Lab SUG 2020 Observatoire des Sciences de l'Univers



Isothermal Metamorphism of Snow:

Measurement of Interface Velocities and Phase-Field Modeling for a Better Understanding of the Involved Mechanisms







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Once deposited on the ground, snow forms a complex porous material whose microstructure constantly transforms over time. These evolutions, which strongly impact the physical and mechanical properties of snow (e.g. Srivastava et al, 2010; Löwe et al, 2013; Calonne et al, 2014; Wautier et al, 2015) need to be considered in details for an accurate snowpack modeling. However, some of the physical mechanisms involved in metamorphism are still poorly understood.

To address this problem, several investigations combining X-ray tomography and 3D micro-modeling have been carried out over the past decade (e.g. Flin et al, 2003; Kämpfer and Plapp, 2009; Pinzer et al, 2012) but precise comparisons between experimentation and modeling remain difficult. One of the difficulties comes from the lack of high resolution time-lapse series for experiments occurring with very well-defined boundary conditions, and from which precise measurements of the interfacial growth rates can be achieved.

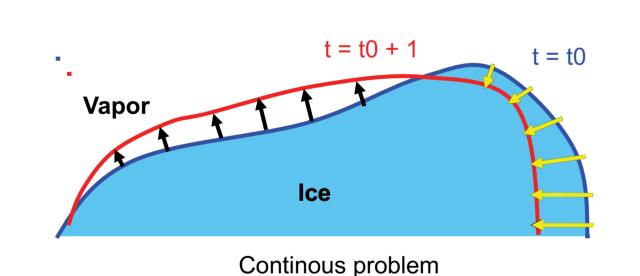
Using CellDyM, a recently developed cryogenic cell (Calonne et al, 2015), we conducted in situ time-lapse tomographic experiments on several snow and ice samples under various conditions. We focus here on a 28 h experiment of isothermal metamorphism at -7°C. The non-destructive nature of X-ray microtomography yielded a series of 8 micron resolution images that were acquired with time steps ranging from 2 to 12 hours. An image analysis method was then implemented to estimate the normal growth rates on each point of the ice-air interface and applied to the series obtained.

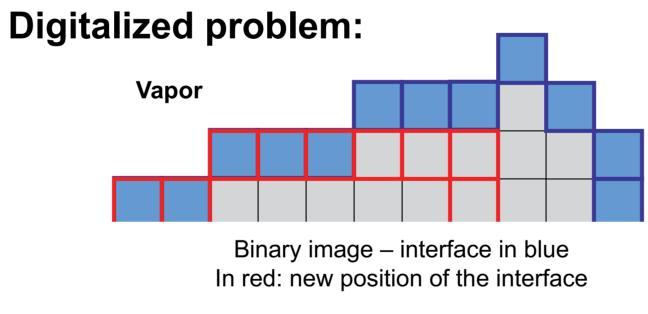
In addition, a curvature-based phase-field model (Bretin et al, 2015) was used on the first image of the experimental series to compute the snow evolution under a pure sublimation-condensation mechanism: the numerical morphologies were then compared to the experimental ones. The analysis of all the results gives interesting outlooks for the understanding of the physical mechanisms involved in snow isothermal metamorphism.

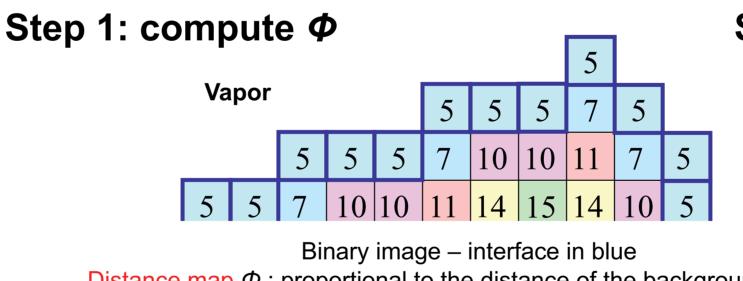
2. CellDyM: a Cryogenic Cell for Time-Lapse Tomography at Room-Temperature **Based on the following principles:** Metamorphism under isothermal conditions at -7°C -thermal regulation using 2 Peltier modules at top and base of the sample -thermal insulation from room temperature using a vacuum chamber -an amovible conductive sample holder that protects specimens during their installation into the cell Hot exchanger Water circulation Peltier module Cold exchanger Copper column t = 0 hour z: direction of gravity Vacuum chamber Temperature field simulated for equi-temperature 3D images with curvature fields Plexiglas cylinder Vertical cross-section (left) and temperature gradient (right) conditions Aluminum sample holder Snow sample Turbo +38 mm⁻¹ $C_{\text{mean}} = 6.3 \,\text{mm}^{-1}$ +38 mm⁻¹ Membrane molecular Sealing cap → Results: 3D time-lapse series with: - high spatial and temporal resolution Box with a dry air circulation - well-defined thermal conditions Enlarged views

