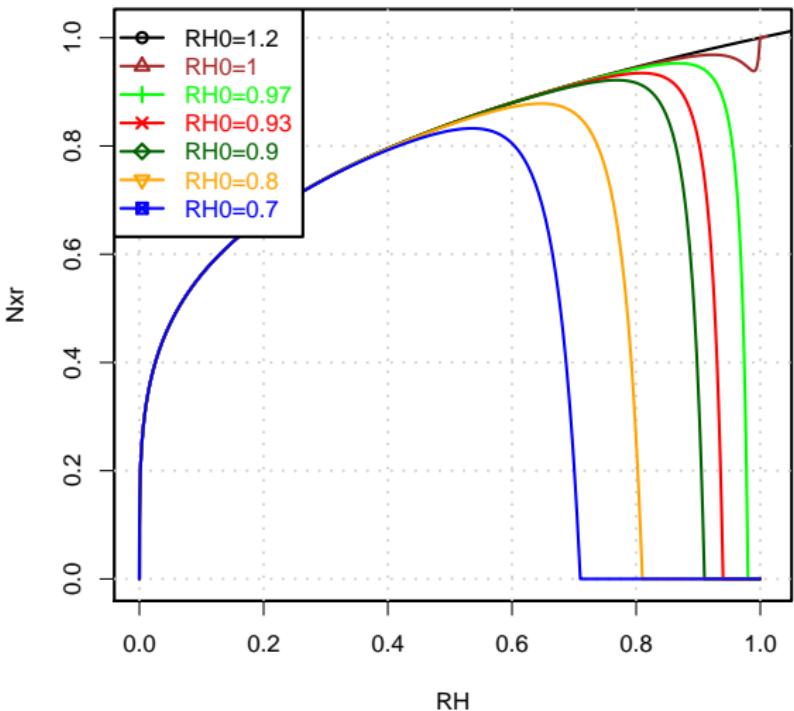
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microphysics should precipitate excess $\widehat{q}_c - \widehat{q}_{csusx}$

Xu-Randall's 1996 cloud fraction

XR formula limited by q_t , $qc_0=2e-04$, $qw=2e-02$



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- $RH = \frac{\overline{q}_v}{\overline{q}_w}$, mean grid-box wet bulb relative humidity
 (T_w, q_w) = attempt to avoid $q_s(T, p)$:
 - evaporate excess condensate, taking the heat from the air mass
 - bring oversaturated air mass to saturation, leaving the heat in the air

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 - bring oversaturated air mass to saturation, leaving the heat in the air*but blowing new air continuously*

Xu-Randall's 1996 cloud fraction

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- $RH = \frac{\overline{q}_v}{\overline{q}_w}$, mean grid-box wet bulb relative humidity
 (T_w, q_w) = attempt to avoid $q_s(T, p)$:
you cannot escape local diabatic effects:

$$L \cdot \delta q_v < \max(N, N_{xn}) \cdot c_p \left. \frac{\partial T}{\partial t} \right|_x \cdot \Delta t, \quad \text{allow around } 10K/day$$

Playing with Xu-Randall's 1996 formula

$$N = (RH)^r \left[1 - \exp \left\{ - \frac{\alpha \bar{q}_c}{(\bar{q}_w(1 - RH))^\delta} \right\} \right] \quad \text{re-defining } RH = \frac{\bar{q}_v}{\bar{q}_w}$$
$$\bar{q}_c = \min[\bar{q}_{cx}, q_t - RH \cdot \bar{q}_w], \quad \bar{q}_t = \bar{q}_w \cdot RH_0 + q_{c0}$$

- known: RH_0 , \bar{q}_w , \bar{q}_{c0} , hence \bar{q}_t
- 3 unknowns: N , RH , \bar{q}_c
- parameters: $\alpha \sim 100$, $r \sim 0.25$, $\delta \sim 0.49$
- q_{cx} to be parameterized, e.g. $\bar{q}_{cx} = \widehat{q_{cx0}} N^\gamma f(\bar{\omega})$

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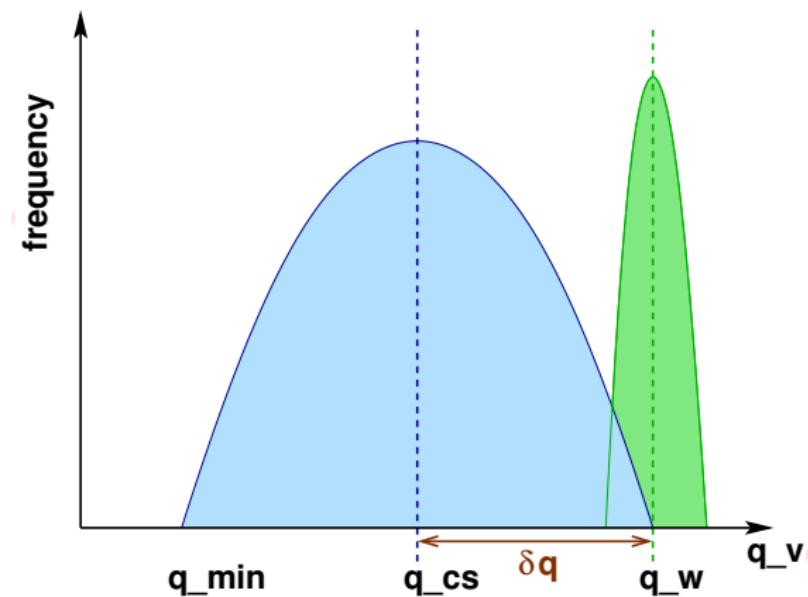
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⇒ We need one more relation between N and RH to close, e.g.

$$RH = N + (1 - N)RH_{cs}$$

i.e. $RH = 1$ in cloud and a mean RH_{cs} in clear sky.

Closure: subgrid variability assumption

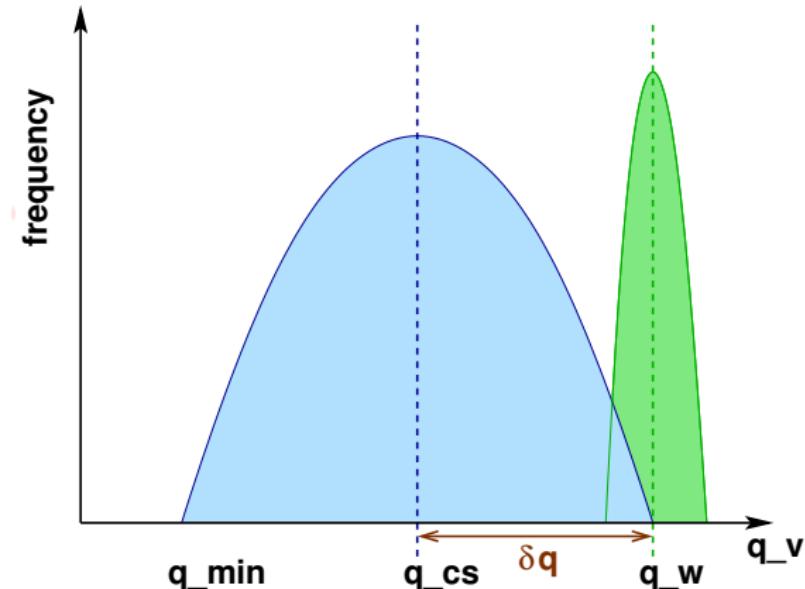


Closure: subgrid variability assumption

The range of RH in clear part ($1 - N$) may vary with its area (granularity)

$$\delta q = [(1 - N) \frac{A}{A_{\text{ref}}}]^p \frac{\overline{q_w} - q_{\min}}{2} \equiv \overline{q_w} (1 - N)^p (1 - H_0)$$

Larger p when more homogeneity (large TKE)

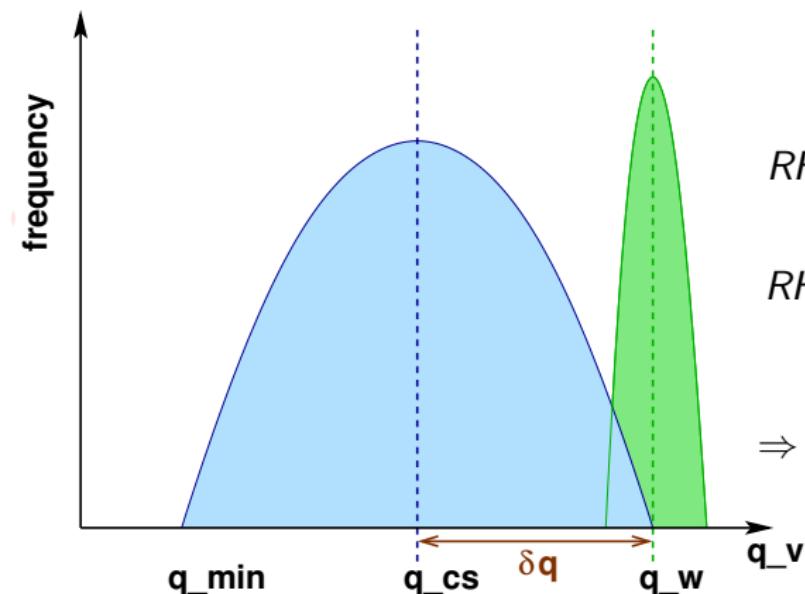


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Larger p when more homogeneity (large TKE)