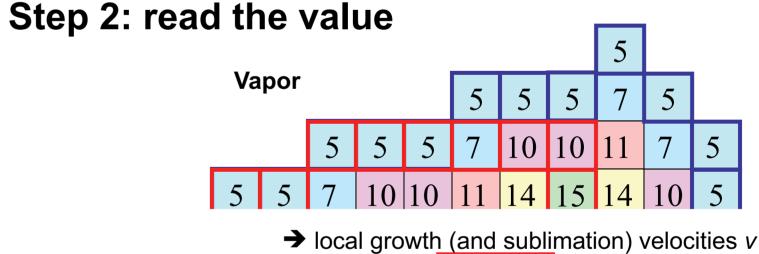
3. A Simple Method to Measure Normal Growth Rates from 3D Time-Lapse Images

The (signed) distance of the red interface from the blue one is directly given by the distance map Φ as follows: $d = (5 - \Phi) / 5$ (pixels)









Distance map Φ : proportional to the distance of the background (pores)

 \rightarrow local fluxes $j = v \cdot \rho_{ice}$ at the interface

From approach 1, we have:

 \rightarrow B = 1.79 x 10⁻¹⁹ m³/s

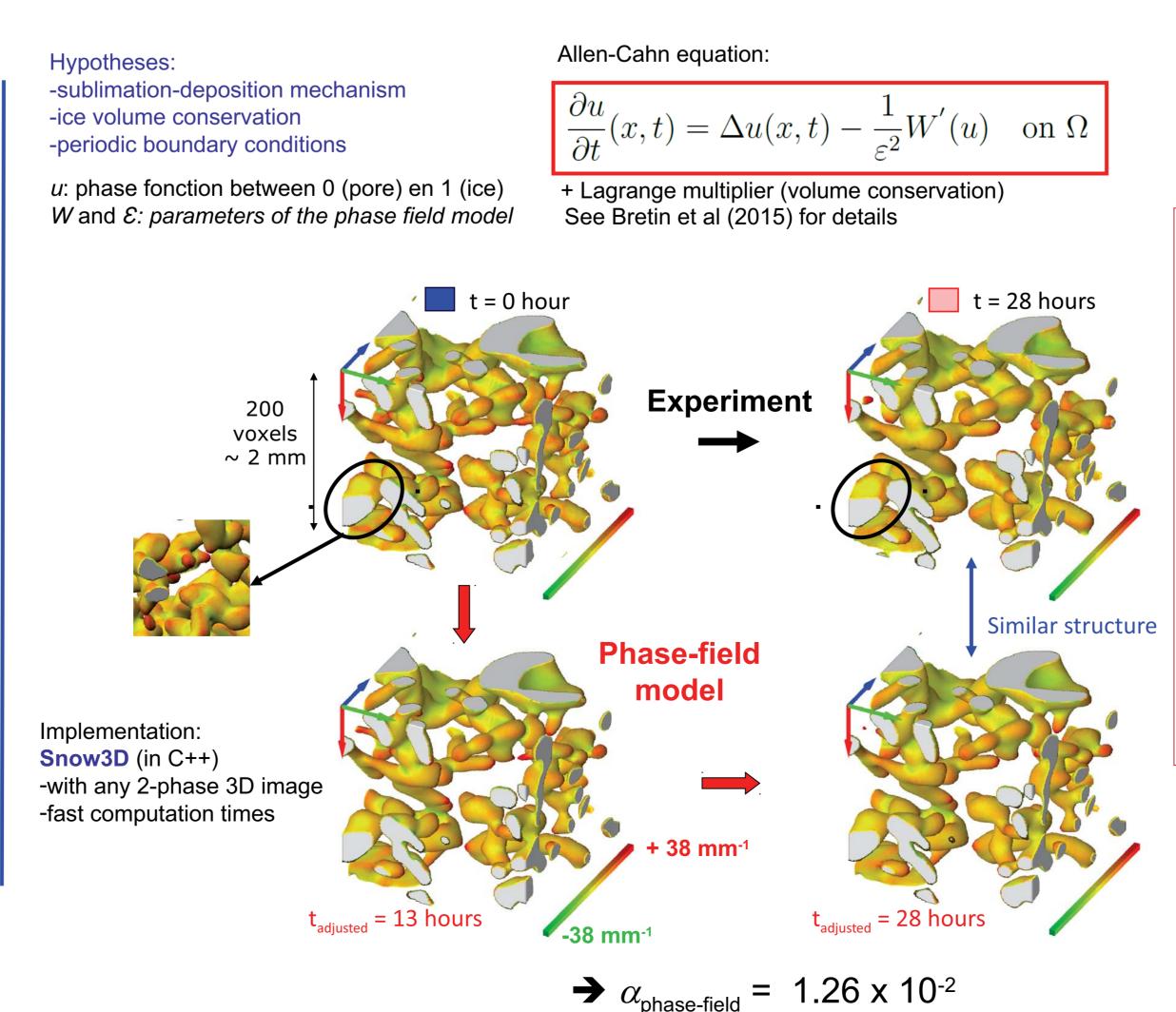
→ it is 3 times larger than the theo-

4. Preliminary Results: Is Sublimation-Deposition a Limiting Mechanism at -7°C? Which Value for $\alpha_{\rm eff}$?

Approach 2: Phase-field model

Approach 1: Growth rate measurements Interface velocity estimated between 0 and 28 h 131 voxels ~ 1 mm deposition: +5.10⁻¹⁰ m s⁻¹ sublimation: -5.10⁻¹⁰ m s⁻¹ $C_{\text{mean}} = 7.4 \, \text{mm}^{-1}$ Sublimation-deposition model (see e.g. Flin et al, 2003; Krol and Löwe, 2016): $v = \alpha k (C_{integral} - C)$ Mean curvature (mm⁻¹

Large scatter in the measurements + digitization effects Determination of α based on the slope of the linear regression: $\rightarrow \alpha_{\text{sublimation-deposition}} = 1.18 \times 10^{-2}$



with $\lambda = (2C^2 - K)^{-1/2}$ retical value at -7°C (5.92 x 10⁻²⁰ m³/s) **→** sublimation-deposition seems more likely than diffusion to be the limiting mechanism **Conclusions:** - We developed a room temperature operating cryogenic cell that provides images series of high spatial and temporal resolution and with well defined boundary conditions. - We realized an isothermal experiment at -7°C with this cell and proposed a numerical method to efficiently measure growth rates in 3D. - We performed growth rates measurements and concluded the sublimation-deposition seems predominant