$$q_{cs} = \bar{q}_w - \delta q$$

$$RH_{cs} = 1 - (1 - N)^p (1 - H_0)$$

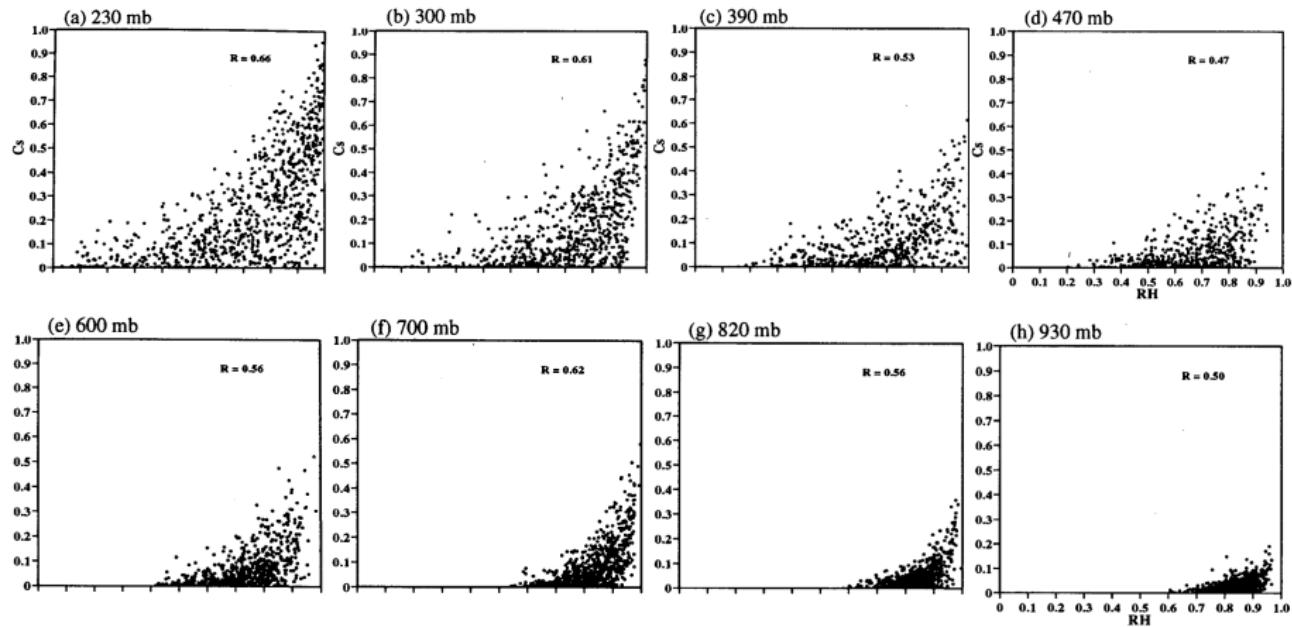
$$\begin{aligned} RH &= N + (1 - N) - RH_{cs} \\ &= 1 - (1 - N)^{p+1} (1 - H_0) \end{aligned}$$

$$\Rightarrow H_0 = \text{'critical' } RH \ (N \rightarrow 0)$$

Critical relative humidity ?

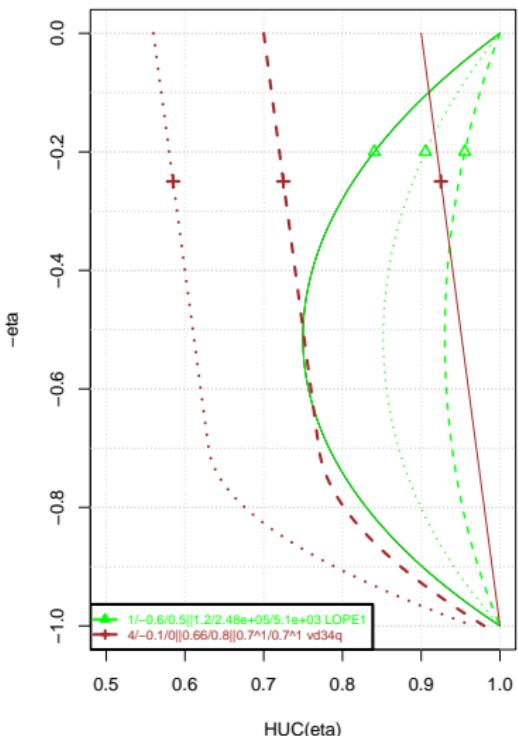
Xu-Randall 1997 Fig 5

Scatterplot: stratiform cloud amount vs large-scale RH (GATE 64km subdomain)



no unique RH_c but higher threshold and less scatter downwards

Profiles for subgrid variability



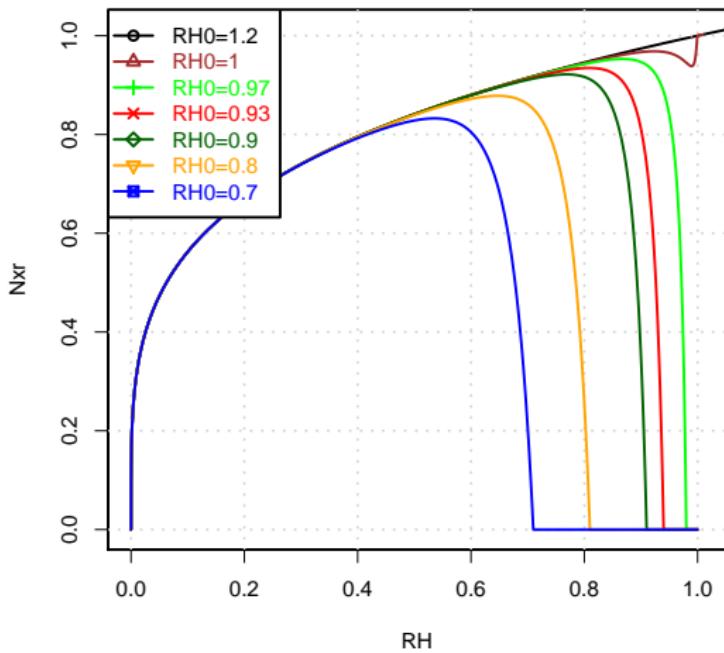
'Critical' relative humidity

- phase dependent
- increasing downwards
- was found necessary to return close to 1 at the surface

p exponent: increase towards surface

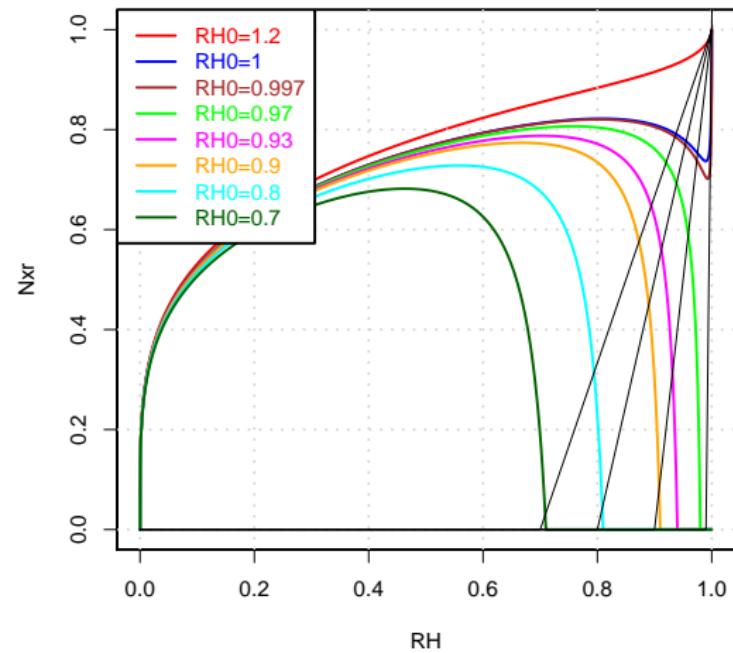
Problems of initial (Alaro-0) implementation

XR formula limited by q_t , $qc_0=2e-04$, $qw=2e-02$



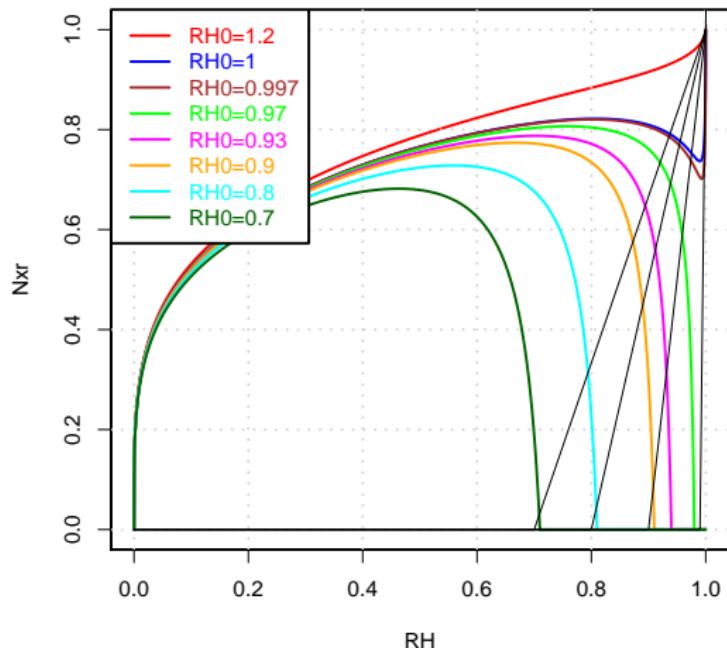
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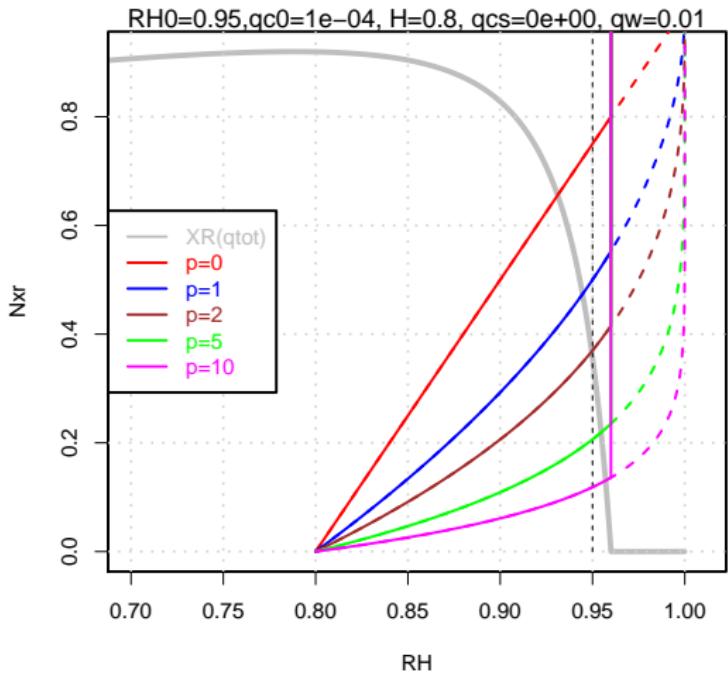
- simplified XR formula

$$N \approx RH^{\frac{1}{4}} \frac{\alpha \bar{q}_c}{\alpha \bar{q}_c + \sqrt{q_w(1-RH)}}$$

- $\bar{q}_c = \bar{q}_t - RH \cdot \bar{q}_w$
- $p = 0 \Rightarrow RH = N + (1 - N)H_0$
- H_0 set very high to limit condensation
⇒ imprecise N , binary.