over the diffusion (at -7°C). It is in agreement with the results of Krol and Löwe (2016). - We confirmed these results using a phase-field model and found an effective accomodation coefficient about 10⁻², which is consistent (1) with the estimation given by our growth rate measurements and (2) with the typical values given in the literature (usually between 10⁻³ and 10⁻¹). **Outlooks:** - Confirm these results (with different samples, sizes, time-steps...) - Check other mechanisms (grain boundaries, surface diffusion, etc.)

Diffusion case?

 $v = B (C_{integral} - C) / \lambda$

- Use this approach with other experimental conditions (temperatures, etc.) Bretin, E., R. Denis, F. Flin, J.-O. Lachaud, E. Oudet & T. Roussillon (2015), Technical Report 4 of ANR DigitalSnow Project. Calonne, N., F. Flin, C. Geindreau, and others (2014), The Cryosphere, 8(6), 2255-2274, doi:10.5194/tc-8-2255-2014. Calonne, N., F. Flin, B. Lesaffre, A. Dufour, J. Roulle, P. Puglièse and others (2015), Geophys. Res. Lett., 42, doi: 10.1002/2015GL063541. Flin, F., J.-B. Brzoska, B. Lesaffre, and others (2003), J. Phys. D. Appl. Phys., 36, A49-A54, doi: 10.1088/0022-3727/36/10A/310. Flin, F. & J.-B. Brzoska (2008), Ann. Glaciol., 49, 17–21, doi:10.3189/172756408787814834. Kämpfer, T. U. & M. Plapp (2009), Phys. Rev. E, 79(3), 031502, doi:10.1103/PhysRevE.79.031502. Krol, Q. & H. Löwe (2016), J. Glaciol, 62(232), 378-390, doi: 10.1017/jog.2016.32 Löwe, H., F. Riche & M. Schneebeli (2013), The Cryosphere, 7(5), 1473-1480, doi:10.5194/tc-7-1473-2013. Pinzer, B., M. Schneebeli & T. Kämpfer (2012), The Cryosphere, 6, 1141–1155. Srivastava, P., P. Mahajan, P. Satyawali & V. Kumar (2010), Ann. Glaciol., 51, 73–82. Wautier, A., C. Geindreau & F. Flin, (2015), Geophys. Res. Lett., 42, 8031-8041, doi: 10.1002/2015GL065227.