Applying the limitations

$$N = (RH)^r \left[1 - \exp \left\{ - \frac{\alpha \overline{q}_c}{(\overline{q}_w(1 - RH))^\delta} \right\} \right]$$

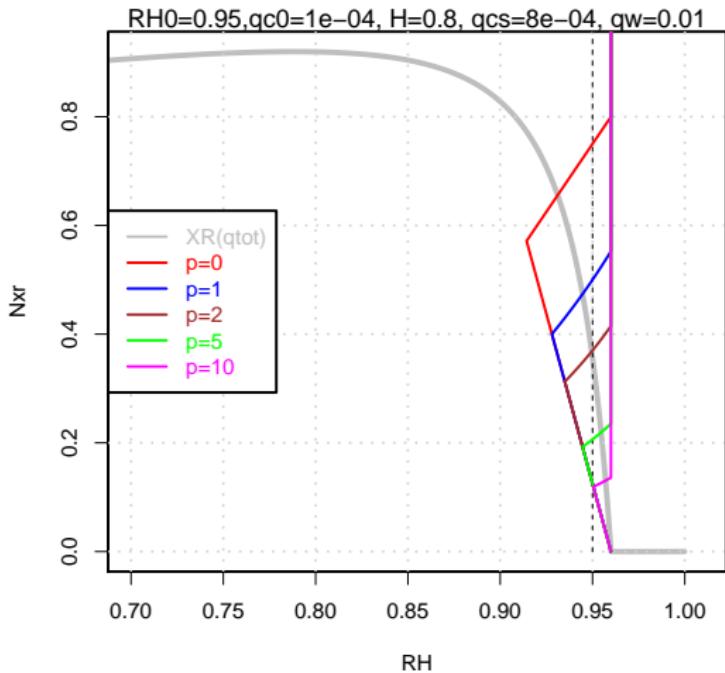


$$\overline{q}_c = \overline{q}_t - RH \overline{q}_w$$

$$RH = N + (1 - N)^{p+1}(1 - H_0)$$

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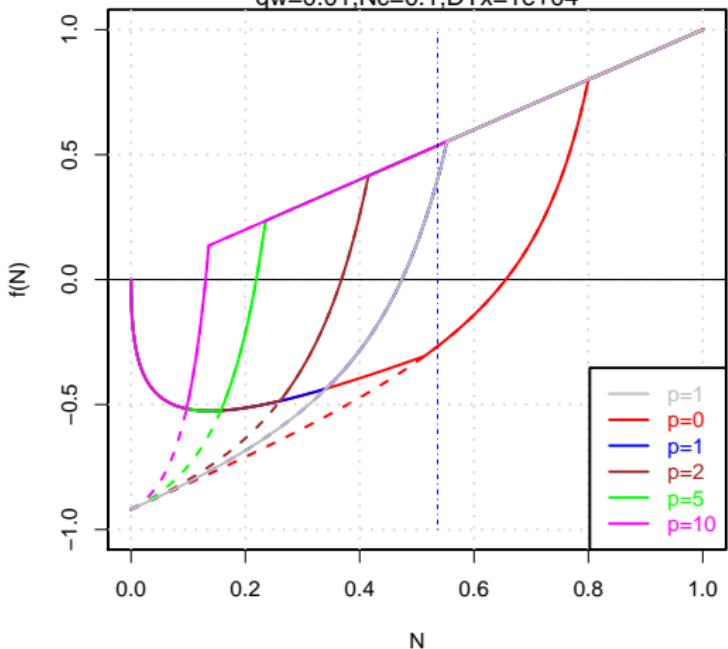
$$\overline{q}_c = \max[\overline{q}_{cx}, \overline{q}_t - RH\overline{q}_w]$$

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Applying the limitations

$$f(N) = N - (RH)^r \left[1 - \exp\left\{-\frac{\alpha \bar{q}_c}{(\bar{q}_w(1-RH))^\delta}\right\} \right]$$

RH0=0.95, qc0=1e-04, H=0.8, qcs=8e-04,
qw=0.01, Nc=0.1, DTx=1e+04



linear part: \bar{q}_t limitation

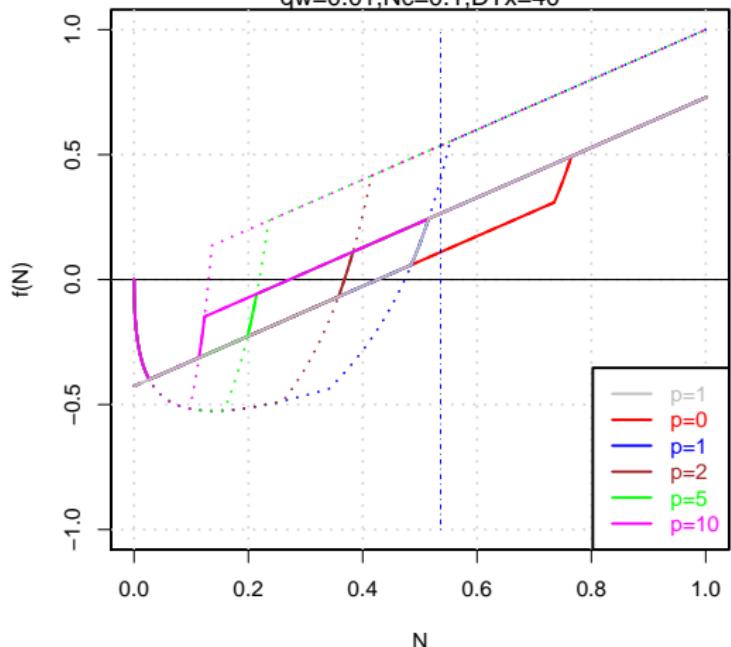
dashed: no \widehat{q}_{csus} limitation

solid: $\widehat{q}_{csus} = 0.8 \text{g/kg}$

Applying the limitations

$$f(N) = N - (RH)^r \left[1 - \exp\left\{-\frac{\alpha \bar{q}_c}{(\bar{q}_w(1-RH))^\delta}\right\} \right]$$

RH₀=0.95, q_{c0}=1e-04, H=0.8, q_{cs}=8e-04,
q_w=0.01, N_c=0.1, DT_x=40



protect convective condensate

$$RH < \frac{\bar{q}_{t0} - \bar{q}_{cc}}{\bar{q}_w}$$

limit cooling:

$$RH < RH_0 + \frac{c_p \Delta T_x}{L \bar{q}_w} g(N)$$

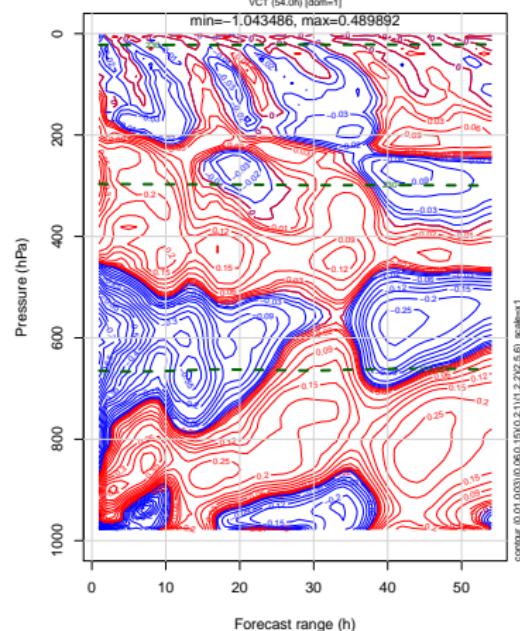
and heating:

$$RH > RH_0 - \frac{c_p \Delta T_x}{L \bar{q}_w} g(N)$$

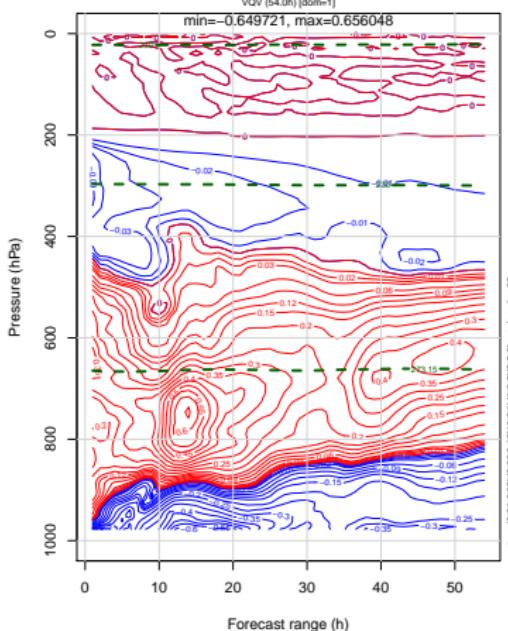
(shown $\Delta T_x = 40K/d$, $g(N) \equiv 1$)

Progress ? DDH tendencies not yet close enough...

vd34r5_0629-od42t1_0629, 2009-06-29 00:00

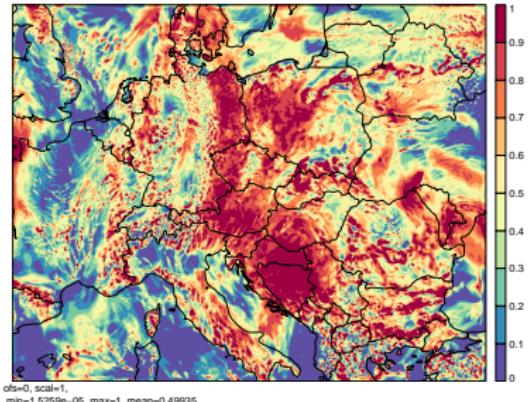


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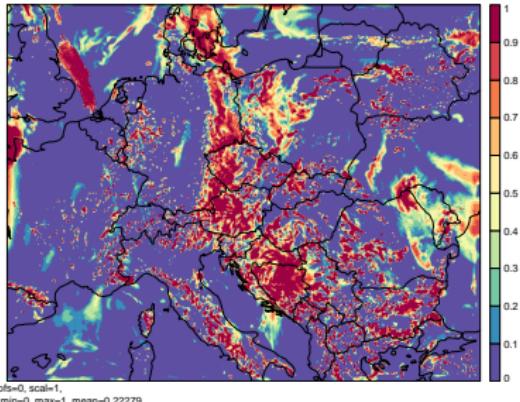


small $\Delta T_x \Rightarrow$ long spinup starting from null condensates
but larger ΔT_x may lead to excessive condensation, precipitation and evaporation.

od42t1 : 2009/06/29 z00:00 +12h
SURFNEBUL.TOT

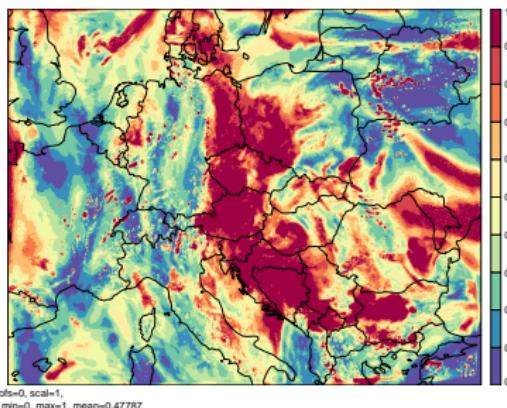


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SURFNEBUL.TOT



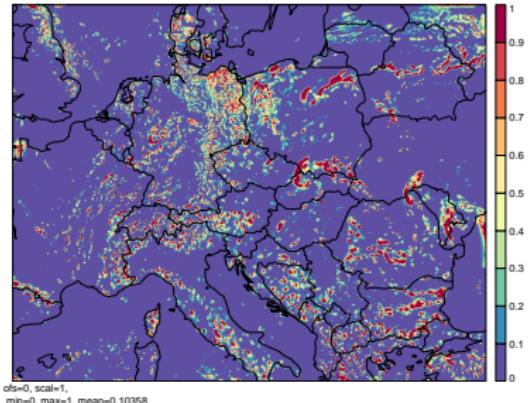
3MT Diag cloud

vd34r5 : 2009/06/29 z00:00 +12h
SURFNEBUL.TOT



3MT μ phys. cloud

od42t1 : 2009/06/29 z00:00 +12h
SURFNEBUL.CONV

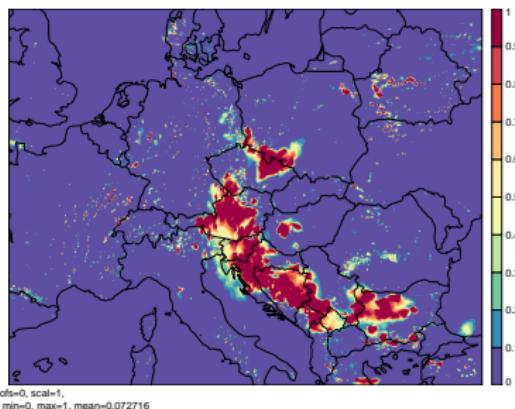


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SURFNEBUL.CONV



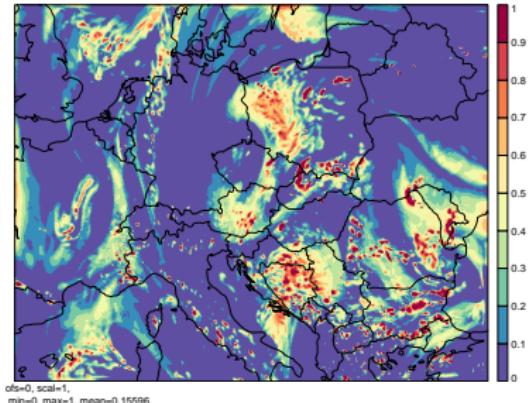
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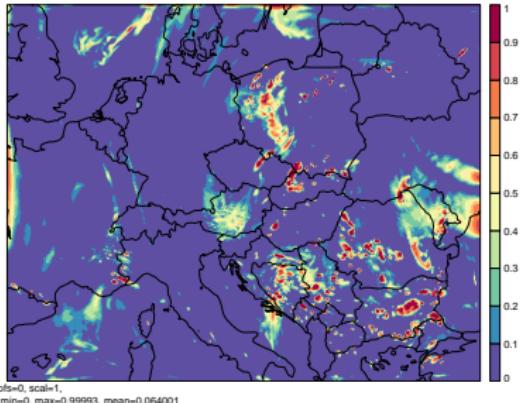


3MT μ phys. cloud

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SURFNEBUL.HAUT

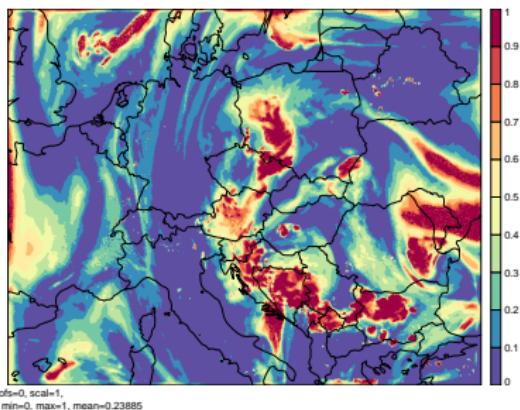


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SURFNEBUL.HAUT



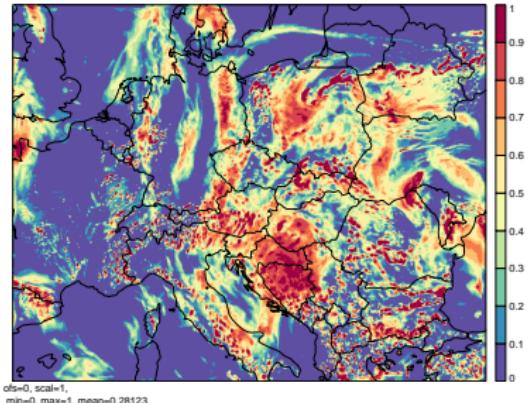
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SURFNEBUL.HAUT

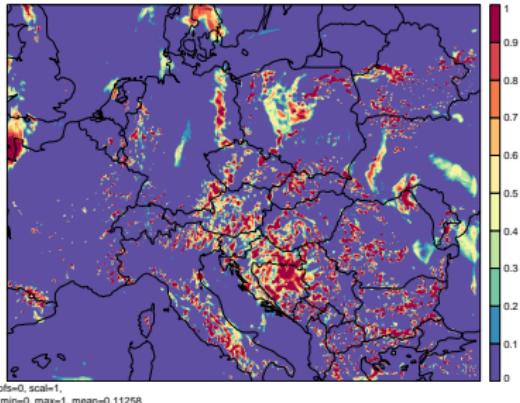


3MT μ phys. cloud

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SURFNEBUL.MOYE

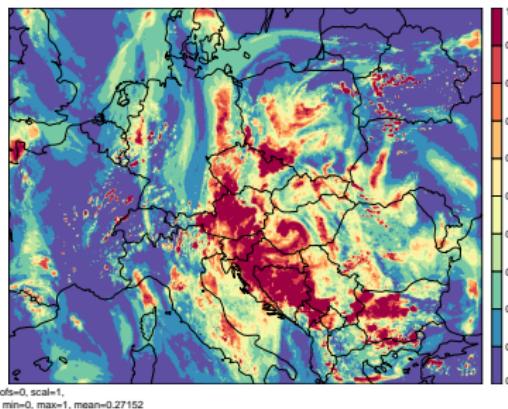


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SURFNEBUL.MOYE



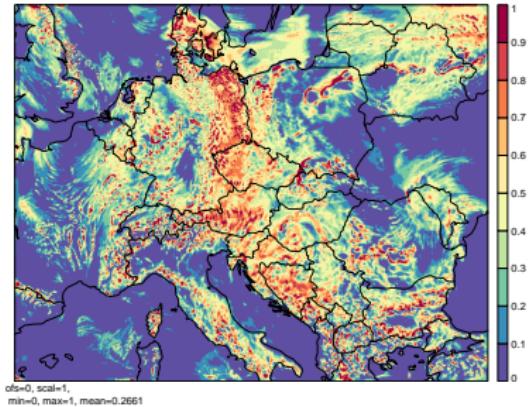
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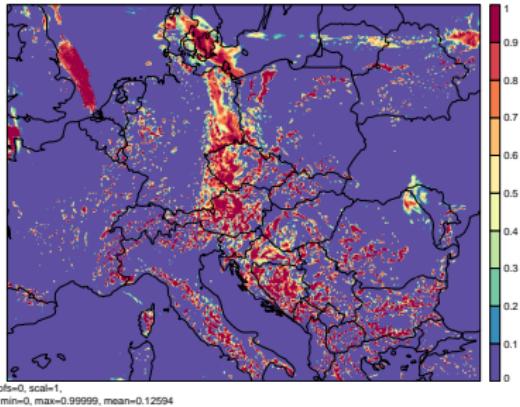


3MT μ phys. cloud

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SURFNEBUL.BASS

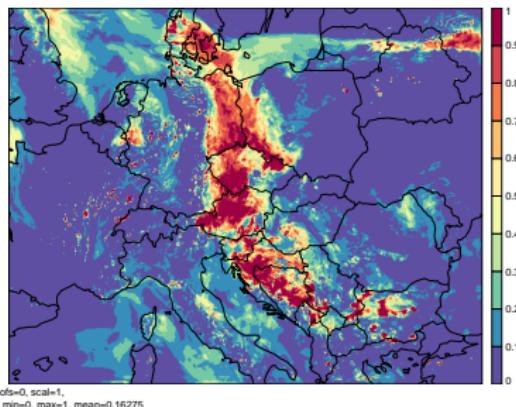


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SURFNEBUL.BASS



3MT Diag cloud

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SURFNEBUL.BASS



3MT μ phys. cloud

CSD+ unified cloud

And further on...

- Spinup problem, starting from zero condensates
- Tuning using CSD convection scheme (using initial 3MT as reference): means different split of condensation between convective and cloud cheme – test with 3MT planned.
- Tuning in parallel on winter and summer events.
- Control DDH components (T, q_v, q_c, q_r) and cloud aspect
- Multiple and complex feedbacks : e.g. more condensation \Rightarrow heating but also increased rr \Rightarrow more evaporative cooling and more downdraught activity, e.g. substantial surface cooling by subgrid transport.
- Most sensitive tunings are ΔT_x , H_0 profile, parameters of autoconversion.